



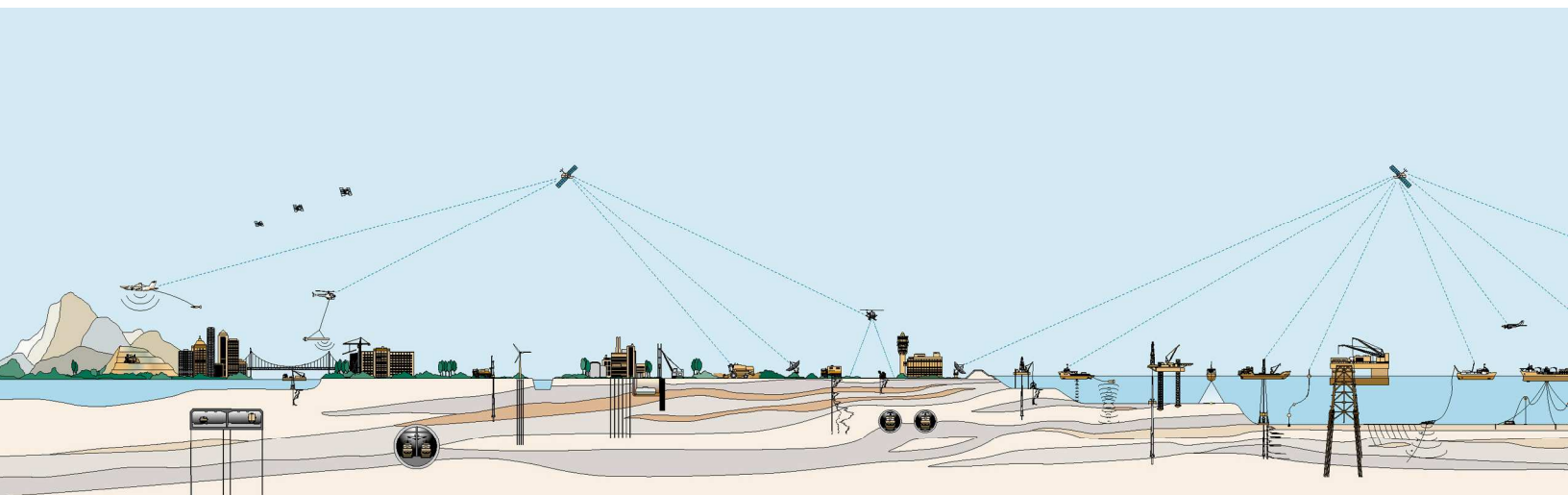
**Fugro Airborne Surveys**

**LOGISTICS AND PROCESSING REPORT  
GEOTEM<sup>®</sup> Airborne Electromagnetic and Magnetic Survey**

**RENABIE  
WAWA, ONTARIO**

**Project No. 13416**

**Conquest Resources Limited**



**LOGISTICS AND PROCESSING REPORT  
GEOTEM<sup>®</sup> AIRBORNE ELECTROMAGNETIC AND  
MAGNETIC SURVEY  
RENABIE  
WAWA, ONTARIO**

**PROJECT NO. 13416**

Client: Conquest Resources Limited  
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Toronto, Ontario  
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Date of Report: March, 2013

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# I

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## Introduction

On March 15, 2013, Fugro Airborne Surveys conducted a GEOTEM<sup>®</sup> airborne electromagnetic and magnetic survey of the Renabie area on behalf of Conquest Resources Limited. Using Wawa, Ontario as the base of operations, a total of 175 line kilometres of data were collected using a Casa 212 modified aircraft (Figure 1).

The survey data were processed and compiled in the Fugro Airborne Surveys Ottawa office. The collected and processed data are presented on colour maps and multi-parameter profiles. The following maps were produced: Residual Magnetic Intensity (RMI), First Vertical Derivative of RMI, Apparent Conductance derived from dB/dt Z-Coil Channel 01, Decay Constant (Tau) derived from B-Field X-Coil Channels 12 to 30, Total Energy Envelope derived from dB/dt Channel 08 and Flight Path with Anomaly Selection. In addition, digital archives of the raw and processed survey data in line format, and gridded EM data were delivered.



Figure 1: Specially modified Casa 212 aircraft used by Fugro Airborne Surveys.

# II

## Survey Operations

### Location of the Survey Area

The Renabie area (Figure 2) was flown with Wawa, Ontario as the base of operations. A total of 35 traverse lines were flown 4 km in length, with a spacing of 200 m between lines, and 5 tie lines were flown with a spacing of 1000 m between tie lines totalling 175 km for the complete survey.

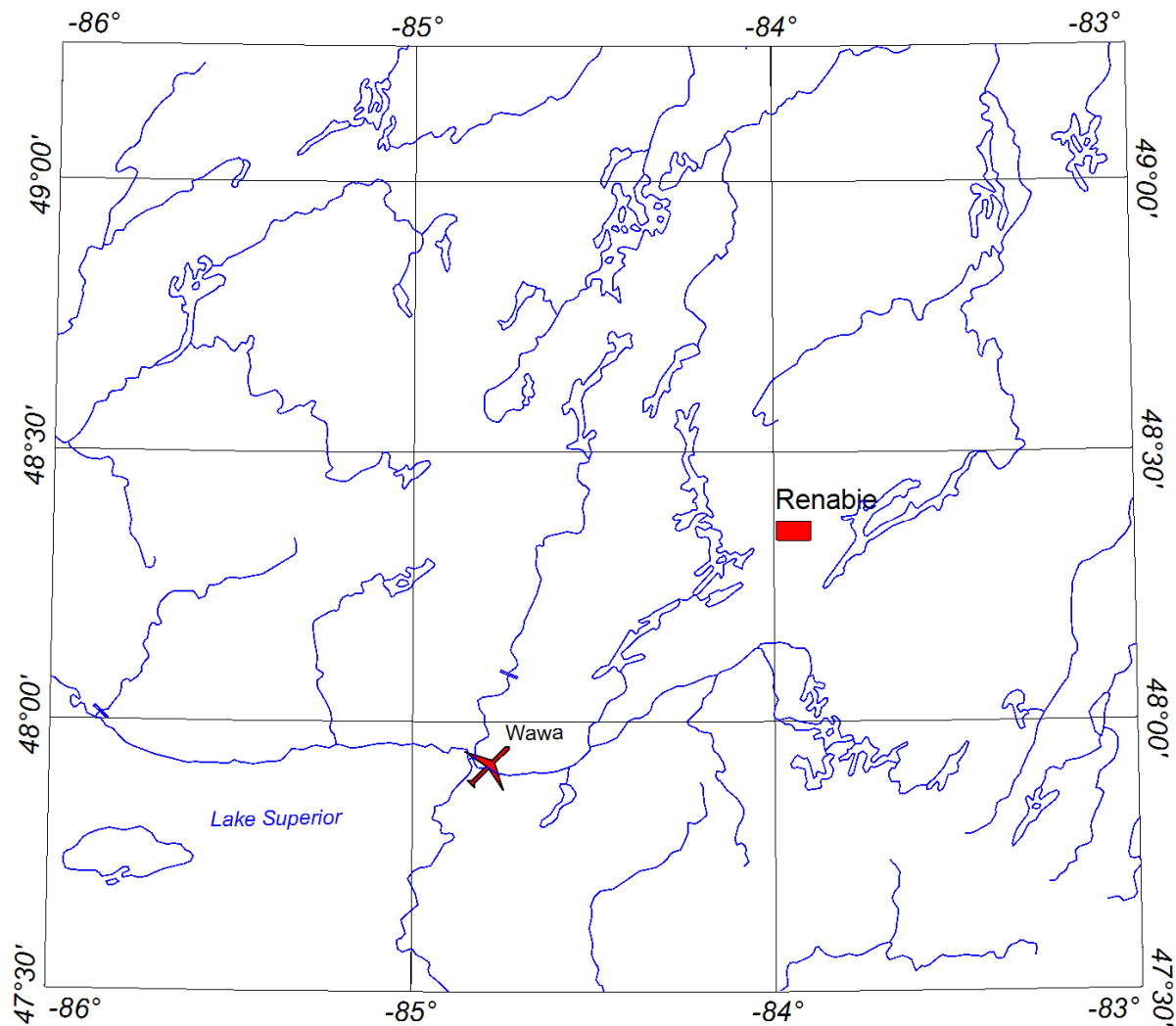


Figure 2: Survey location.

## **Aircraft and Geophysical On-Board Equipment**

Aircraft: Casa 212 (Twin Turbo Propeller)

Operator: FUGRO AIRBORNE SURVEYS

Registration: C-GDPP

Survey Speed: 125 knots / 145 mph / 65 m/s

Magnetometer: Scintrex CS-2 single cell caesium vapour, towed-bird installation, sensitivity =  $0.01 \text{ nT}^1$ , sampling rate = 0.1 s, ambient range 20,000 to 100,000 nT. The general noise envelope was kept below 0.5 nT. The nominal sensor height was ~75 m above ground.

Electromagnetic system: GEOTEM® 30 channel Multi-coil System

Transmitter: Vertical axis loop mounted on aircraft of  $231 \text{ m}^2$

Number of turns 6

Nominal height above ground of 120 m

Receiver: Multi-coil system (x, y and z) with a final recording rate of 4 samples per second, for the recording of 30 channels of x, y and z-coil data. The nominal height above ground is ~75 m, placed ~130 m behind the centre of the transmitter loop.

Base frequency: 90 Hz

Pulse delay:  $22 \mu\text{s}$

Pulse width:  $2127 \mu\text{s}$

Off-time:  $3407 \mu\text{s}$

Point value:  $10.85 \mu\text{s}$

Transmitter Current: 530 A

Dipole moment:  $0.73 \times 10^6 \text{ Am}^2$



Figure 3: Mag and GEOTEM® Receivers



Figure 4: Modified Casa 212 in flight.

<sup>1</sup> One nanotesla (nT) is the S.I. equivalent of one gamma.

Table 1: Electromagnetic Data Windows.

Channel	Start (p)	End (p)	Width (p)	Start (ms)	End (ms)	Width (ms)	Mid (ms)
1	3	13	11	0.022	0.141	0.119	0.081
2	14	73	60	0.141	0.792	0.651	0.467
3	74	132	59	0.792	1.432	0.640	1.112
4	133	192	60	1.432	2.083	0.651	1.758
5	193	203	11	2.083	2.203	0.119	2.143
6	204	205	2	2.203	2.224	0.022	2.214
7	206	207	2	2.224	2.246	0.022	2.235
8	208	209	2	2.246	2.268	0.022	2.257
9	210	211	2	2.268	2.289	0.022	2.279
10	212	214	3	2.289	2.322	0.033	2.306
11	215	217	3	2.322	2.355	0.033	2.338
12	218	220	3	2.355	2.387	0.033	2.371
13	221	224	4	2.387	2.431	0.043	2.409
14	225	228	4	2.431	2.474	0.043	2.452
15	229	233	5	2.474	2.528	0.054	2.501
16	234	239	6	2.528	2.593	0.065	2.561
17	240	246	7	2.593	2.669	0.076	2.631
18	247	254	8	2.669	2.756	0.087	2.713
19	255	264	10	2.756	2.865	0.109	2.810
20	265	276	12	2.865	2.995	0.130	2.930
21	277	290	14	2.995	3.147	0.152	3.071
22	291	306	16	3.147	3.320	0.174	3.234
23	307	324	18	3.320	3.516	0.195	3.418
24	325	344	20	3.516	3.733	0.217	3.624
25	345	366	22	3.733	3.971	0.239	3.852
26	367	390	24	3.971	4.232	0.260	4.102
27	391	416	26	4.232	4.514	0.282	4.373
28	417	444	28	4.514	4.818	0.304	4.666
29	445	474	30	4.818	5.143	0.326	4.980
30	475	512	38	5.143	5.556	0.412	5.349



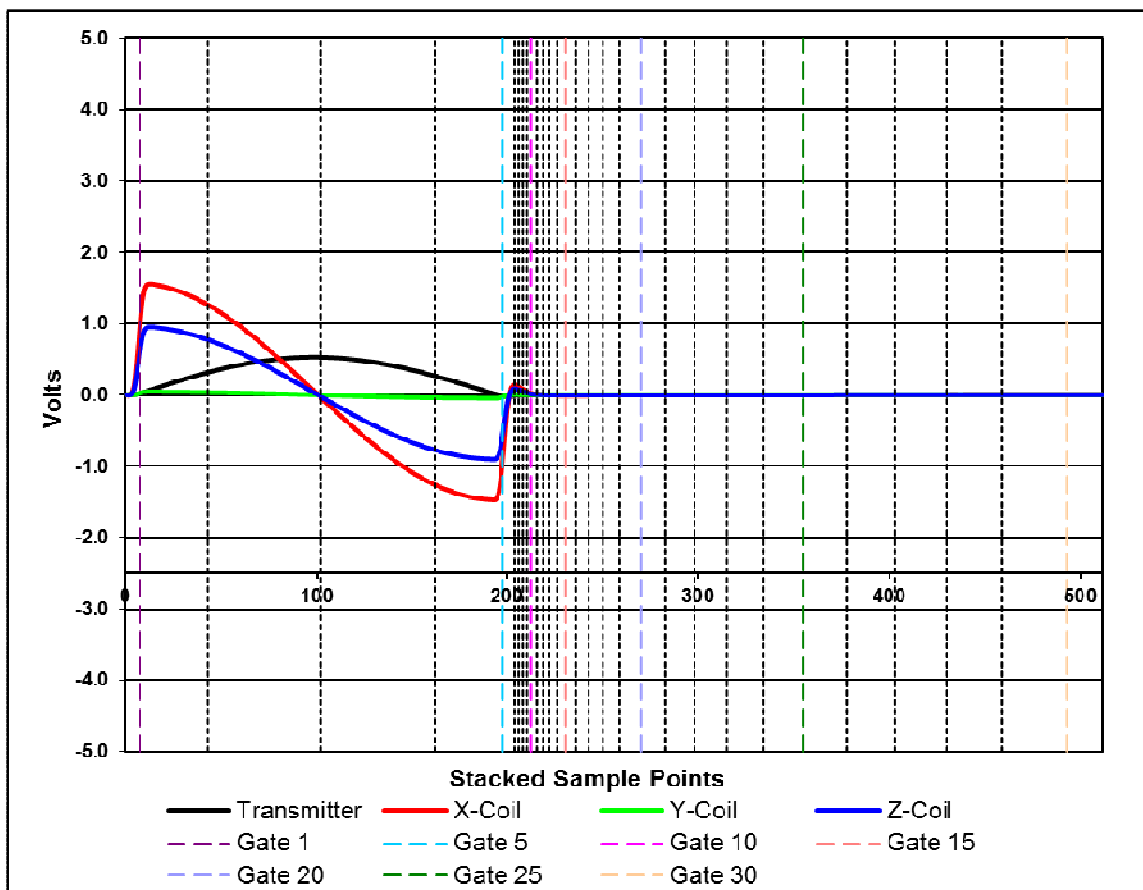


Figure 5: GEOTEM® Waveform and response with gate centres showing positions in sample points.

Digital Acquisition:	FUGRO AIRBORNE SURVEYS GEODAS SYSTEM.
Barometric Altimeter:	Rosemount 1241M, sensitivity 1 ft, 0.5 sec recording interval.
Radar Altimeter:	King, accuracy 2%, sensitivity 1 ft, range 0 to 2500 ft, 0.5 sec recording interval.
Camera:	Digital video
Electronic Navigation:	NovAtel OEM4, 1 sec recording interval, with a resolution of 0.00001 degree and an accuracy of $\pm 5m$ .

### **Base Station Equipment**

Magnetometer:	Natural Resources Canada's Ottawa Geomagnetic Observatory
GPS Receiver:	NovAtel OEM4, measuring all GPS channels, for up to 12 satellites.
Computer:	Laptop, Pentium model or better.
Data Logger:	CF1, SBBS (single board base station).

### **Field Office Equipment**

Computer:	Dell Laptop.
Printer:	Portable printer.
DVD writer Drive:	Internal DVD+RW format.
Hard Drives:	Internal hard drive + two external hard drives for redundant backups.

### **Survey Specifications**

Traverse Line Direction:	004° - 184°
Traverse Line Spacing:	200 m
Tie Line direction:	094° - 274°
Tie Line spacing:	1000 m
Navigation:	Real-time differential GPS. Traverse and tie line spacing was not to exceed the nominal by > 50% for more than 3 km.
Altitude:	The survey was flown at a mean terrain clearance of 120 m. Altitude was not to exceed 130 m over 3 km.
Magnetic Noise Levels:	The noise envelope on the magnetic data was not to exceed $\pm 0.25$ nT over 3 km.
EM Noise Levels:	The noise envelope on the last raw electromagnetic dB/dt X- and Z-coil channel was not to exceed $\pm 3500$ pT/s over a distance greater than 3 km.

### **Personnel**

Data Processor:	E. Rooen, D. Wilson, K. Zawadzki, L. Lafrenière
Pilots:	D. Wiens, T. Gaillot
Electronics Operator:	F. Beillard

### **Production Statistics**

Flying dates:	March 15 <sup>th</sup> , 2013
Total production:	175 line kilometres
Number of production flights:	1
Days lost weather:	none

# III

## Quality Control and Compilation Procedures

Important checks were performed during the data acquisition stage to ensure that the data quality was in keeping with the survey specifications. The following outlines the Quality Control measures conducted throughout the acquisition phase of the survey.

### Initial Field QC

At the completion of each day's flying an initial review of the data was performed in the field. This process was primarily to ensure all the equipment was functioning properly and enables the crew to immediately ascertain that production can resume the following day. This process does not necessarily determine if the data were within specifications. Priority was given to getting the data back to the office where a more thorough analysis of the data was performed. A list of the steps of the initial field review of the data follows:

- 1) All digital files were confirmed to be readable and free of defects.
- 2) The integrity of the airborne electromagnetic and magnetometer data was checked through statistical analysis and graphically viewed in profile form. Any null values or unreasonable noise levels were identified.
- 3) All altimeter and positional data were checked for any inconsistency, invalid values and spikes.
- 4) The base station files were examined for validity and continuity. The data extent was confirmed to cover the entire acquisition period.
- 5) The diurnal data were examined for any noise events or spiking.
- 6) Flight path video files were visually checked for quality and to confirm the full coverage for the survey flight.
- 7) Duplicate backups of all digital files were created.

### Transmission of Data from Field to Office

At the completion of each day's flying the raw data was uploaded to a secure FTP site. This enabled the office processing staff to immediately conduct more thorough data quality checks and start the processing with a minimum duplication of procedures or loss of time. This also enabled the direct supervision and involvement by senior processors and the availability of a greater depth of knowledge to be applied to any problems with the minimum of delay.

# IV

## Data Processing

### **Flight Path Recovery**

GPS Recovery:	Real-time differentially corrected GPS.
Projection:	Universal Transverse Mercator (UTM Zone 17N)
Datum:	NAD83
Central meridian:	81° West
False Easting:	500000 metres
False Northing:	0 metres
Scale factor:	0.9996

### **Altitude Data**

Noise editing:	Alfatrim median filter used to eliminate the highest and lowest values from the two highest and lowest values from a 9 point sample window for the radar and barometric altimeters before applying a 3 point averaging filter.
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### **Base Station Diurnal Magnetics**

Data properties:	Data from Natural Resources of Canada's Ottawa Geomagnetic Observatory were obtained at a recording interval of 60 seconds.
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### **Airborne Magnetics**

Lag correction:	1.5 s
Noise editing:	4th difference editing routine set to remove spikes greater than 0.5 nT.
Noise filtering:	Triangular filter set to remove noise events having a wavelength less than 0.9 seconds.
Diurnal subtraction:	The diurnal was removed from the data with a base value of 54690 nT added back.
IGRF model:	2010
IGRF removal date:	2013/03/15
Gridding:	The data were gridded using a bi-directional routine with a grid cell size of 50 m.

## **Residual Magnetic Intensity**

The residual magnetic intensity (RMI) was calculated from the total magnetic intensity (TMI), the diurnal, and the regional magnetic field. The TMI was measured in the aircraft, the diurnal was measured from the Ottawa Observatory and the regional magnetic field was calculated from the International Geomagnetic Reference Field (IGRF). The low frequency component of the diurnal was extracted from the filtered ground station data and removed from the TMI. The average of the diurnal was then added back in to obtain the resultant TMI. The regional magnetic field, calculated for the specific survey location and the time of the survey, was removed from the resultant TMI to obtain the RMI. The final step was to Tie line level and microlevel the RMI data.

## **Magnetic First Vertical Derivative**

The first vertical derivative was calculated in the frequency domain from the final grid values to enhance subtleties related to geological structures.

A first vertical derivative has also been displayed in profile form. This was calculated from the line data by combining the transfer functions of the 1st vertical derivative and a low-pass filter (cut-off wavelength = 2.5 seconds, roll-off wavelength = 4.5 seconds). The low-pass filter was designed to attenuate the high frequencies representing non-geological signal, which are normally enhanced by the derivative operator. This parameter is also stored in the final digital archive.

## **Electromagnetics**

### **dB/dt data**

Lag correction: 2.25 s

Data correction: The x, y and z-coil data were processed from the 30 raw channels recorded at 4 samples per second.

The following processing steps were applied to the dB/dt data from all coil sets:

- a) The data were edited for residual spheric spikes by examining the decay pattern of each individual EM transient. Bad decays (i.e. not fitting a normal exponential function) were deleted and replaced by interpolation;
- b) Noise filtering was done using an adaptive filter technique based on time domain triangular operators. Using a 2nd difference value to identify changes in gradient along each channel, minimal filtering (3 point convolution) is applied over the peaks of the anomalies, ranging in set increments up to a maximum amount of filtering in the resistive background areas (31 points for both the x-coil and the z-coil data);
- c) The data from channels 1 to 5 (on-time) and 6 to 30 (off-time) were corrected for drift in flight form (prior to cutting the recorded data back to the correct line limits) by passing a low order polynomial function through the baseline minima along each channel, via a graphic screen display;
- d) The filtered data from the x, y and z-coils were then re-sampled to a rate of 5 samples per second and combined into a common file for archiving.

## B-field data

**Processing steps:** The processing of the B-Field data stream is very similar to the processing for the regular dB/dt data. The lag adjustment used was the same, followed by:

- 1) Spike editing for spheric events;
- 2) Correction for coherent noise. By nature, the B-Field data will contain a higher degree of coherency of the noise that automatically gets eliminated (or considerably attenuated) in the regular dB/dt, since this is the time derivative of the signal;
- 3) Final noise filtering with an adaptive filter;
- 4) Drift adjustments.

**Note:** The introduction of the B-Field data stream, as part of the GEOTEM® system, provides the explorationist with a more effective tool for exploration in a broader range of geological environments and for a larger class of target priorities.

The advantage of the B-Field data compared with the normal voltage data (dB/dt) are as follows:

1. A broader range of target conductance that the system is sensitive to. (The B-Field is sensitive to bodies with conductance as great as 100,000 siemens);
2. Enhancement of the slowly decaying response of good conductors;
3. Suppression of rapidly decaying response of less conductive overburden;
4. Reduction in the effect of spherics on the data;
5. An enhanced ability to interpret anomalies due to conductors below thick conductive overburden;
6. Reduced dynamic range of the measured response (easier data processing and display).

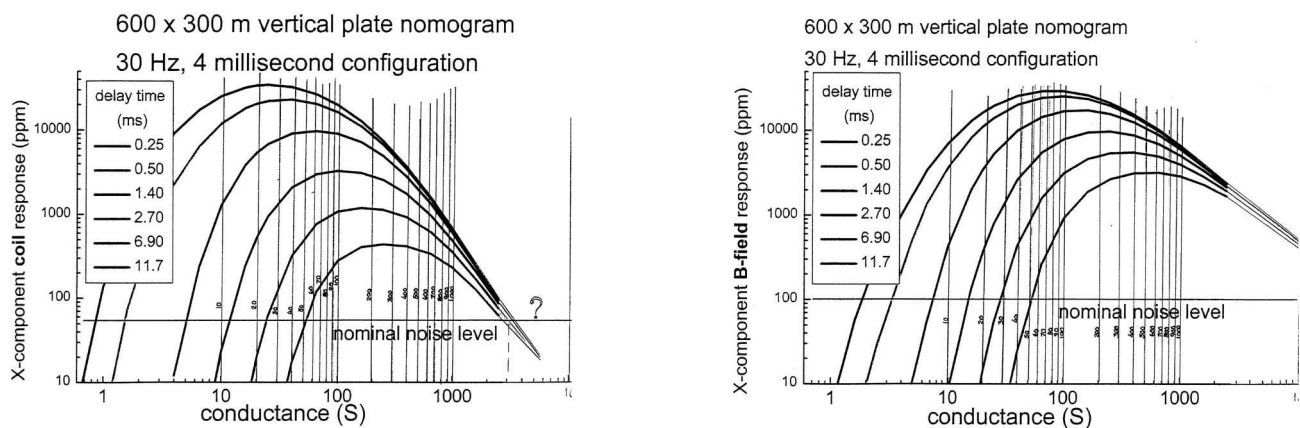


Figure 6: dB/dt vertical plate nomogram (left), B-field vertical plate nomogram (right).

Figure 6 displays the calculated vertical plate response for the GEOTEM<sup>®</sup> signal for the dB/dt and B-Field. For the dB/dt response, you will note that the amplitude of the early channel peaks at about 25 siemens, and the late channels at about 250 siemens. As the conductance exceeds 1000 siemens the response curves quickly roll back into the noise level. For the B-Field response, the early channel amplitude peaks at about 80 siemens and the late channel at about 550 siemens. The projected extension of the graph in the direction of increasing conductance, where the response would roll back into the noise level, would be close to 100,000 siemens. Thus, a strong conductor, having a conductance of several thousand siemens, would be difficult to interpret on the dB/dt data, since the response would be mixed in with the background noise. However, this strong conductor would stand out clearly on the B-Field data, although it would have an unusual character, being a moderate to high amplitude response, exhibiting almost no decay.

In theory, the response from a super conductor (50,000 to 100,000 siemens) would be seen on the B-Field data as a low amplitude, non-decaying anomaly, not visible in the off-time channels of the dB/dt stream. Caution must be exercised here, as this signature can also reflect a residual noise event in the B-Field data. In this situation, careful examination of the dB/dt on-time (in-pulse) data is required to resolve the ambiguity. If the feature were strictly a noise event, it would not be present in the dB/dt off-time data stream. This would locate the response at the resistive limit, and the mid in-pulse channel (normally identified as channel 3) would reflect little but background noise, or at best a weak negative peak. If, on the other hand, the feature does indeed reflect a superconductor, then this would locate the response at the inductive limit. In this situation, channel 3 of the dB/dt stream will be a mirror image of the transmitted pulse, i.e. a large negative.

### Coil Oscillation Correction

The electromagnetic receiver sensor is housed in a bird, which is towed behind the aircraft using a cable. Any changes in airspeed of the aircraft, variable crosswinds, or other turbulence will result in the bird swinging from side to side. This can result in the induction sensors inside the bird rotating about their mean orientation. The rotation is most marked when the air is particularly turbulent. The changes in orientation result in variable coupling of the induction coils to the primary and secondary fields. For example, if the sensor that is normally aligned to measure the x-axis response pitches upward, it will be measuring a response that will include a mixture of the X and Z-component responses. The effect of coil oscillation on the data increases as the signal from the ground (conductivity) increases and may not be noticeable when flying over areas which are generally resistive. This becomes more of a concern when flying over highly conductive ground.

Using the changes in the coupling of the primary field, it is possible to estimate the pitch, roll and yaw of the receiver sensors. In the estimation process, it is assumed that a smoothed version of the primary field represents the primary field that would be measured when the sensors are in the mean orientation. The orientations are estimated using a non-linear inversion procedure, so erroneous orientations are sometimes obtained. These are reviewed and edited to insure smoothly varying values of orientations. These orientations can then be used to unmix the measured data to generate a response that would be measured if the sensors were in the correct orientation. For more information on this procedure please refer to *Advances in Time-Domain EM Technology* published in Proceedings of Exploration 97; Fourth Decennial International Conference on Mineral Exploration by R. Smith and P. Annan.

For the present dataset, the data from all 30 channels of dB/dt and B-Field parameters have been corrected for coil oscillation.



## Total Energy Envelope

To combine the benefits of the measurement from both the X and Z-coil data and reduce the asymmetry in the signature of the anomalies, the channel data from both the X and Z-coils can be used to compute the Total Energy Envelope of the response. This is done through a Hilbert Transform and essentially reflects the square root of the sum of the squares of each component.

For the present dataset, channel 08 of the dB/dt or component (mid-time position of 108 µsec after turn-off) was selected as the optimum window along the transient to map the response over features of interest. This channel is sufficiently late in time to avoid minor variations in the signal due to overburden conductivity and altitude effects and yet, not so late in time that the possible weak response over features of interest would get lost in the background

## Decay Constant (TAU)

The decay constant values are obtained by fitting the channel data from either the complete off-time signal of the decay transient or only a selected portion of it (as defined by specific channels) to a single exponential of the form

$$Y = A e^{-t/\tau}$$

where **A** is amplitude at time zero, **t** is time in microseconds and **τ** is the decay constant, expressed in microseconds. A semi-log plot of this exponential function will be displayed as a straight line, the slope of which will reflect the rate of decay and therefore the strength of the conductivity. A slow rate of decay, reflecting a high conductivity, will be represented by a high decay constant.

As a single parameter, the decay constant provides more useful information than the amplitude data of any given single channel, as it indicates not only the peak position of the response but also the relative strength of the conductor. It also allows better discrimination of conductive axes within a broad formational group of conductors.

For the present dataset, the decay constant was calculated by fitting the X-coil response from channels 12 to 30 (mean delay times of 222 to 3200 µsec after turn-off) of the B-Field component to the exponential function.

## Apparent Conductance

Fugro has developed an algorithm that converts the response in any measurement window (on or off-time) into an apparent or conductance. This is performed using a look-up table that contains the response at a range of half-space conductivities and altimeter heights.

The apparent conductance for the present dataset was calculated using dB/dt Z-coil channel 01 to provide the maximum information on the near-surface conductivity of the ground which, when combined with the magnetic signature, provides good geological mapping.

# V

## Final Products

### **Digital Archives**

Line and grid data in the form of an ASCII text file (\*.xyz), Geosoft database (\*.gdb), Geosoft grids (\*.grd), geotiff grids (\*.tif) and KMZ files (\*.kmz) have been written to DVD. The formats and layouts of these archives are further described in Appendix E (Data Archive Description). Hardcopies of all maps have been created as outlined below.

### **Maps**

Scale: 1:20,000  
Parameters: Residual Magnetic Intensity  
First Vertical Derivative of the Residual Magnetic Intensity  
Decay Constant (Tau) Derived from B-Field X-Coil Channels 12 to 30  
Apparent Conductance Derived from dB/dt Z-Coil Channel 01  
Total Energy Envelope Derived from dB/dt X and Z-Coil Channel 08  
Media/Copies: 1 paper & 1 digital (Geosoft, Geotiff, and PDF format)

### **Profile Plots**

Scale: 1:20,000  
Parameters: Multi-channel presentation with 23 channels of both dB/dt and B-field X and Z-coil, Residual Magnetic Intensity, Calculated Magnetic Vertical Gradient, GPS Altitude, EM Primary Field, Hz Monitor and Terrain.  
Media/Copies: 1 Digital of Each Line (PNG format)

### **Report**

Media/Copies: 2 paper & 1 digital (PDF format)

# **Appendix A**

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## **Fixed-Wing Airborne Electromagnetic Systems**

## **FIXED-WING AIRBORNE ELECTROMAGNETIC SYSTEMS**

### **General**

The operation of a towed-bird time-domain electromagnetic system (EM) involves the measurement of decaying secondary electromagnetic fields induced in the ground by a series of short current pulses generated from an aircraft-mounted transmitter. Variations in the decay characteristics of the secondary field (sampled and displayed as windows) are analyzed and interpreted to provide information about the subsurface geology. The response of such a system utilizing a vertical-axis transmitter dipole and a multi-component receiver coil has been documented by various authors including Smith and Keating (1991, *Geophysics* v.61, p. 74-81).

A number of factors combine to give the fixed-wing platforms excellent signal-to-noise ratio and depth of penetration: 1) the principle of sampling the induced secondary field in the absence of the primary field (during the "off-time"), 2) the large separation of the receiver coils from the transmitter, 3) the large dipole moment and 4) the power available from the fixed wing platform. Such a system is also relatively free of noise due to air turbulence. However, also sampling in the "on-time" can result in excellent sensitivity for mapping very resistive features and very conductive features, and thus mapping the geology (Annan et al., 1991, *Geophysics* v.61, p. 93-99). The on-time and off-time parts of the half-cycle waveform are shown in Figure 1.

Through free-air model studies using the University of Toronto's Plate and Layered Earth programs it may be shown that the "depth of investigation" depends upon the geometry of the target. Typical depth limits would be 400 m below surface for a homogeneous half-space, 550 m for a flat-lying inductively thin sheet or 300 m for a large vertical plate conductor. These depth estimates are based on the assumptions that the overlying or surrounding material is resistive.

The method also offers very good discrimination of conductor geometry. This ability to distinguish between flat-lying and vertical conductors combined with excellent depth penetration results in good differentiation of bedrock conductors from surficial conductors (Appendix C).

### **Methodology**

The Fugro time-domain fixed-wing electromagnetic systems (GEOTEM® and MEGATEM®) incorporate a high-speed digital EM receiver. The primary electromagnetic pulses are created by a series of discontinuous sinusoidal current pulses fed into a three- or six-turn transmitting loop surrounding the aircraft and fixed to the nose, tail and wing tips. The base frequency rate is selectable: 25, 30, 75, 90, 125, 150, 225 and 270 Hz. The length of the pulse can be tailored to suit the targets. Standard pulse widths available are 0.6, 1.0, 2.0 and 4.0 ms. The available off-time can be selected to be as great as 16 ms. The dipole moment depends on the pulse width, base frequency and aircraft used on the survey. Example pulse widths and off-time windows at different base frequencies are shown on Figure 2. The specific dipole moment, waveform and gate settings for this survey are given in the main body of the report.

The receiver is a three-axis (x, y, z) induction coil. In the fixed-wing systems, this is towed by the aircraft on a 135-metre cable. The tow cable is non-magnetic, to reduce noise levels. The usual mean terrain clearance for the aircraft is 120 m with the EM bird being situated nominally 50 m below and 130 m behind the aircraft (see Figure 3).

Each primary pulse causes decaying eddy currents in the ground to produce a secondary magnetic field. This secondary magnetic field, in turn, induces a voltage in the receiver coils, which is the

electromagnetic response. Good conductors decay slowly, while poor conductors more rapidly (see Figure 1).

The measured signals pass through anti-aliasing filters and are then digitized with an A/D converter at sampling rates of up to 80 kHz. The digital data flows from the A/D converter into an industrial-grade computer where the data are processed to reduce the noise.

Operations, which are carried out in the receiver, are:

1. *Primary-field removal:* In addition to measuring the secondary response from the ground, the receiver sensor coils also measure the primary response from the transmitter. During flight, the bird position and orientation changes slightly, and this has a very strong effect on the magnitude of the total response (primary plus secondary) measured at the receiver coils. The variable primary field response is distracting because it is unrelated to the ground response. The primary field can be measured by flying at an altitude such that no ground response is measurable. These calibration signals are used to define the shape of the primary waveform. By definition this primary field includes the response of the current in the transmitter loop plus the response of any slowly decaying eddy currents induced in the aircraft. We assume that the shape of the primary will be unchanged as the bird position changes, but that the amplitude will vary. The primary-field-removal procedure involves solving for the amplitude of the primary field in the measured response and removing this from the total response to leave a secondary response. Note that this procedure removes any ("in-phase") response from the ground that has the same shape as the primary field. For more details on the primary-field removal procedure please refer to *On removing the primary field from fixed-wing time-domain airborne electromagnetic data: some consequences for quantitative modeling, estimating bird position and detecting perfect conductors* published in *Geophysical Prospecting*, 2001, Vol. 49; 405-416 by R. Smith.
2. *Digital Stacking:* Stacking is carried out to reduce the effect of broadband noise on the data.
3. *Windowing of data:* The digital receiver samples the secondary and primary electromagnetic field at 512, 1024, or 2048 points per EM pulse and windows the signal in up to 20 time gates whose centres and widths are software selectable and which may be placed anywhere within or outside the transmitter pulse. This flexibility offers the advantage of arranging the gates to suit the goals of a particular survey, ensuring that the signal is appropriately sampled through its entire dynamic range. Example off-time windows are shown on Figure 1.
4. *Power Line Filtering:* Digital comb filters are applied to the data during real-time processing to remove power line interference while leaving the EM signal undisturbed. The RMS power line voltage (at all harmonics in the receiver passband) are computed, displayed and recorded for each data stack.
5. *Primary Field:* The primary field at the towed sensor is measured for each stack and recorded as a separate data channel to assess the variation in coupling between the transmitter and the towed sensor induced by changes in system geometry.
6. *Earth Field Monitor:* A monitor of sensor coil motion noise induced by coil motion in the Earth's magnetic field is also extracted in the course of the real-time digital processing. This information is also displayed on the real-time chart as well as being recorded for post-survey diagnostic processes.

7. *Noise/Performance:* A monitor computes the RMS signal level on an early off-time window over a running 10-second window. This monitor provides a measure of noise levels in areas of low ground response. This information is printed at regular intervals on the side of the flight record and is recorded for every data stack.

One of the major roles of the digital receiver is to provide diagnostic information on system functions and to allow for identification of noise events, such as spherics, which may be selectively removed from the EM signal. The high digital sampling rate yields maximum resolution of the secondary field.

### **System Hardware**

The airborne EM system consists of the aircraft, the on-board hardware, and the software packages controlling the hardware. The software packages in the data acquisition system and in the EM receiver were developed in-house, as were, certain elements of the hardware (transmitter, system timing clock, towed-bird sensor system).

### **Transmitter System**

The transmitter system drives high-current pulses of an appropriate shape and duration through the coils mounted on the aircraft.

### **System Timing Clock**

This subsystem provides appropriate timing signals to the transmitter, and also to the analog-to-digital converter, in order to produce output pulses and capture the ground response. All systems are synchronized to GPS time.

### **Towed-Bird Systems**

A three-axis induction coil sensor is mounted inside a towed bird, which is typically 50 metres below and 130 metres behind the aircraft. (A second bird, housing the magnetometer sensor, is typically 50 metres below and 80 metres behind the aircraft.)

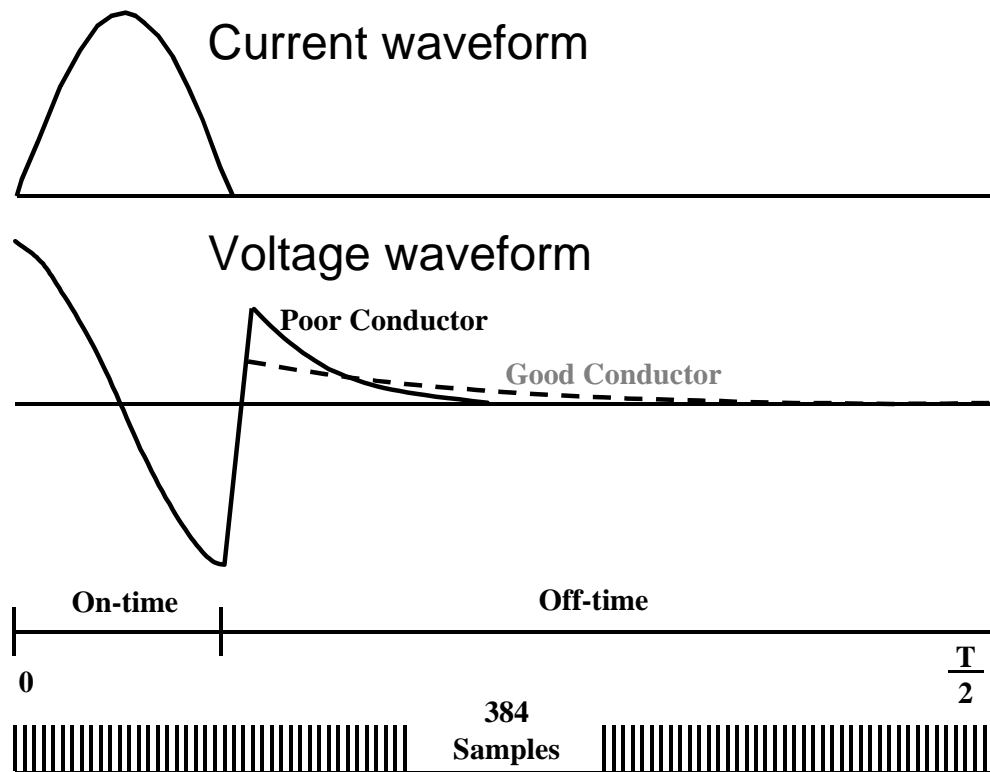


Figure 1. The waveforms and data sampling throughout the transmitter on and off-time.

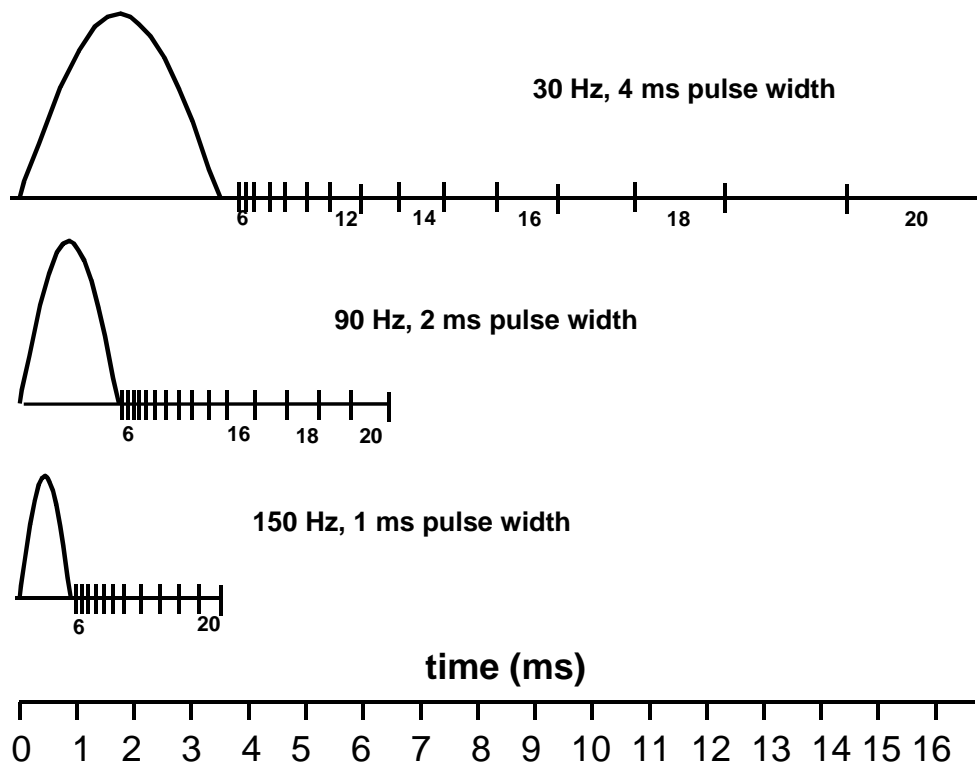


Figure 2. Pulse width and measurement windows for 150, 90 and 30 Hz base frequencies.

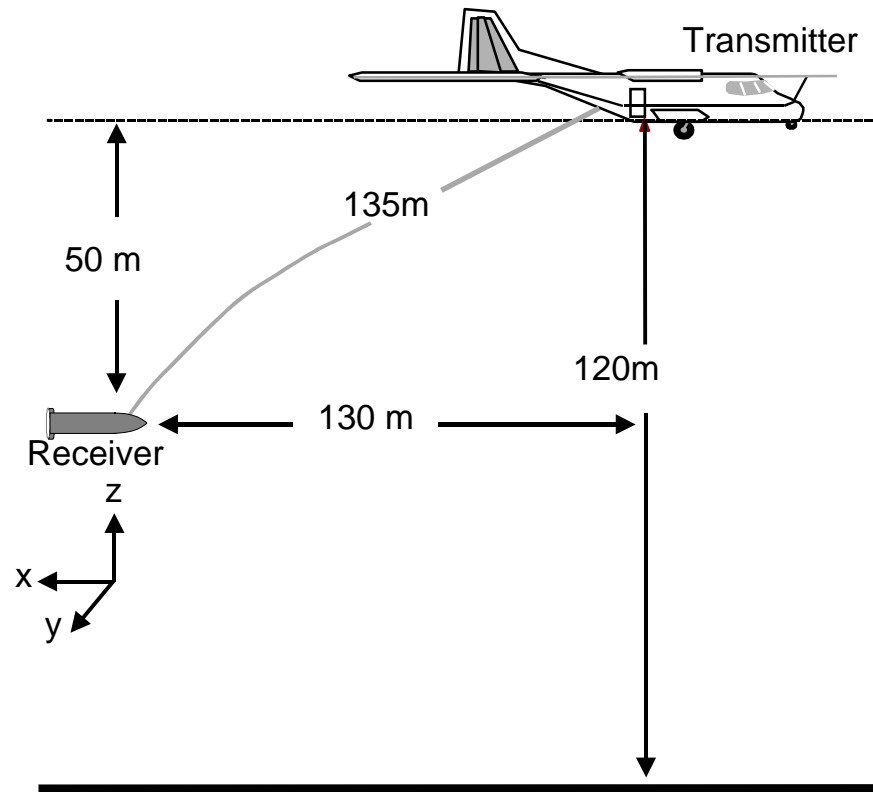


Figure 3. Nominal geometry of the fixed-wing electromagnetic system.



## **Appendix B**

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### Airborne Transient EM Interpretation

## Interpretation of transient electromagnetic data

### Introduction

The basis of the transient electromagnetic (EM) geophysical surveying technique relies on the premise that changes in the primary EM field produced in the transmitting loop will result in eddy currents being generated in any conductors in the ground. The eddy currents then decay to produce a secondary EM field that may be sensed in the receiver coil.

MEGATEM® and GEOTEM® are airborne transient (or time-domain) towed-bird EM systems incorporating a high-speed digital receiver which records the secondary field response with a high degree of accuracy. Most often the earth's total magnetic field is recorded concurrently.

Although the approach to interpretation varies from one survey to another depending on the type of data presentation, objectives and local conditions, the following generalizations may provide the reader with some helpful background information.

The main purpose of the interpretation is to determine the probable origin of the responses detected during the survey and to suggest recommendations for further exploration. This is possible through an objective analysis of all characteristics of the different types of responses and associated magnetic anomalies, if any. If possible the airborne results are compared to other available data. Certitude is seldom reached, but a high probability is achieved in identifying the causes in most cases. One of the most difficult problems is usually the differentiation between surface conductor responses and bedrock conductor responses.

### Types of Conductors

#### Bedrock Conductors

The different types of bedrock conductors normally encountered are the following:

1. Graphites. Graphitic horizons (including a large variety of carbonaceous rocks) occur in sedimentary formations of the Precambrian as well as in volcanic tuffs, often concentrated in shear zones. They correspond generally to long, multiple conductors lying in parallel bands. They have no magnetic expression unless associated with pyrrhotite or magnetite. Their conductivity is variable but generally high.
2. Massive sulphides. Massive sulphide deposits usually manifest themselves as short conductors of high conductivity, often with a coincident magnetic anomaly. Some massive sulphides, however, are not magnetic, others are not very conductive (discontinuous mineralization or sphalerite), and some may be located among formational conductors so that one must not be too rigid in applying the selection criteria.

In addition, there are syngenetic sulphides whose conductive pattern may be similar to that of graphitic horizons but these are generally not as prevalent as graphites.

3. Magnetite and some serpentized ultrabasics. These rocks are conductive and very magnetic.
4. Manganese oxides. This mineralization may give rise to a weak EM response.

## Surficial Conductors

1. Beds of clay and alluvium, some swamps, and brackish ground water are usually poorly conductive to moderately conductive.
2. Lateritic formations, residual soils and the weathered layer of the bedrock may cause surface anomalous zones, the conductivity of which is generally low to medium but can occasionally be high. Their presence is often related to the underlying bedrock.

## Cultural Conductors (Man-Made)

3. Power lines. These frequently, but not always, produce a conductive type of response. In the case when the power line comb filter does not remove the radiated field, the anomalous response can exhibit phase changes between different windows. In the case of current induced by the EM system in a grounded wire, or steel pylon, the anomaly may look very much like a bedrock conductor.
4. Grounded fences or pipelines. These will invariably produce responses much like a bedrock conductor. Whenever they cannot be identified positively, a ground check is recommended.
5. General culture. Other localized sources such as certain buildings, bridges, irrigation systems, tailings ponds etc., may produce EM anomalies. Their instances, however, are rare and often they can be identified on the visual path recovery system.

## Analysis of the Conductors

The conductance of a plate is generally estimated assuming the plate is vertical and 600 m by 300 m. Hence the conductance alone is not generally a decisive criterion in the analysis of a conductor. In particular, one should note:

- Its shape and size,
- All local variations of characteristics within a conductive zone,
- Any associated geophysical parameter (e.g. magnetics),
- The geological environment,
- The structural context, and
- The pattern of surrounding conductors.

The first objective of the interpretation is to classify each conductive zone according to one of the three categories which best defines its probable origin. The categories are cultural, surficial and bedrock. A second objective is to assign to each zone a priority rating as to its potential as an economic prospect.

## Bedrock Conductors

This category comprises those anomalies that cannot be classified according to the criteria established for cultural and surficial responses. It is difficult to assign a universal set of values that typify bedrock conductivity because any individual zone or anomaly might exhibit some, but not all, of these values and still be a bedrock conductor. The following criteria are considered indicative of a bedrock conductor:

1. An intermediate to high conductivity identified by a response with slow decay, with an anomalous response present in the later windows.
2. For vertical conductors, the anomaly should be narrow, relatively symmetrical, with a well-defined x-component peak.
3. If the conductor is thin, the response should show the characteristics evident in Figures 2 to 4. These figures illustrate how the response varies as a function of the flight direction for three bodies with different dips. The alternating character of the response as a result of line direction can be diagnostic of conductor geometry.
4. A small to intermediate amplitude. Large amplitudes are normally associated with surficial conductors. The amplitude varies according to the depth of the source.
5. A degree of continuity of the EM characteristics across several lines.
6. An associated magnetic response of similar dimensions. One should note, however, that those magnetic rocks that weather to produce a conductive upper layer would possess this magnetic association. In the absence of one or more of the characteristics defined in 1, 2, 3, 4 and 5, the related magnetic response cannot be considered significant.

Most obvious bedrock conductors occur in long, relatively monotonous, sometimes multiple zones following formational strike. Graphitic material is usually the most probable source. Massive syngenetic sulphides extending for many kilometres are known in nature but, in general, they are not common. Long formational structures associated with a strong magnetic expression may be indicative of banded iron formations.

In summary, a bedrock conductor reflecting the presence of a massive sulphide would normally exhibit the following characteristics:

- A high conductivity,
- A good anomaly shape (narrow and well-defined peak),
- A small to intermediate amplitude,
- An isolated setting,
- A short strike length (in general, not exceeding one kilometre), and
- Preferably, with a localized magnetic anomaly of matching dimensions.

### **Surficial Conductors**

This term is used for geological conductors in the overburden, either glacial or residual in origin, and in the weathered layer of the bedrock. Most surficial conductors are probably caused by clay minerals. In some environments the presence of salts will contribute to the conductivity. Other possible electrolytic conductors are residual soils, swamps, brackish ground water and alluvium such as lake or river-bottom deposits, flood plains and estuaries.

Normally, most surficial materials have low to intermediate conductivity so they are not easily mistaken for highly conductive bedrock features. Also, many of them are wide and their anomaly shapes are typical of broad horizontal sheets.

When surficial conductivity is high it is usually still possible to distinguish between a horizontal plate

(more likely to be surficial material) and a vertical body (more likely to be a bedrock source) thanks to the asymmetry of the fixed-wing system responses observed at the edges of a broad conductor when flying adjacent lines in opposite directions. The configuration of the system is such that the response recorded at the leading edge is more pronounced than that registered at the trailing edge. Figure 1 illustrates the "edge effect". In practice there are many variations on this very diagnostic phenomenon.

One of the more ambiguous situations as to the true source of the response is when surface conductivity is related to bedrock lithology as for example, surface alteration of an underlying bedrock unit. At times, it is also difficult to distinguish between a weak conductor within the bedrock (e.g. near-massive sulphides) and a surficial source.

In the search for massive sulphides or other bedrock targets, surficial conductivity is generally considered as interference but there are situations where the interpretation of surficial-type conductors is the primary goal. When soils, weathered or altered products are conductive, and in-situ, the responses are a very useful aid to geologic mapping. Shears and faults are often identified by weak, usually narrow, anomalies.

Analysis of surficial conductivity can be used in the exploration for such features as lignite deposits, kimberlites, palaeochannels and ground water. In coastal or arid areas, surficial responses may serve to define the limits of fresh, brackish and salty water.

### **Cultural Conductors**

The majority of cultural anomalies occurs along roads and is accompanied by a response on the power line monitor. (This monitor is set to 50 or 60 Hz, depending on the local power grid.) In some cases, the current induced in the power line results in anomalies that could be mistaken for bedrock responses. There are also some power lines that have no response whatsoever.

The power line monitor, of course, is of great assistance in identifying cultural anomalies of this type. It is important to note, however, that geological conductors in the vicinity of power lines may exhibit a weak response on the monitor because of current induction via the earth.

Fences, pipelines, communication lines, railways and other man-made conductors can give rise to responses, the strength of which will depend on the grounding of these objects.

Another facet of this analysis is the line-to-line comparison of anomaly character along suspected man-made conductors. In general, the amplitude, the rate of decay, and the anomaly width should not vary a great deal along any one conductor, except for the change in amplitude related to terrain clearance variation. A marked departure from the average response character along any given feature gives rise to the possibility of a second conductor.

In most cases a visual examination of the site will suffice to verify the presence of a man-made conductor. If a second conductor is suspected the ground check is more difficult to accomplish. The object would be to determine if there is (i) a change in the man-made construction, (ii) a difference in the grounding conditions, (iii) a second cultural source, or (iv) if there is, indeed, a geological conductor in addition to the known man-made source.

The selection of targets from within extensive (formational) belts is much more difficult than in the case of isolated conductors. Local variations in the EM characteristics, such as in the amplitude,

decay, shape etc., can be used as evidence for a relatively localized occurrence. Changes in the character of the EM responses, however, may be simply reflecting differences in the conductive formations themselves rather than indicating the presence of massive sulphides and, for this reason, the degree of confidence is reduced.

Another useful guide for identifying localized variations within formational conductors is to examine the magnetic data in map or image form. Further study of the magnetic data can reveal the presence of faults, contacts, and other features, which, in turn, help define areas of potential economic interest.

Finally, once ground investigations begin, it must be remembered that the continual comparison of ground knowledge to the airborne information is an essential step in maximizing the usefulness of the airborne EM data.

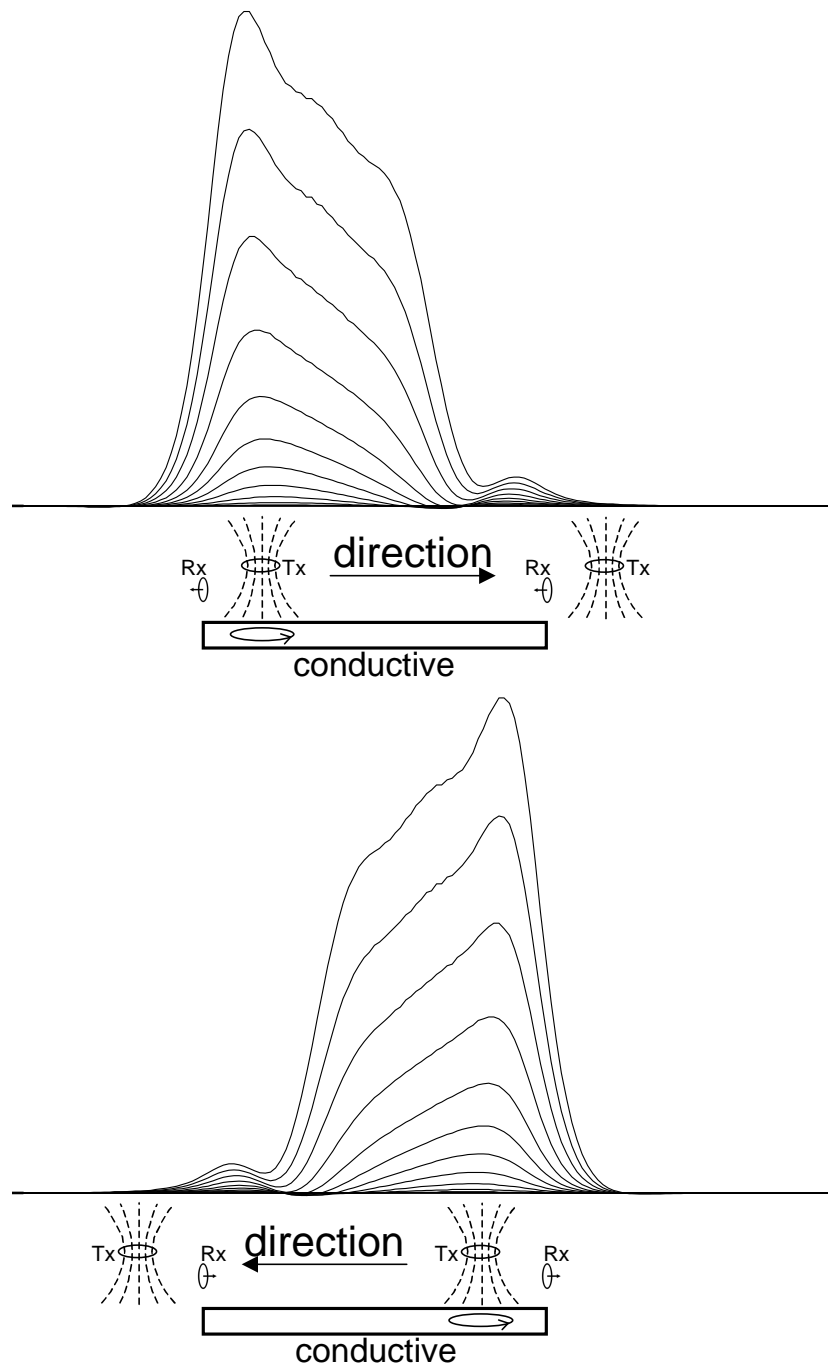


Figure 1. Illustration of how the x-component response varies depending on the flight direction. When the receiver flies onto the conductor, the transmitter is over the conductor and current is induced in the conductive material, resulting in a large response. When the receiver flies off the conductor, the transmitter is not over conductive material, so the response is small.

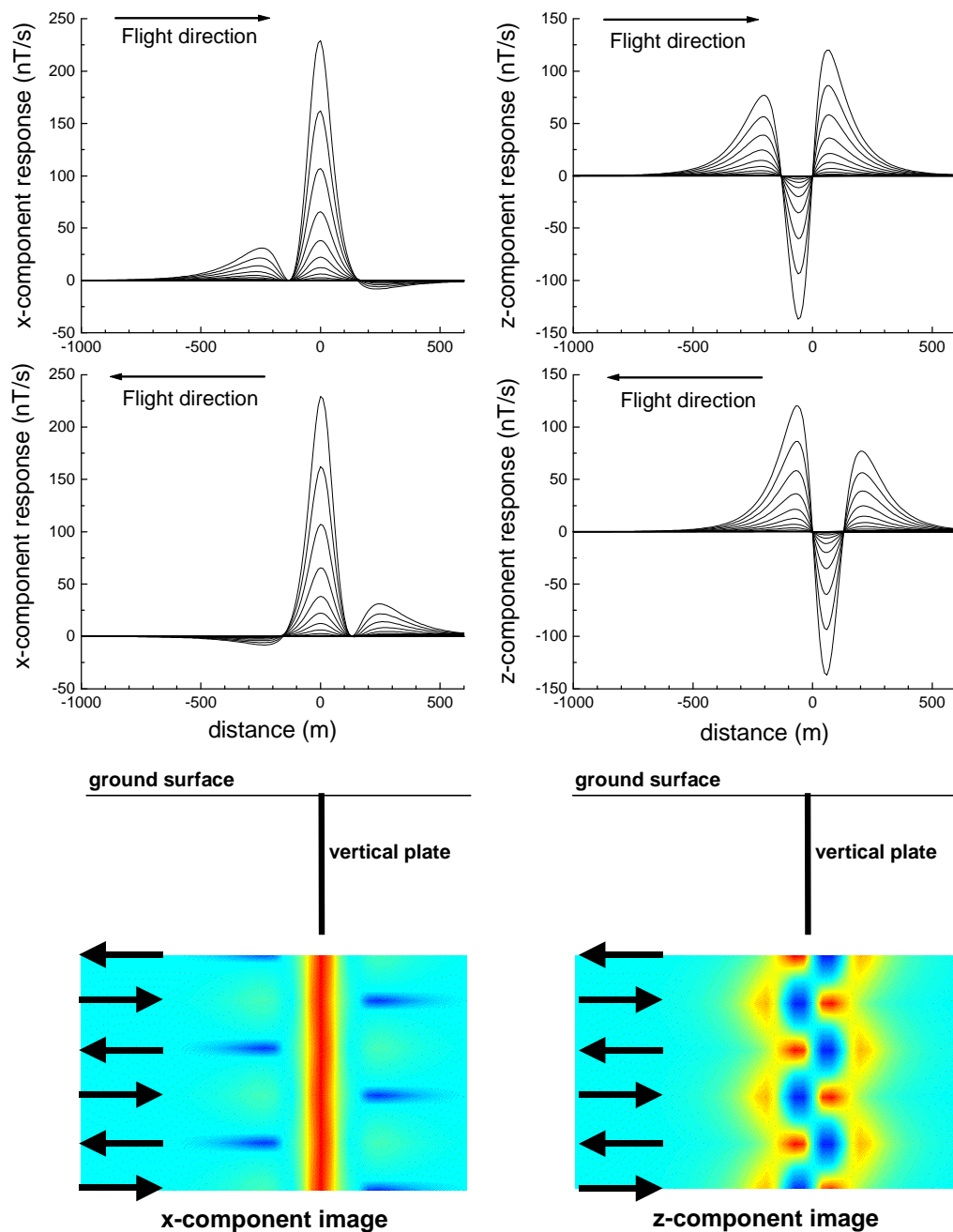


Figure 2. The response over a vertical plate. The left panels show the x-component, the right panels the z-component. The top is flying left to right, the middle is right to left, the bottom is a plan image with the alternating flight directions shown with arrows.



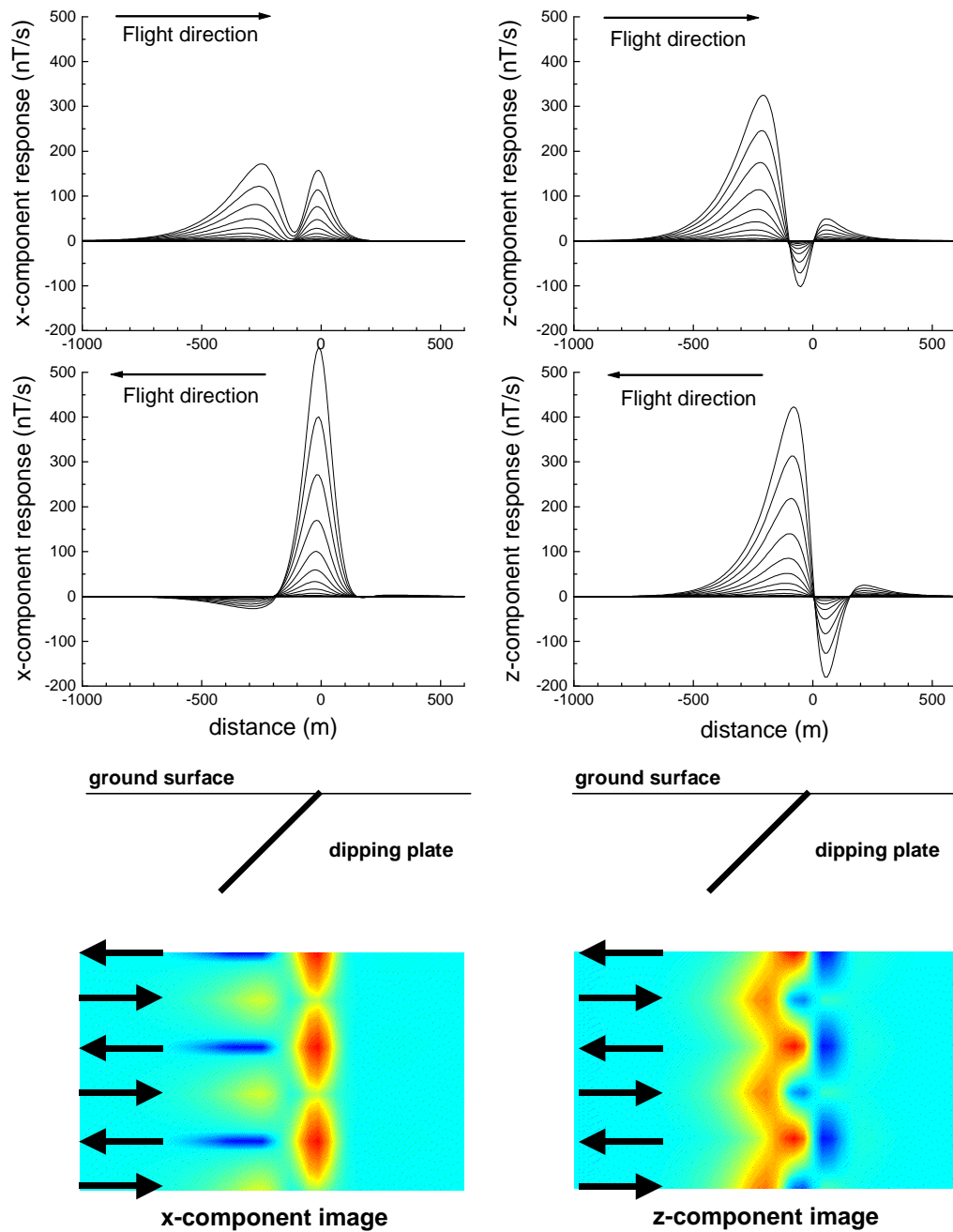


Figure 3. The response over a 45 degree dipping plate. The left panels show the x-component, the right panels the z-component. The top is flying left to right, the middle is right to left, the bottom is a plan image with the alternating flight directions shown with arrows.

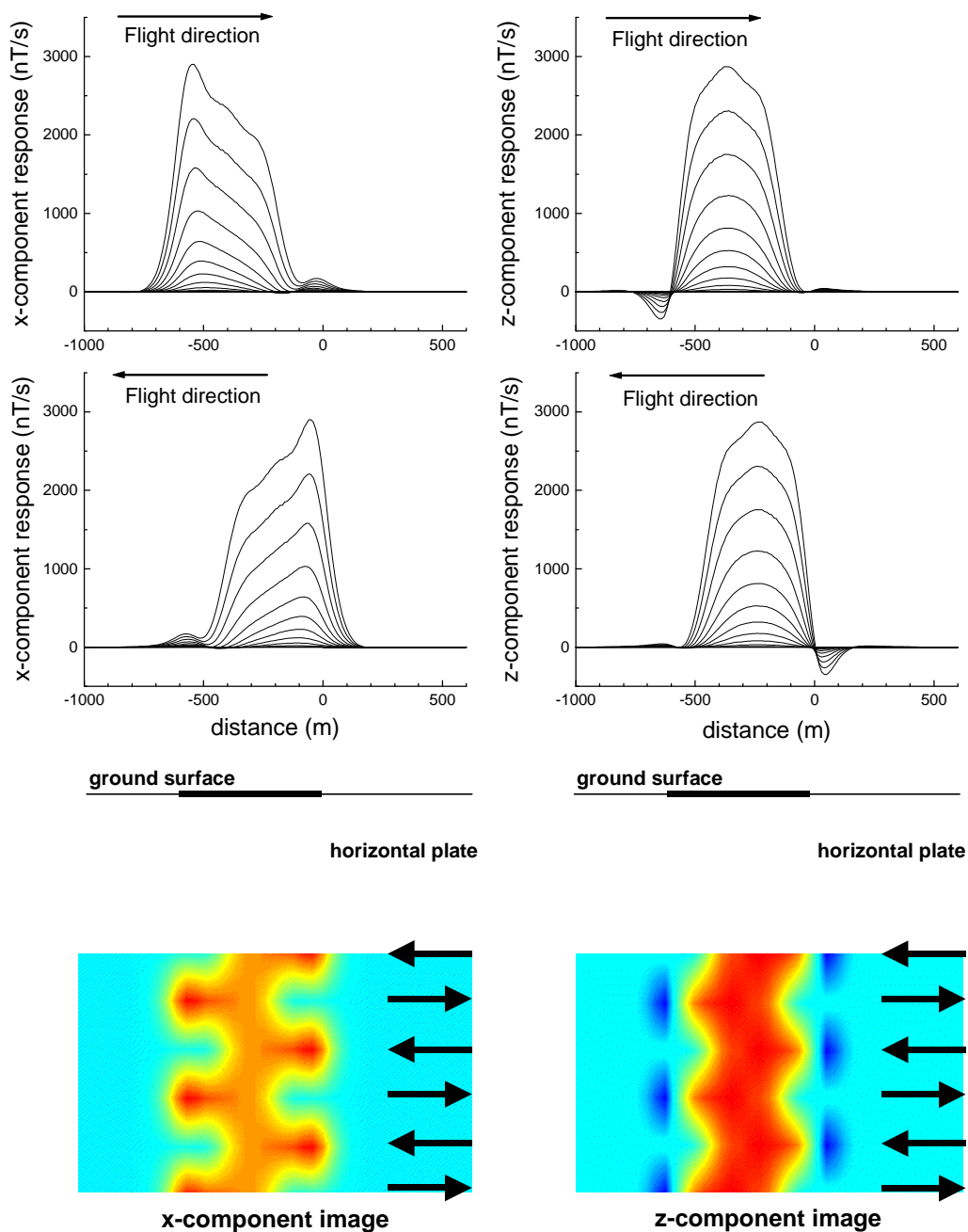


Figure 4. The response over a horizontal plate. The left panels show the x-component, the right panels the z-component. The top is flying left to right, the middle is right to left, the bottom is a plan image with the alternating flight directions shown with arrows.

## **Appendix C**

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### Multi-component Modeling

## Multi-component fixed-wing airborne EM modeling

### PLATE MODELING

The PLATE program has been used to generate synthetic responses over a number of plate models with varying depth of burial (0, 150 and 300 m) and dips (0, 45, 90 and 135 degrees). The geometry assumed for the fixed-wing airborne EM system is shown on the following page (Figure 1), and the transmitter waveform on the subsequent page (Figure 2). In these models, the receiver is 130 m behind and 50 m below the transmitter center.

In all cases the plate has a strike length of 600 m, with a strike direction into the page. The width of the plate is 300 m. As the flight path traverses the center of the plate, the y-component is zero and has not been plotted.

The conductance of the plate is 20 S. In cases when the conductance is different, an indication of how the amplitudes may vary can be obtained from the nomogram included (Figure 3).

In the following profile plots (Figure 4 to 15) the plotting point is the receiver location and all of the component values are in nT/s, assuming a transmitter dipole moment of  $900\,000\text{ Am}^2$ . If the dipole moment is larger or smaller than  $900\,000\text{ Am}^2$ , then the response would be scaled up or down appropriately.

In the following profile plots (Figure 4 to 15) all components are in nT/s, for a transmitter dipole moment of  $900\,000\text{ Am}^2$ . If the dipole moment is larger or smaller, then the response should be scaled up or down appropriately.

The plotting point is the receiver location.

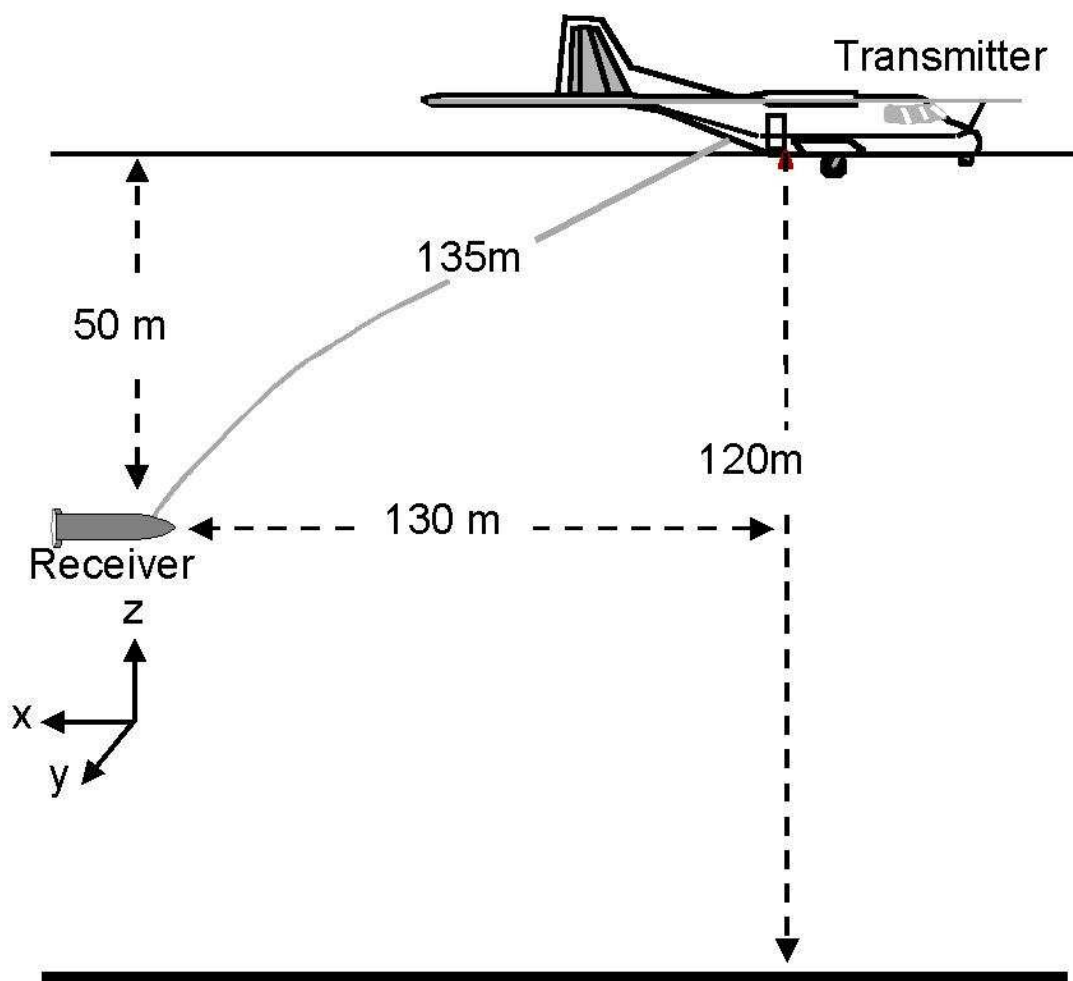


Figure 1. Nominal geometry of the MEGATEM®/GEOTEM® system.

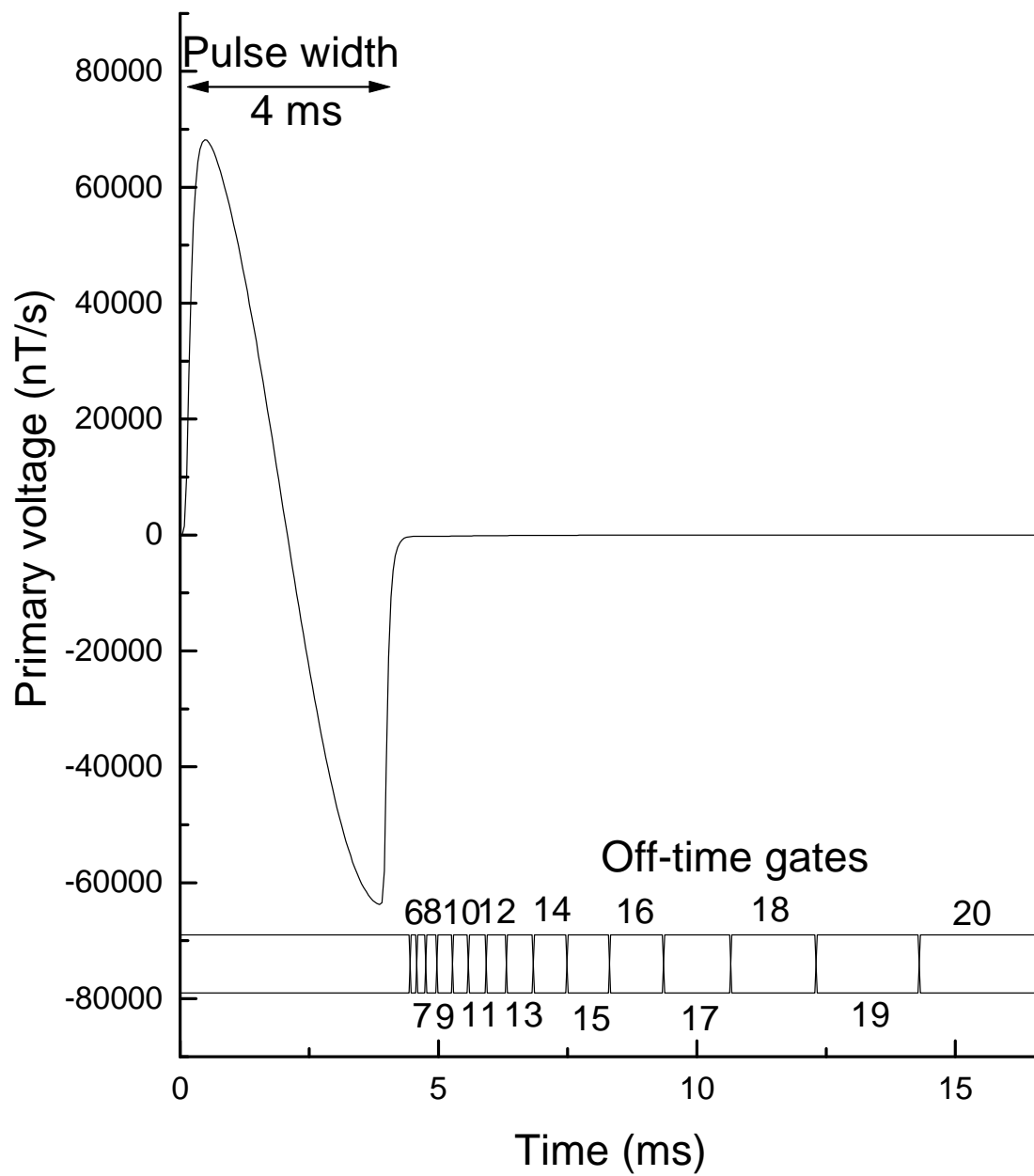


Figure 2. Theoretical transmitter waveform response in the receiver.

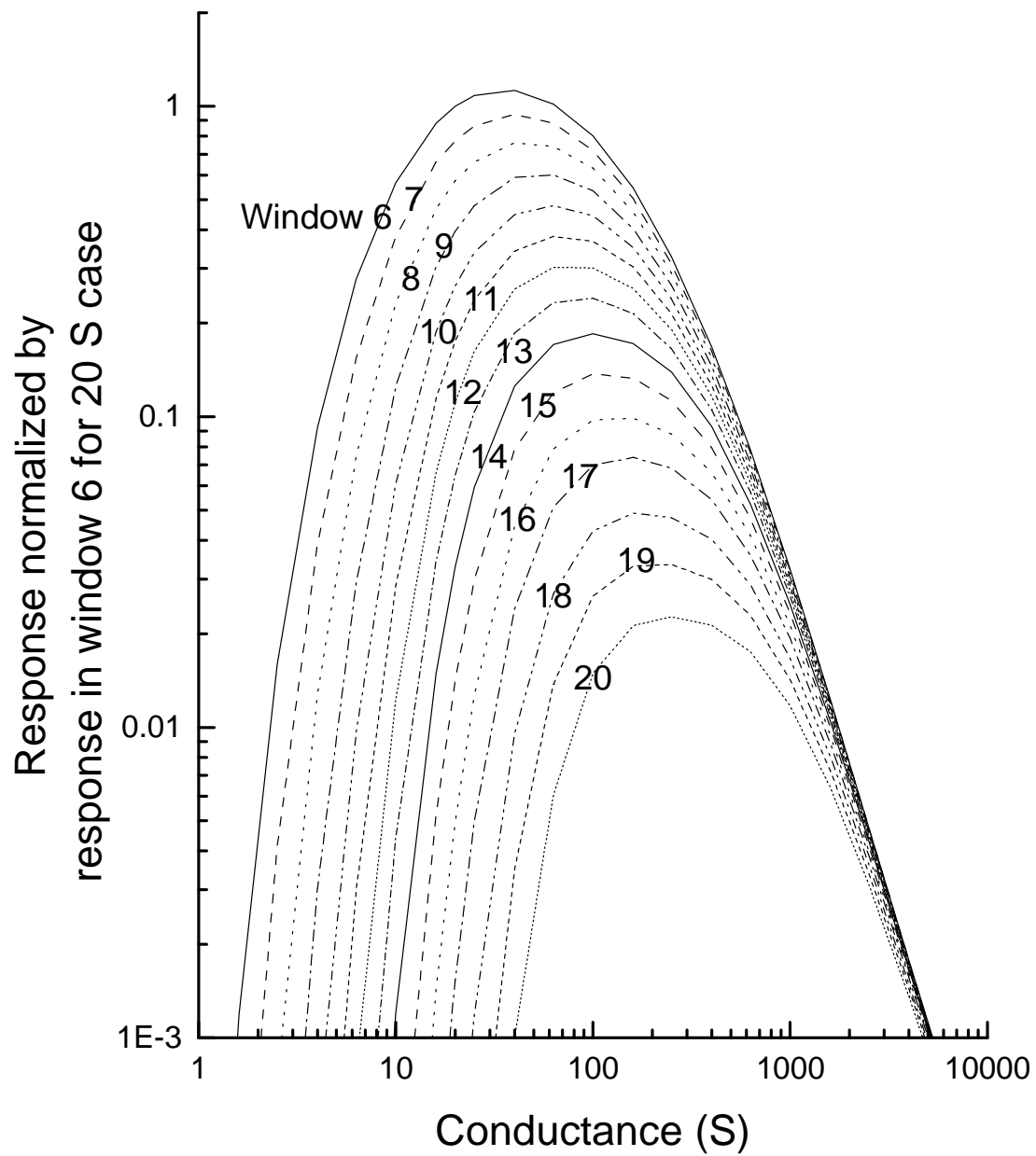


Figure 3. Nomogram for windows 6-20 normalized to a response from a 20-siemen conductor in window 6.

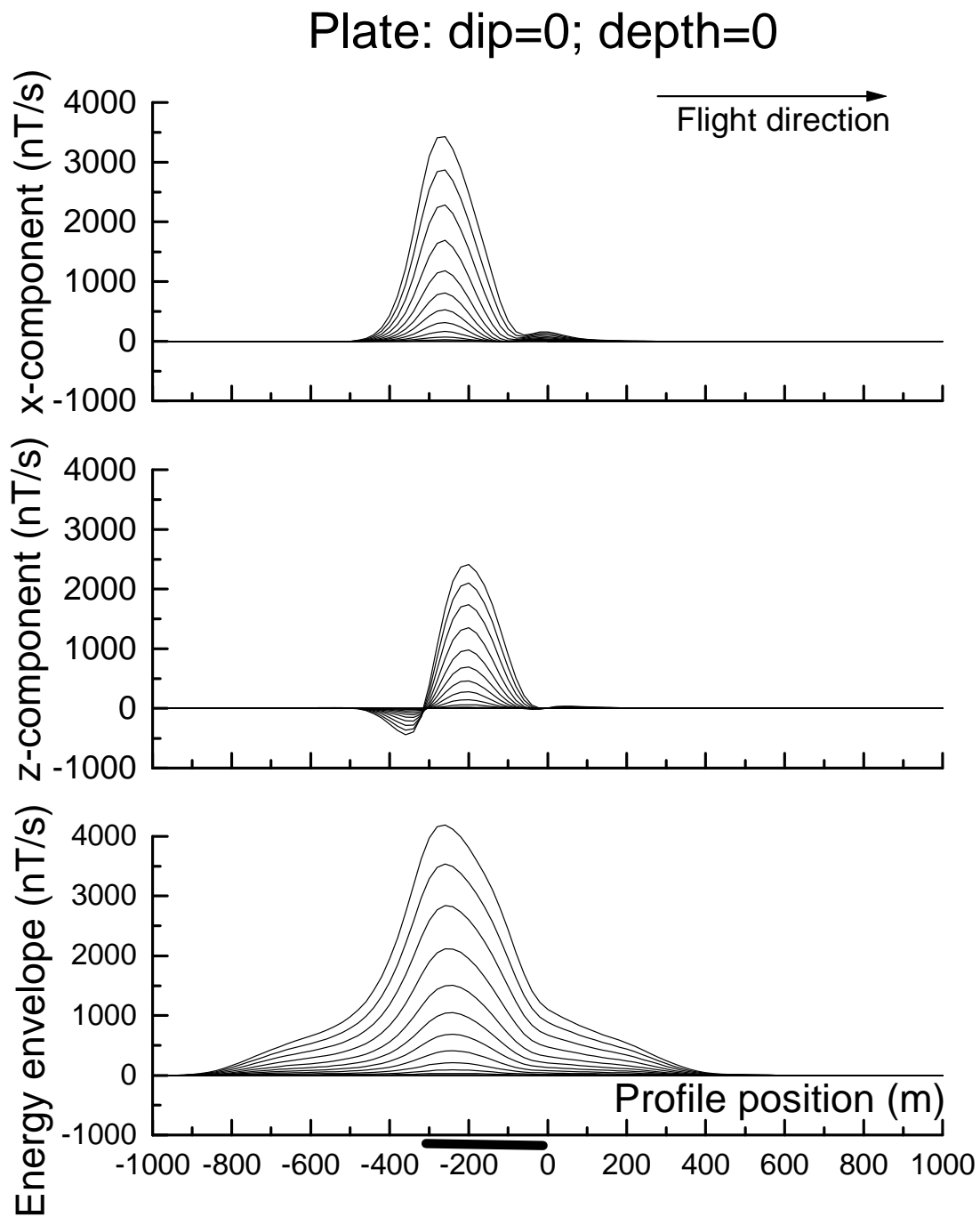


Figure 4.



Plate: dip=0; depth=150

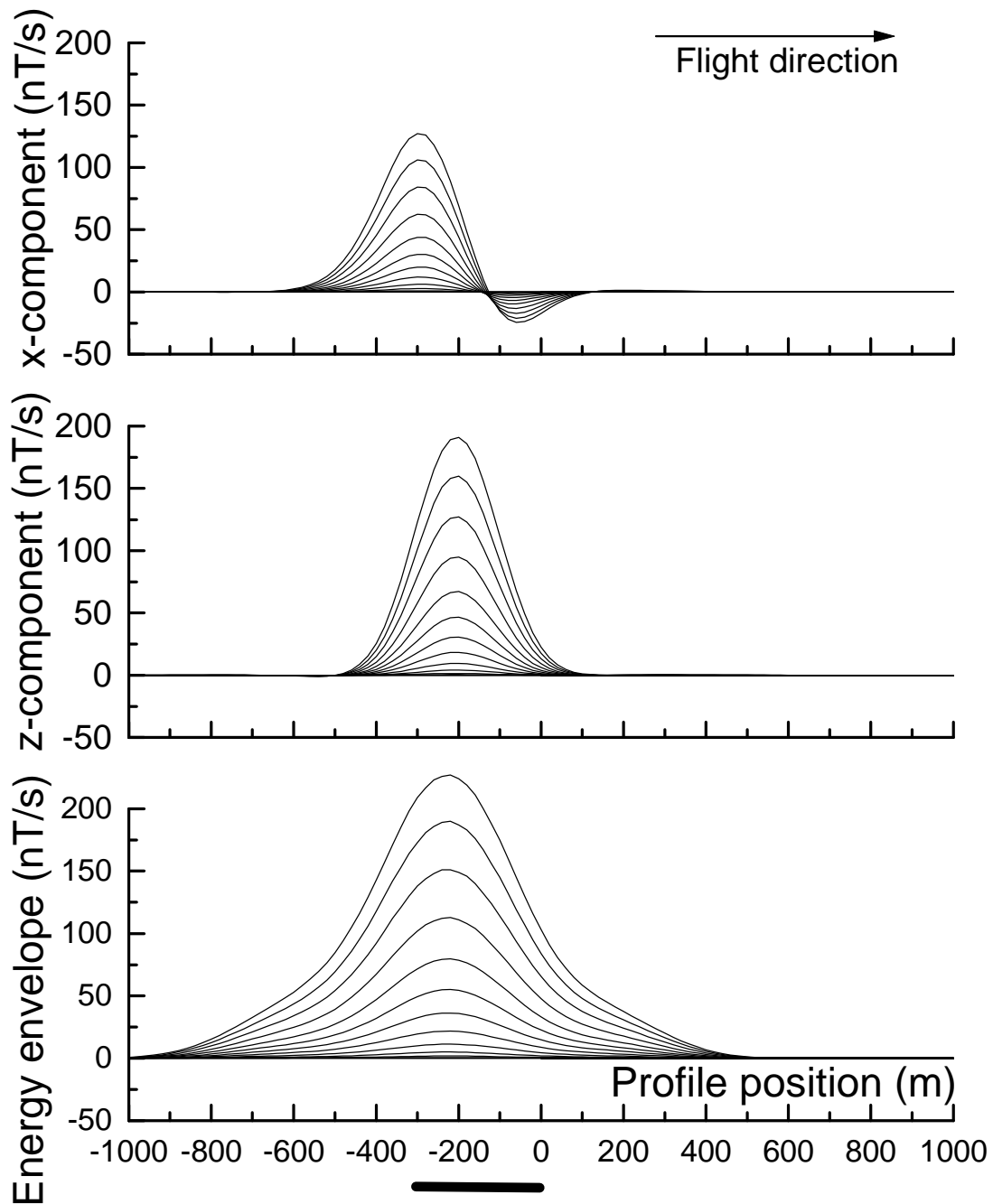


Figure 5.

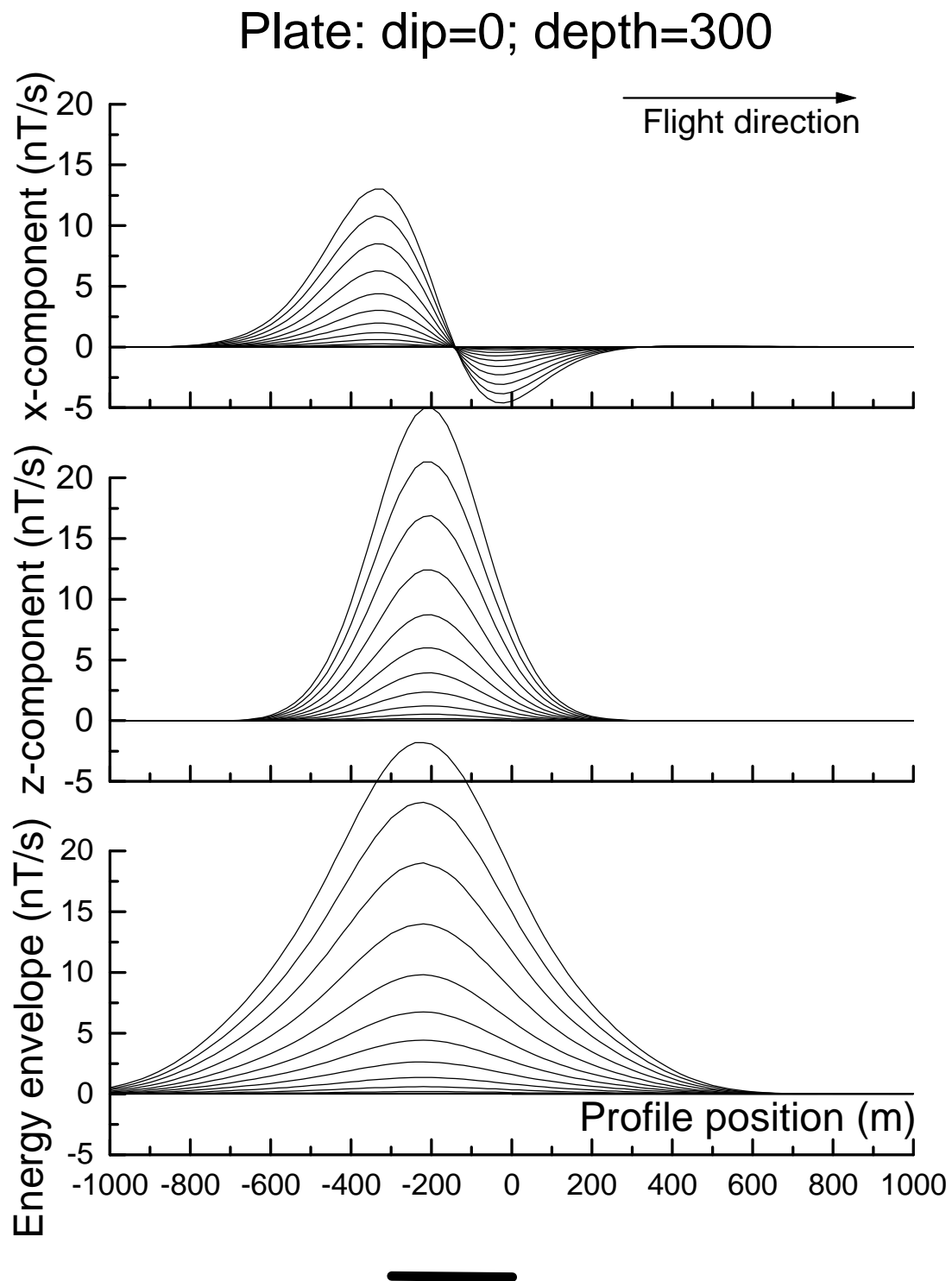


Figure 6.

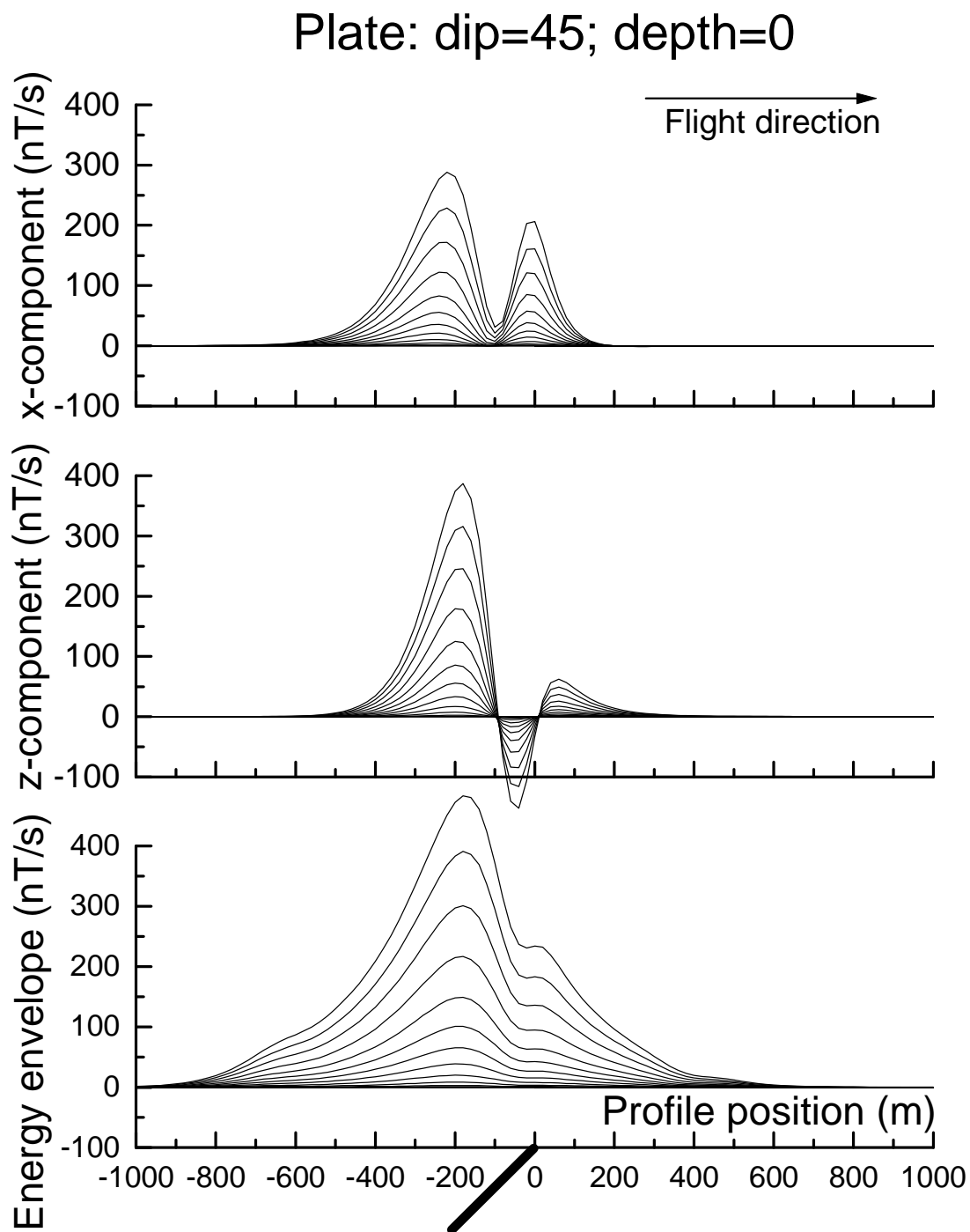


Figure 7.

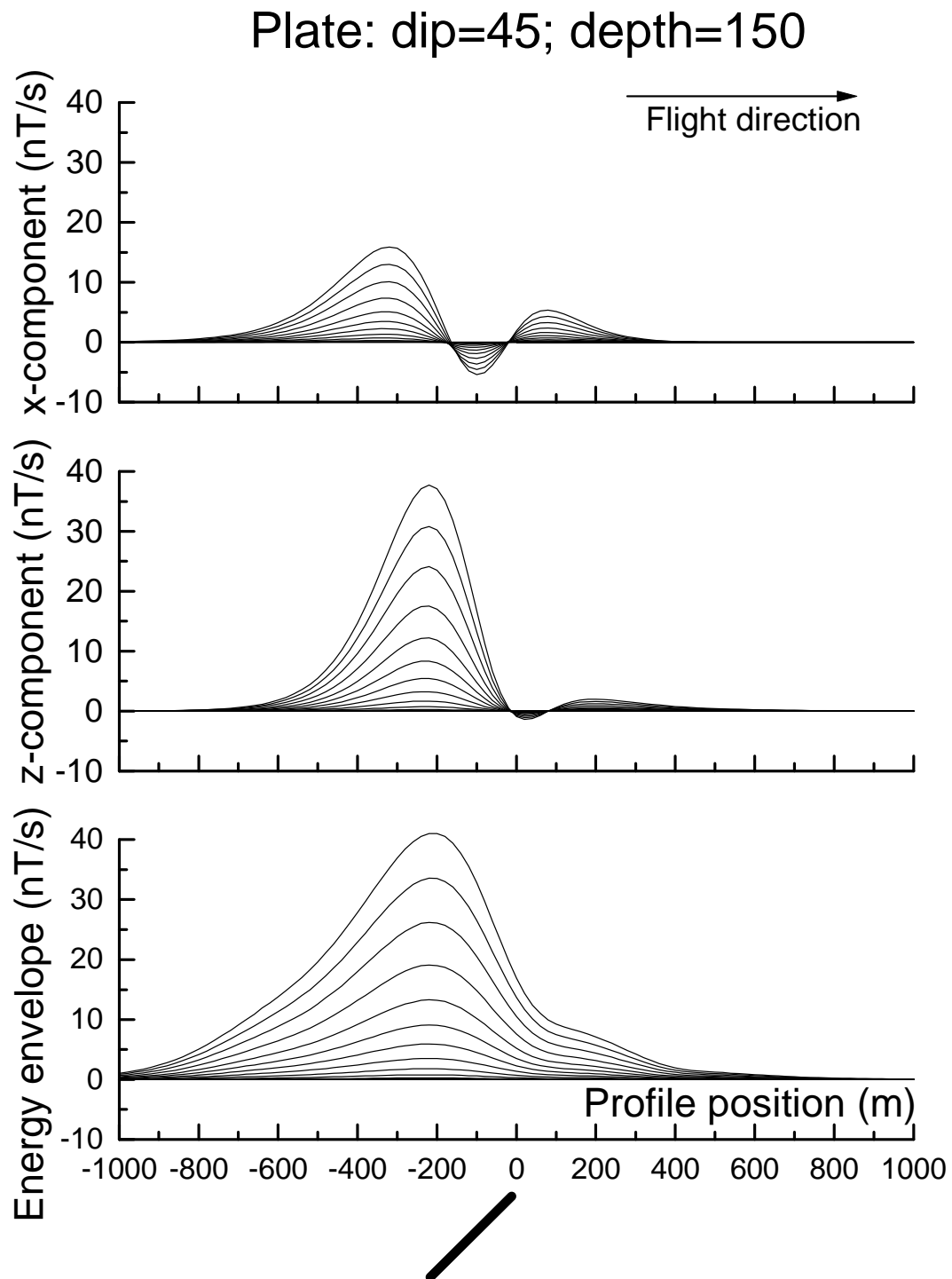


Figure 8.

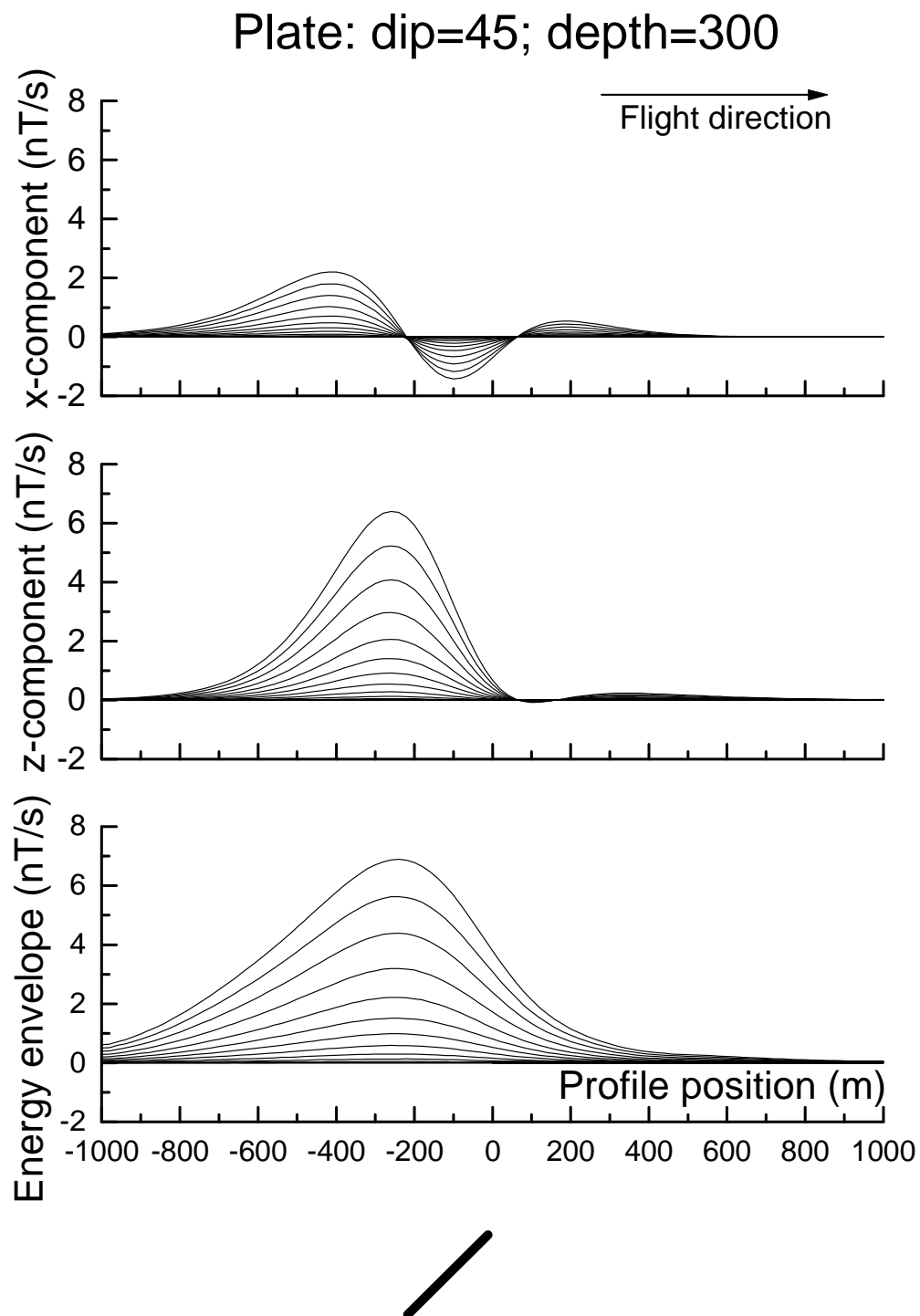


Figure 9.

Plate: dip=90; depth=0

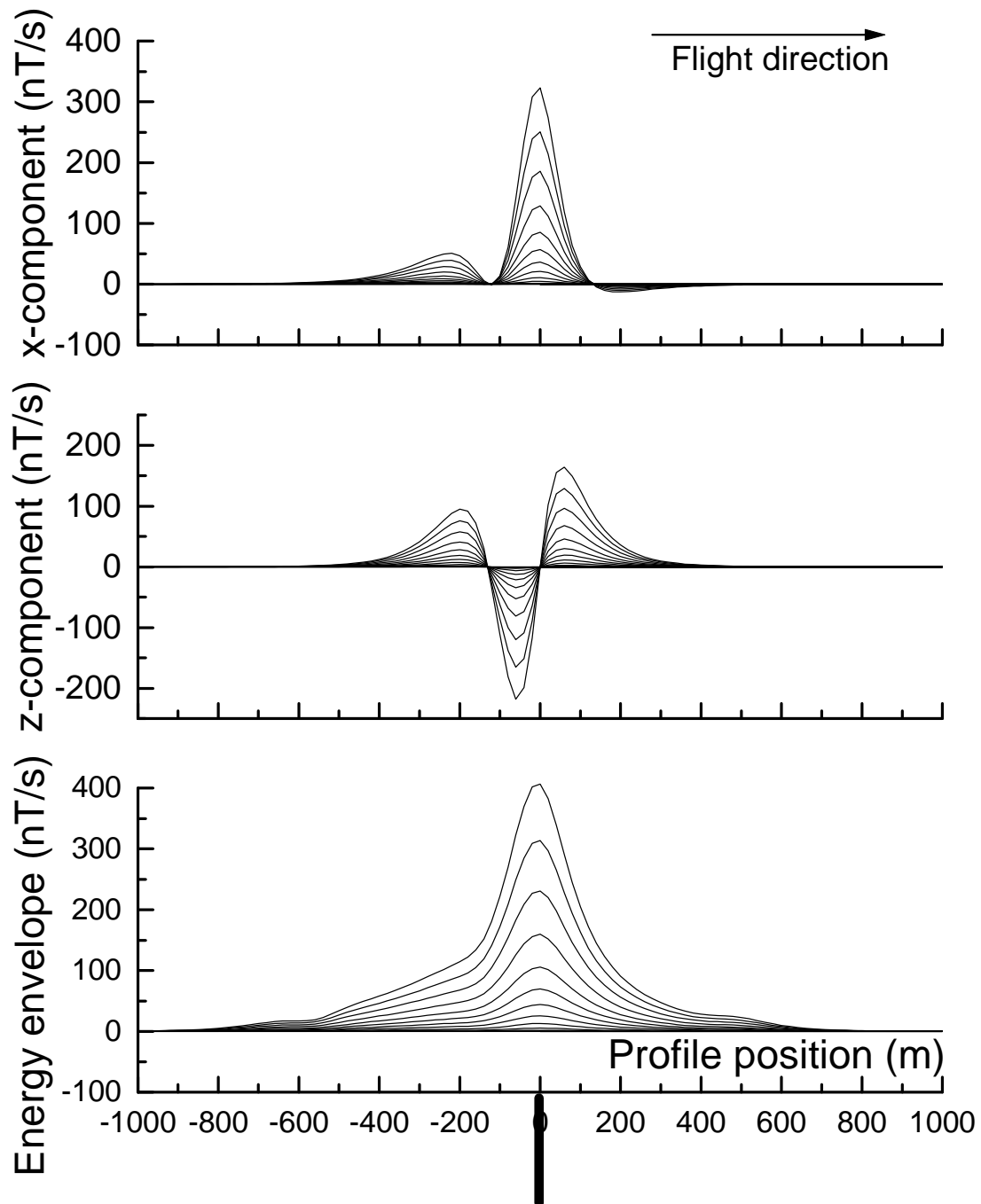


Figure 10.

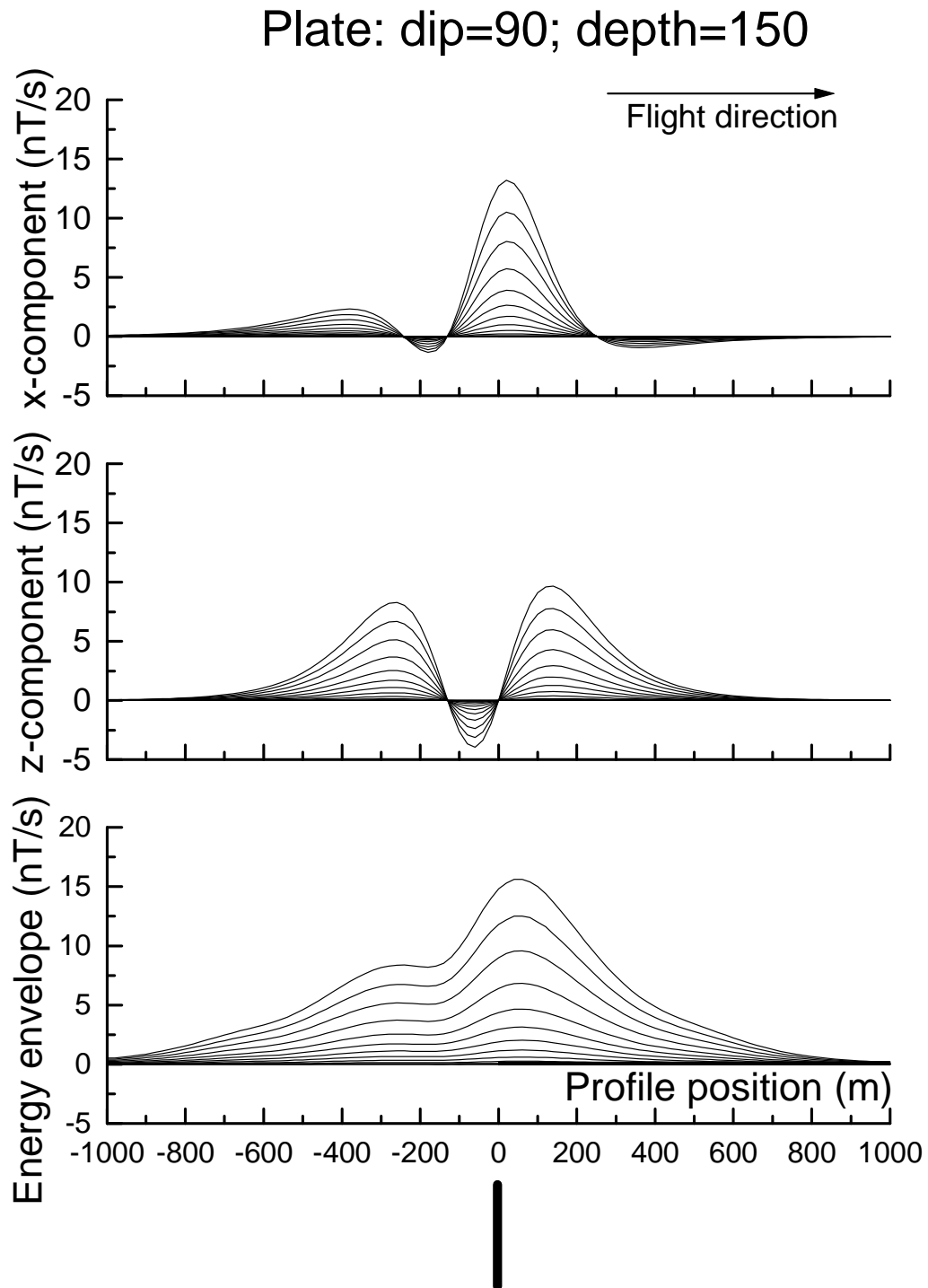


Figure 11.

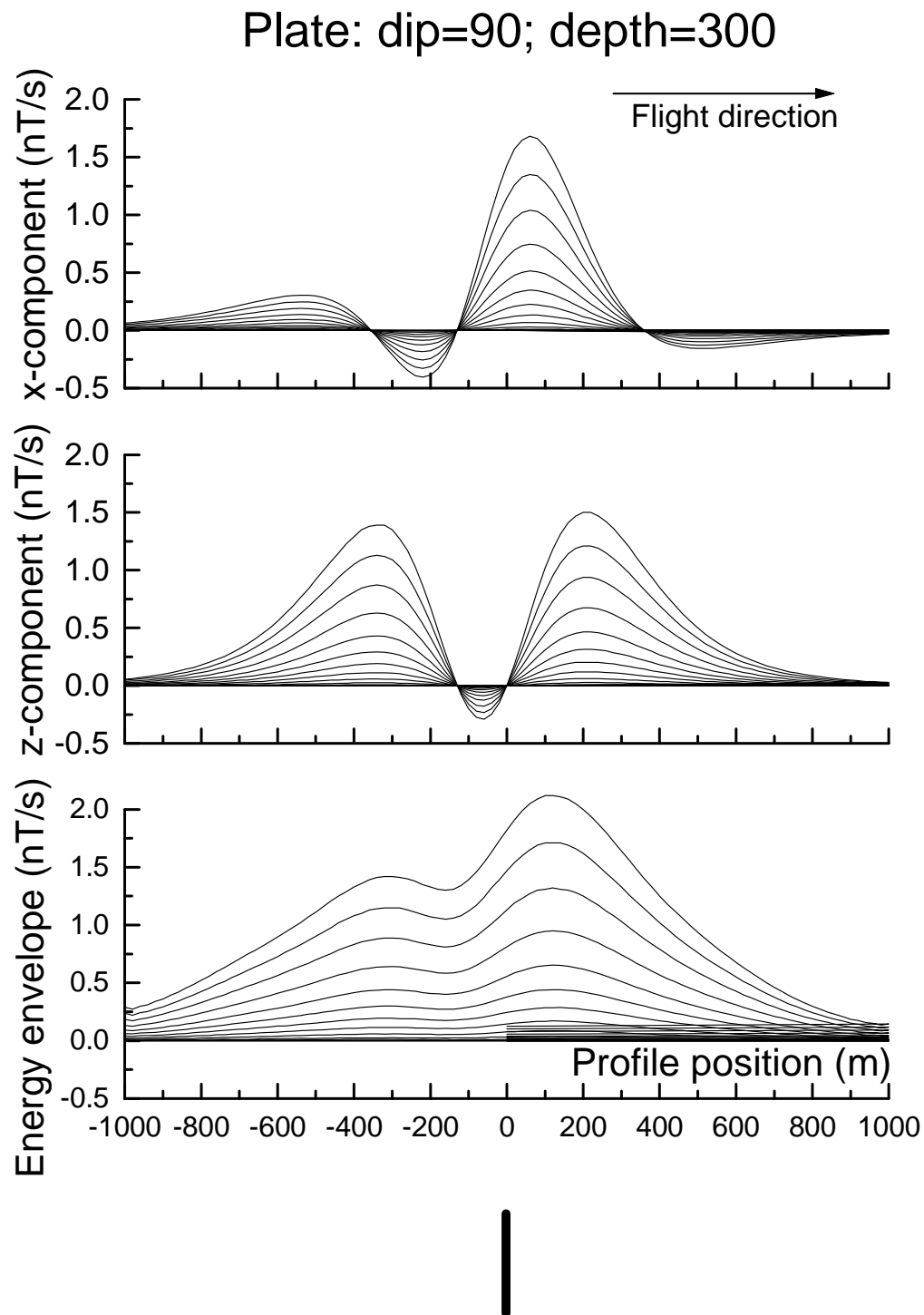


Figure 12.



Plate: dip=135; depth=0

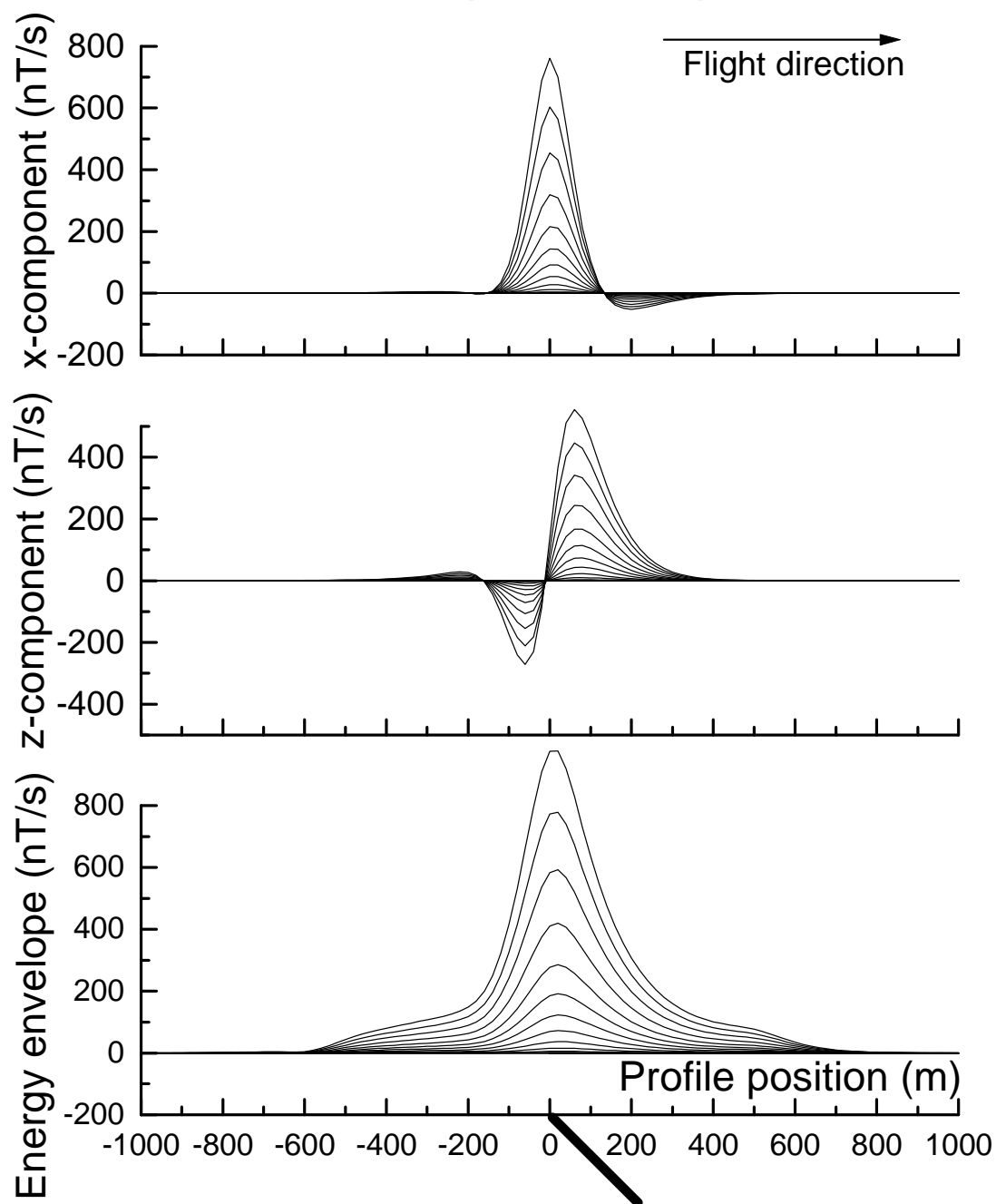


Figure 13.

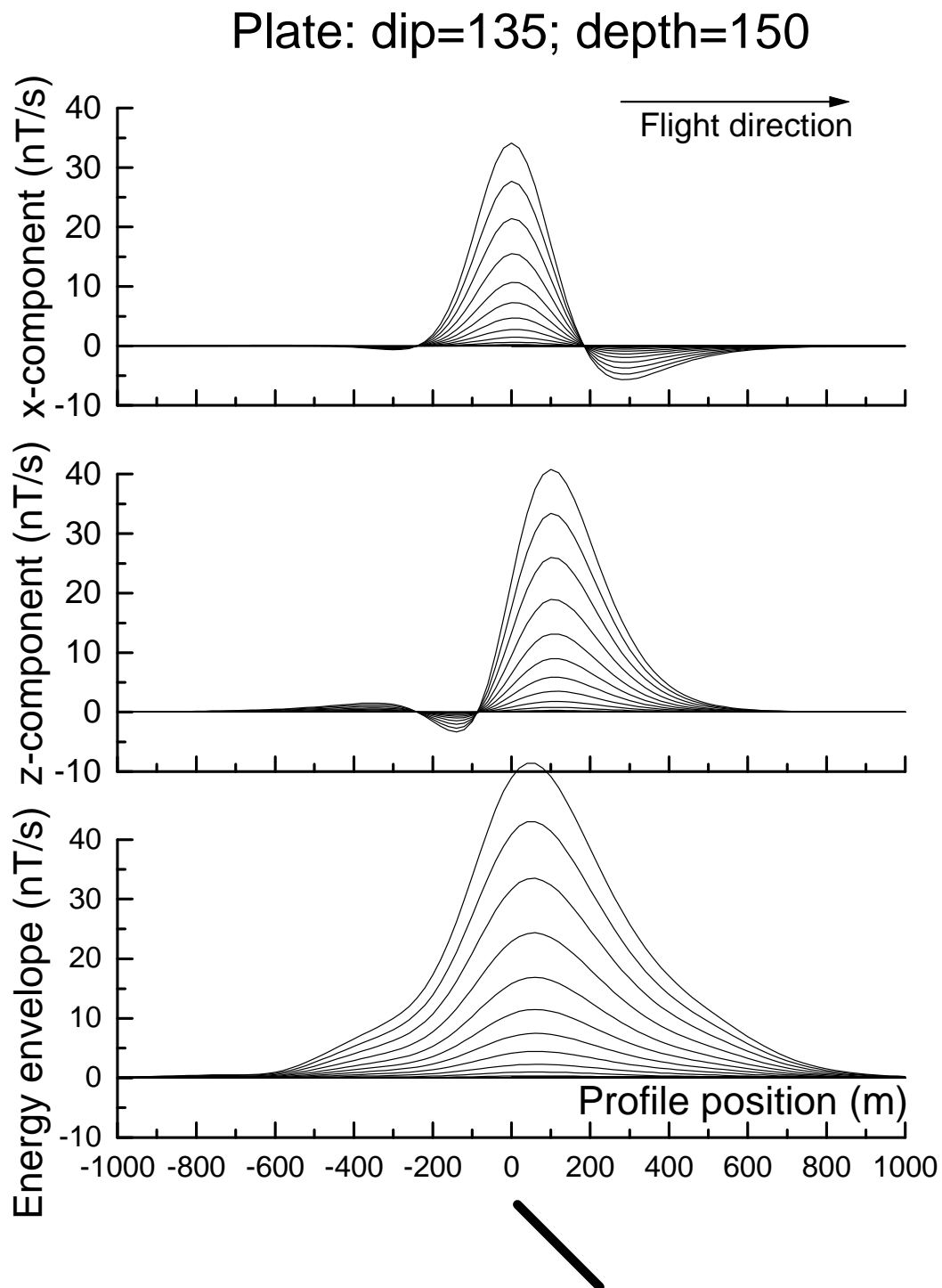


Figure 14.

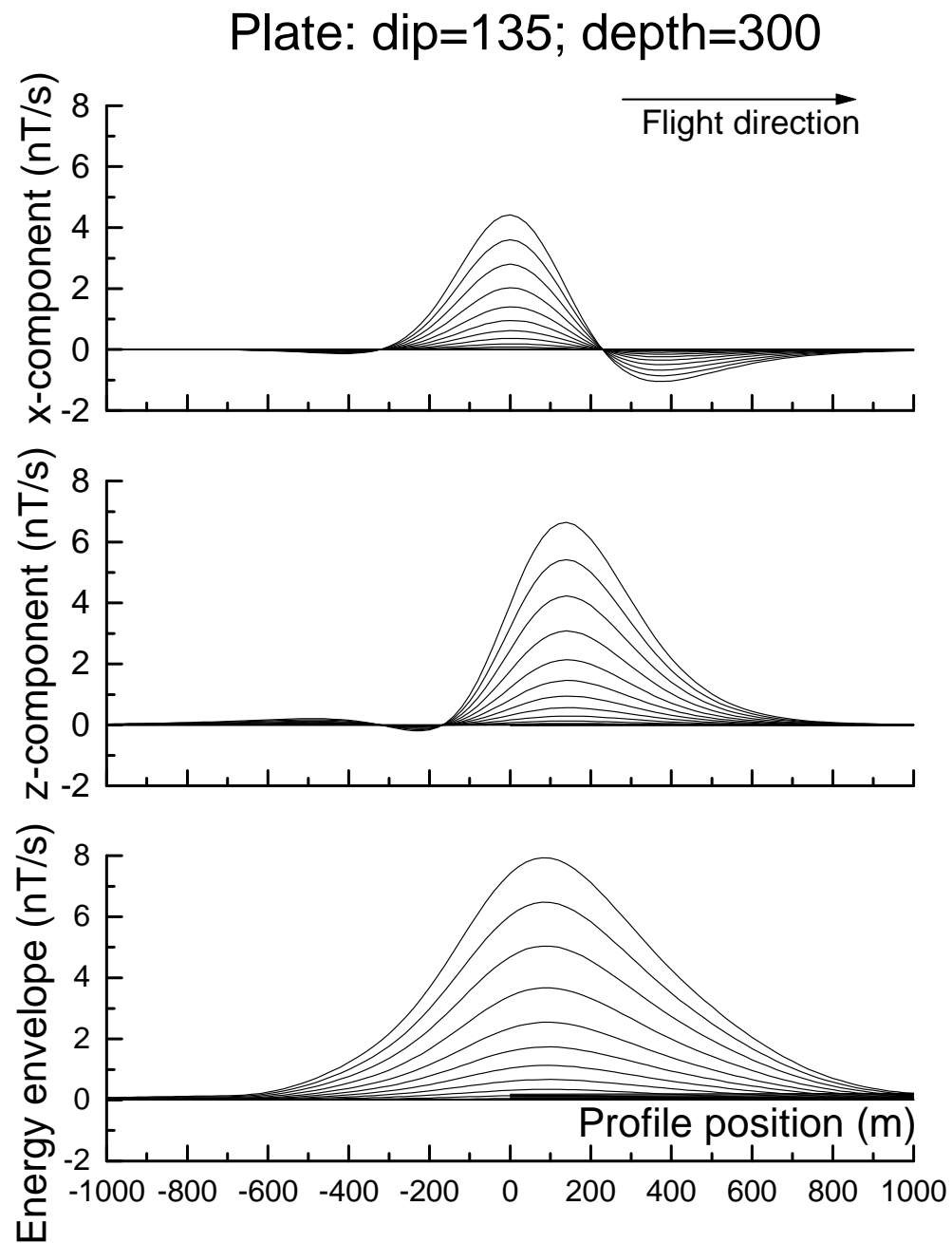


Figure 15.

## **SPHERE MODELING**

The sphere in a uniform field program (Smith and Lee, Exploration Geophysics, 2001, pp 113-118) has been used to generate synthetic responses over a number of sphere models with varying depth of burial (0, 150 and 300 m). The geometry assumed for the fixed-wing airborne EM system and the waveform are as shown in Figures 1 and 2 above.

In all cases the sphere has a radius of 112 m. As the flight path traverses the center of the sphere, the y component is zero and has not been plotted.

The conductivity of the sphere is 1 S/m. In cases when the conductivity is different, an indication of how the amplitudes may vary can be obtained from the nomogram that follows (Figure 16).

In the following profile plots (Figure 17 to 19) all components are in nT/s, for a transmitter dipole moment of 900 000 Am<sup>2</sup>. If the dipole moment is larger or smaller, then the response should be scaled up or down appropriately.

The plotting point is the receiver location.

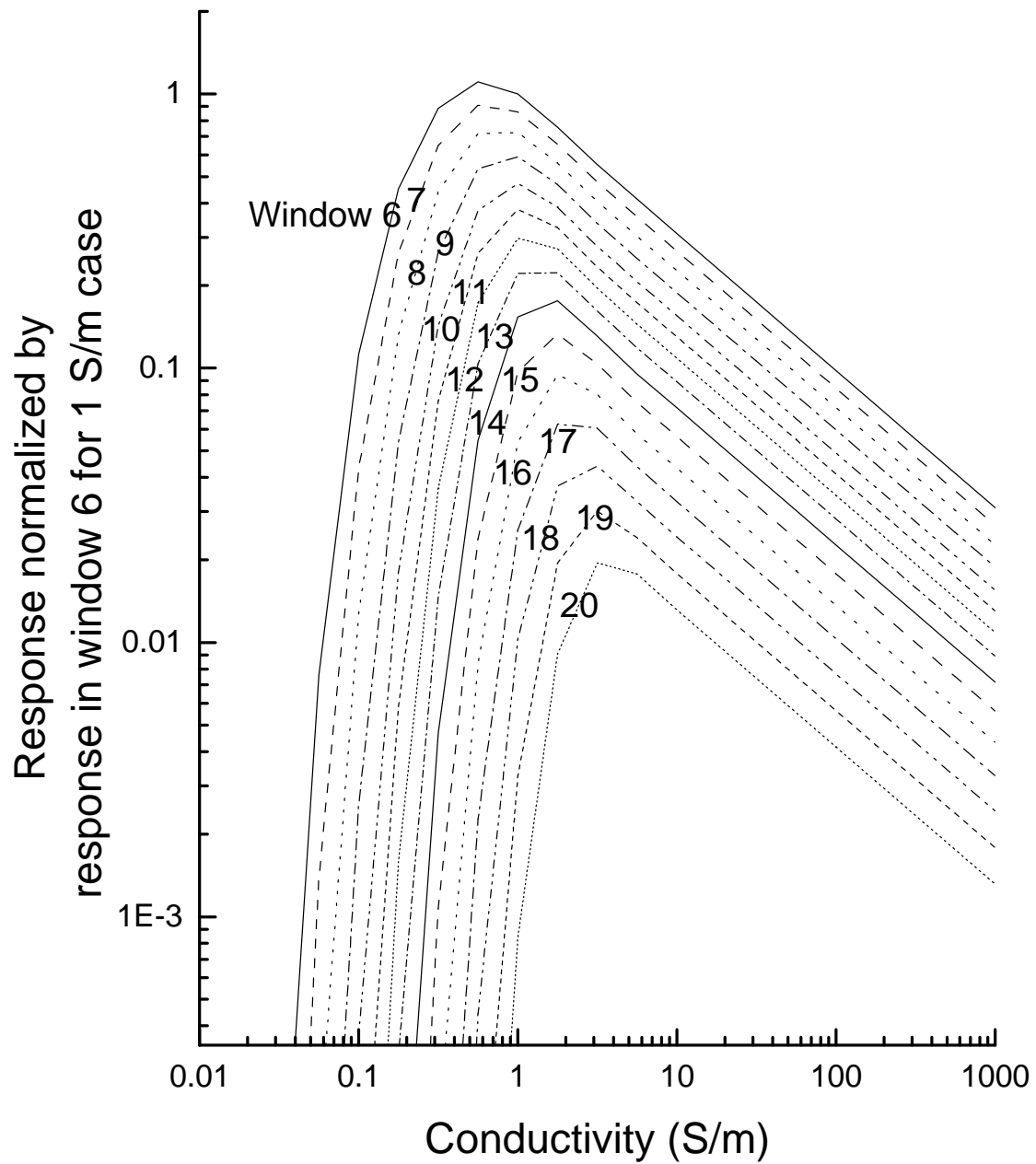


Figure 16. Nomogram for windows 6-20 normalized to a response from a 1-siemens conductor in window 6.

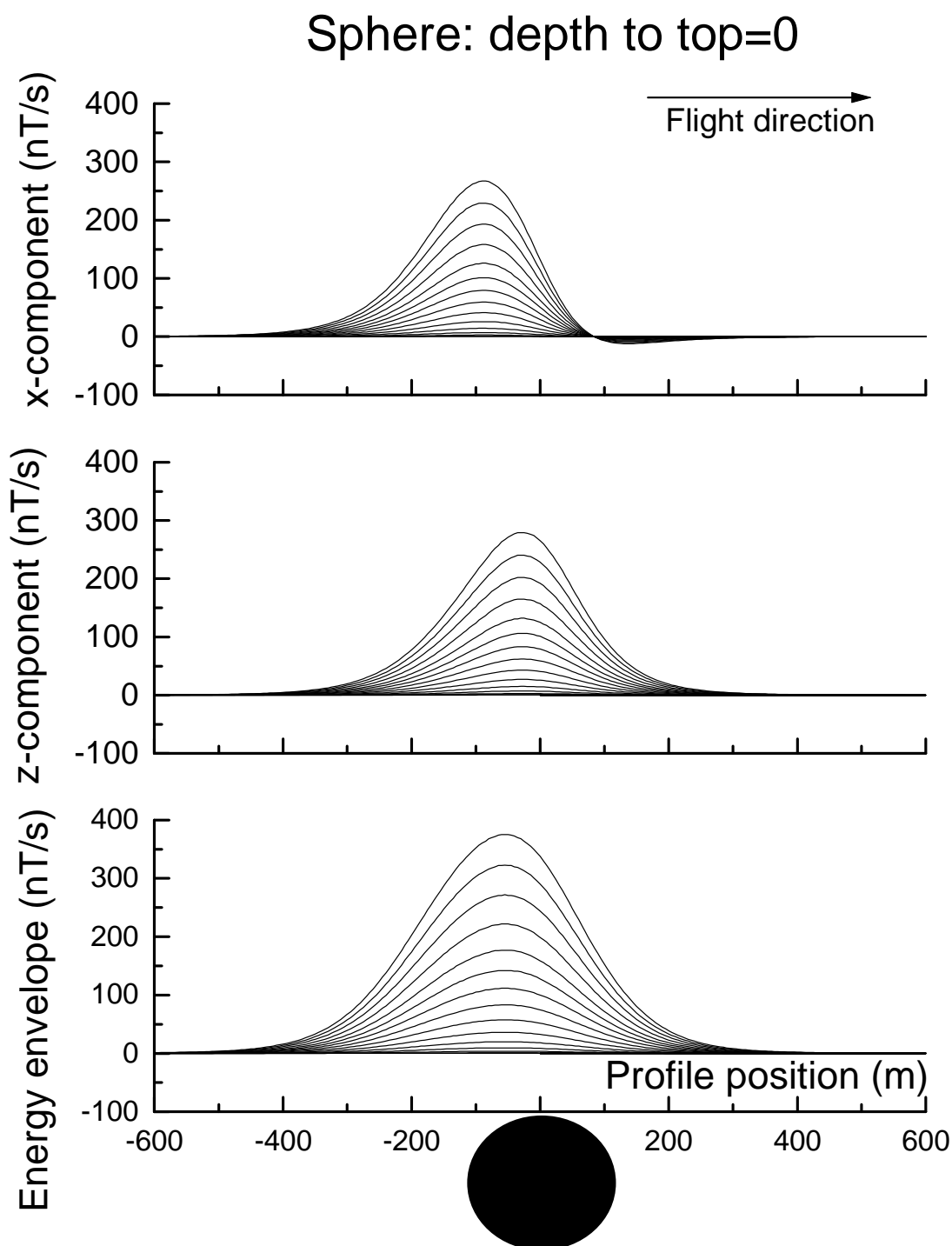


Figure 17.

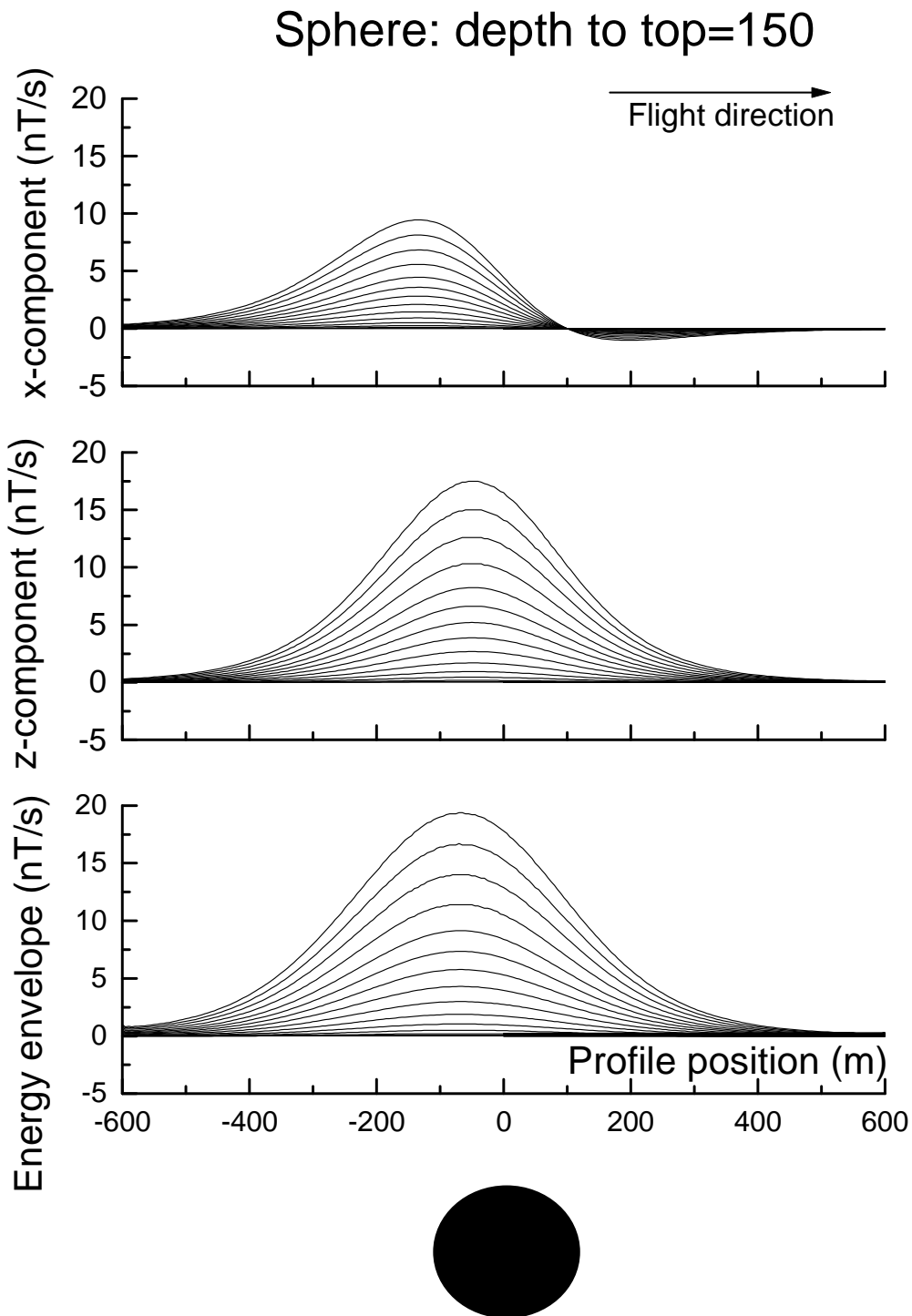


Figure 18.

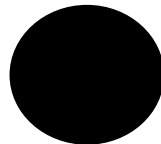
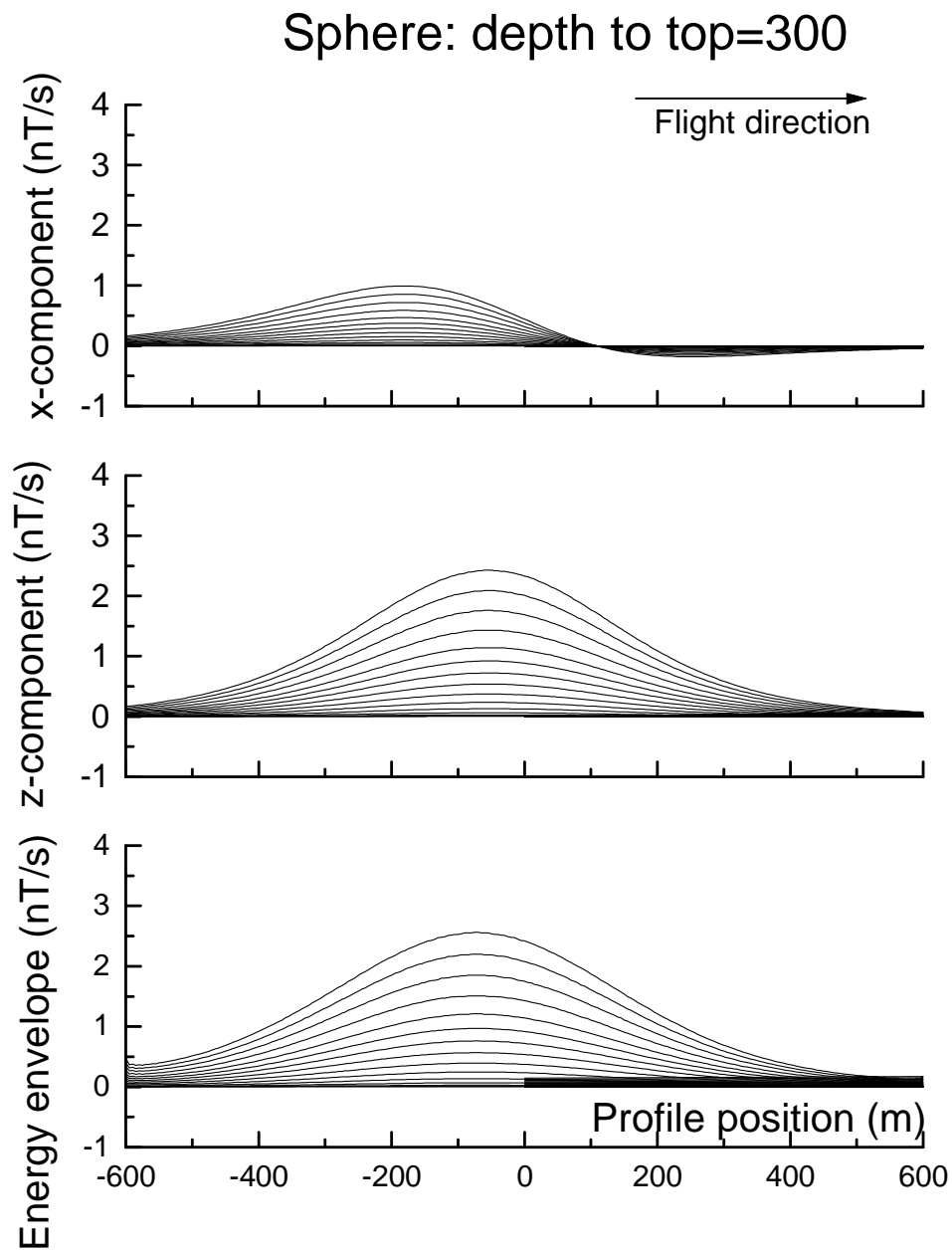


Figure 19.



## **Appendix D**

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# **The Usefulness of Multi-component, Time-Domain Airborne Electromagnetic Measurement**

GEOPHYSICS, VOL 61, NO. 1 (JANUARY-FEBRUARY 1996); P. 74-81, 17 FIGS.

## **The usefulness of multi-component, time-domain airborne electromagnetic measurements**

Richard S. Smith\* and Pierre B. Keating ‡

### **ABSTRACT**

Time-domain airborne electromagnetic (AEM) systems historically measure the inline horizontal (x) component. New versions of the electromagnetic systems are designed to collect two additional components [the vertical (z) and the lateral horizontal (y) component] to provide greater diagnostic information.

In areas where the geology is near horizontal, the z component response provides greater signal to noise, particularly at late delay times. This allows the conductivity to be determined to greater depth. In a layered environment, the symmetry implies that the y component will be zero; hence a non-zero y component will indicate a lateral inhomogeneity.

The three components can be combined to give the “energy envelope” of the response. Over a vertical plate, the response profile of this envelope has a single positive peak and no side lobes. The shape of the energy envelope is dependent on the flight direction, but less so than the shape of the x component response profile.

In the interpretation of discrete conductors, the z component data can be used to ascertain the dip and depth to the conductor using simple rules of thumb. When the profile line is perpendicular to the strike direction and over the center of the conductor, the y component will be zero; otherwise it appears to be a combination of the x and z components. The extent of the contamination of the y component by the x and z components can be used to ascertain the strike direction and the lateral offset of the target, respectively.

Having the z and y component data increases the total response when the profile line has not traversed the target. This increases the possibility of detecting a target located between adjacent flight lines or beyond a survey boundary.

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## INTRODUCTION

The acquisition of multiple-component electromagnetic (EM) data is becoming more commonplace. In some techniques, such as those which use the plane-wave assumption (MT, CSAMT and VLF) more than one component has been acquired as a matter of routine for some time (see reviews by Vozoff, 1990, 1991; Zonge and Hughes, 1991; McNeill and Labson, 1991). Historically, commercially available controlled-waveform finite-source systems generally measure only one component. The only systems designed to acquire multiple component data are generally experimental [e.g., those described in the appendixes of Spies and Frischknecht (1991) or proprietary (the EMP system of Newmont Exploration)].

Slingram EM systems, comprising a moving dipolar transmitter and a moving receiver, generally only measure one component of the response. Although the MaxMin system was designed with a capability to measure a second (minimum coupled) component, this capability is not used extensively in practice. The only systems that use two receiver coils in practice are those that measure the wavetilt or polarization ellipse (Frischknecht et al., 1991).

Historically, time-domain EM systems have been capable of collecting multi-component data in a sequential manner by reorienting the sensor for each component direction. The usefulness of additional components is discussed by Macnae (1984) for the case of the UTEM system. Macnae concluded that, as extra time was required to acquire the additional components, this time was better spent collecting more densely spaced vertical-component data. The vertical-component, which is less subject to spheric noise, could subsequently be converted to the horizontal components using the Hilbert transform operators.

Recent instrument developments have been towards multi-component systems. For example, commercially available ground-EM systems such as the Geonics PROTEM, the Zonge GDP-32 and the SIROTEM have been expanded to include multiple input channels that allow three (or more) components to be acquired simultaneously. There is also a version of the UTEM system currently being developed at Lamontagne Geophysics Ltd. These multichannel receivers require complimentary multi-component sensors -- for ground-based systems these have been developed by Geonics Ltd and Zonge Engineering and Research Organization. The interpretation of fixed-source, multi-component ground-EM data is described in Barnett (1984) and Macnae (1984).

In the past, multi-component borehole measurements have been hindered by the lack of availability of multi-component sensor probes. Following the development of two prototype probes (Lee, 1986; Hodges et al., 1991), multi-component sensors are now available from Crone Geophysics and Exploration Ltd and Geonics. Three component UTEM and SIROTEM borehole sensors are also in development at Lamontagne and Monash University (Cull, 1993), respectively. Hodges et al. (1991) present an excellent discussion of techniques that can be used to interpret three-component borehole data.

Airborne systems such as frequency-domain helicopter electromagnetic methods acquire data using multiple sensors. However, each receiver has a corresponding transmitter that either operates at a different frequency or has a different coil orientation (Palacky and West, 1991). Hence, these systems are essentially multiple single-component systems. The exception to this rule is the now superseded Dighem III system (Fraser, 1972) which used one transmitter and three receivers.

The only multi-component airborne EM (AEM) system currently in operation is the SPECTREM system (Macnae, et al., 1991). This is a proprietary system (owned and operated by

Anglo-American Corporation of South Africa Ltd.), based on the PROSPECT system (Annan, 1986). The Prospect system was originally designed to acquire the  $x$ ,  $y$  and  $z$  components, but SPECTREM is apparently only collecting two components ( $x$  and  $z$ ) at the time of writing. Other multi-component systems currently in development are:

- 1) the SALTMAP system,
- 2) a helicopter time-domain system (Hogg, 1986), and
- 3) a new version of the GEOTEM<sup>®</sup> system (GEOTEM is a registered trademark of Geoterrex).

Apart from a few type curves in Hogg (1986), there is little literature available which describes how to interpret data from these systems.

This paper is intended to give an insight into the types of responses expected with the new multi-component AEM systems, and the information that can be extracted from the data. The insight could be of some assistance in interpreting data from multi-component moving-source ground EM systems (should this type of data be acquired).

The use of multi-component data will be discussed for a number of different applications. For illustration purposes, this paper will use the transmitter-receiver geometry of the GEOTEM<sup>®</sup> system (Figure 1), which is comparable to the other fixed-wing geometries (SPECTREM and SALTMAP). The GEOTEM<sup>®</sup> system is a digital transient EM system utilizing a bipolar half-sinusoidal current waveform [for more details refer to Annan and Lockwood (1991)]. The sign convention used in this paper is shown in Figure 1, with the  $y$  component being into the page. In a practical EM system, the receiver coils will rotate in flight. We will assume that the three components of the measured primary field and an assumed bird position have been used to correct for any rotation of the coil.

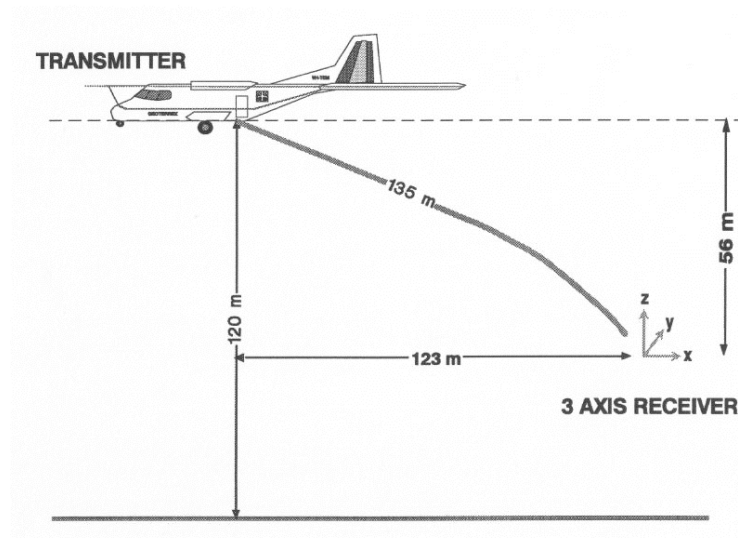


Fig. 1: The geometric configuration of the GEOTEM<sup>®</sup> system. The system comprises a transmitter on the aircraft and a receiver sensor in a “bird” towed behind the aircraft. The  $z$  direction is positive up,  $x$  is positive behind the aircraft, and  $y$  is into the page (forming a right-hand coordinate system).

## SOUNDING IN LAYERED ENVIRONMENTS

In a layered environment, the induced current flow is horizontal (Morrison et al., 1969) so the  $z$  component of the secondary response ( $V_z$ ) is much larger than the  $x$  component ( $V_x$ ), particularly in resistive ground and/or at late delay times. At the same time, the spheric noise in the  $z$  direction is 5 to 10 times less than in the horizontal directions (Macnae, 1984; McCracken et al., 1986), so  $V_z$  has a greater signal-to-noise ratio. Figure 2 shows theoretical curves over two different, but similar, layered earth models. One model is a half-space of  $500 \Omega \cdot m$  and the other is a  $350 \text{ m}$  thick layer of  $500 \Omega \cdot m$  overlying a highly resistive basement. In this plot the data have been normalized by the total primary field. The  $z$  component ( $V_z$ ) is 6 to 10 times larger than  $V_x$ , and both curves are above the noise level, at least for part of the measured transient. On this plot, a noise level of  $30 \text{ ppm}$  has been assumed, which would be a typical noise level for both components when the spheric activity is low. To distinguish between the response of the half-space and thick layer, the difference between the response of one model and the response of the other model must be greater than the noise level. Figure 3 shows this difference for both components. Only the  $V_z$  difference is above the noise level. Hence for the case shown,  $V_z$  is more useful than  $V_x$  for determining whether there is a resistive layer at  $350 \text{ m}$  depth. Because  $V_z$  is generally larger in a layered environment, the vertical component will generally be better at resolving the conductivity at depth.

In the above discussion, we have assumed that corrections have been made for the coil rotation. An alternative approach is to calculate and model the magnitude of the total field, as this quantity is independent of the receiver orientation. Macnae et al. (1991) used this strategy when calculating the conductivity depth sections for SPECTREM data.

The symmetry of the secondary field of a layered environment is such that the  $y$  component response ( $V_y$ ) will always be zero. In fact, the  $V_y$  component will be zero whenever the conductivity structure on both sides of the aircraft is the same. A non-zero  $V_y$  is therefore useful in identifying off-line lateral inhomogeneities in the ground.

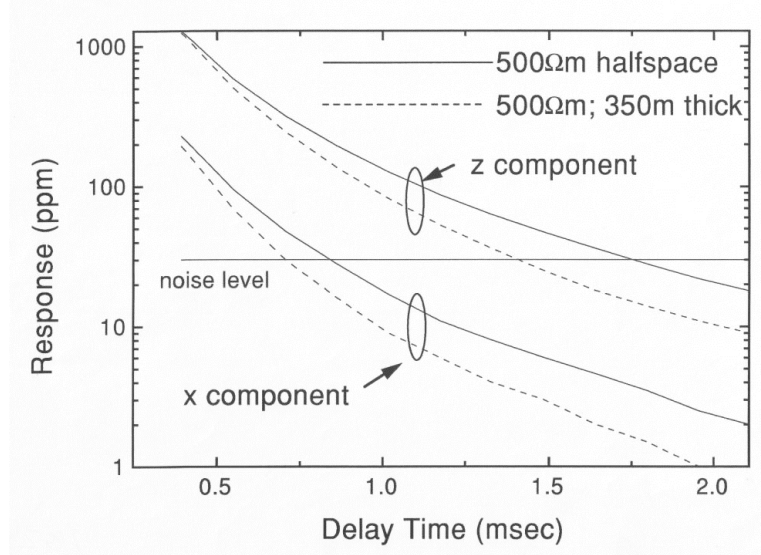


Fig. 2. The response for a  $500 \Omega \cdot m$  half-space (solid line) and a  $500 \Omega \cdot m$  layer of thickness  $350 \text{ m}$  overlying a resistive half-space (dashed line). The  $z$ -component responses are the two curves with the larger amplitudes and the two  $x$ -component response curves are 6 to 10 times smaller than the corresponding  $z$  component. A noise level of  $30 \text{ ppm}$  is considered to be typical of both components in the absence of strong spherics.

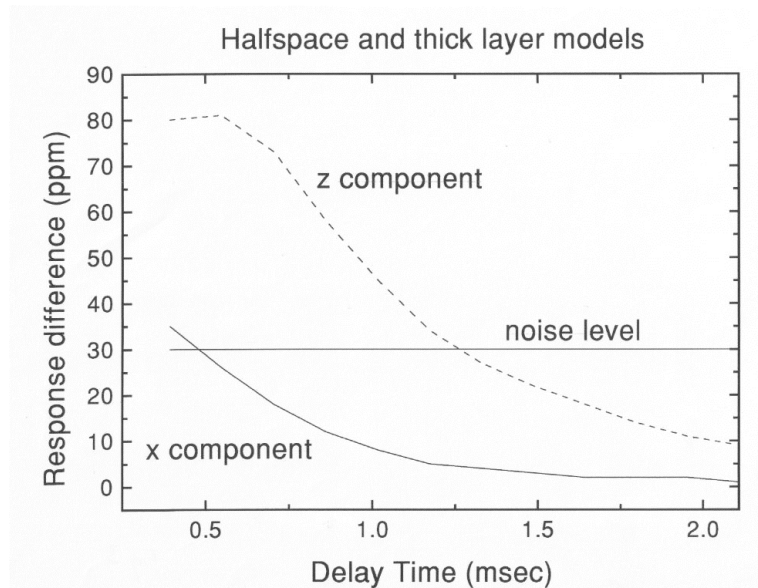


Fig. 3: The difference in the response of each component for the half-space and thick layer models of Figure 2. Only the z-component difference is above the noise level for a significant portion of the transient. Therefore, this is the only component capable of distinguishing between the responses of the two models.

## DISCRETE CONDUCTORS

In our discrete conductor study, models have been calculated using a simple plate in free-space model (Dyck and West, 1984) to provide some insight into the geometry of the induced field. The extension to more complex models, such as those incorporating current gathering, will not be considered in this paper.

Historically, airborne transient electromagnetic (TEM) data have been used for conductor detection. The old INPUT system was designed to measure  $V_x$  because this component gave a large response when the receiver passed over the top of a vertical conductor. The bottom part of Figure 4 shows the response over a vertical conductor, which has been plotted at the receiver position. The  $V_x$  profile (smaller of the two solid lines) has a large peak corresponding with the conductor position. Note that there is also a peak at 200 m, just before the transmitter passes over the conductor, and a trailing edge negative to the left of the conductor. The z component (dashed line) has two peaks and a large negative trough just before the conductor. Because of the symmetry, the  $V_y$  response (dotted line) is zero.

All the peaks, troughs and negatives make the response of a single conductor complicated to display and hence interpret. The display can be simplified by plotting the "energy envelope" (EE) of the response. This quantity is defined as follows:

$$EE = \sqrt{V_x^2 + \bar{V}_x^2 + V_y^2 + \bar{V}_y^2 + V_z^2 + \bar{V}_z^2},$$

where  $\bar{\phantom{x}}$  denotes the Hilbert transform of the quantity. The energy envelope plotted on Figure 4 (the larger of the two solid curves) is almost symmetric, and would be a good quantity to present in plan form (as contours or as an image). For flat-lying conductors, the energy envelope has a maximum at the leading edge (just after the aircraft flies onto the conductor).



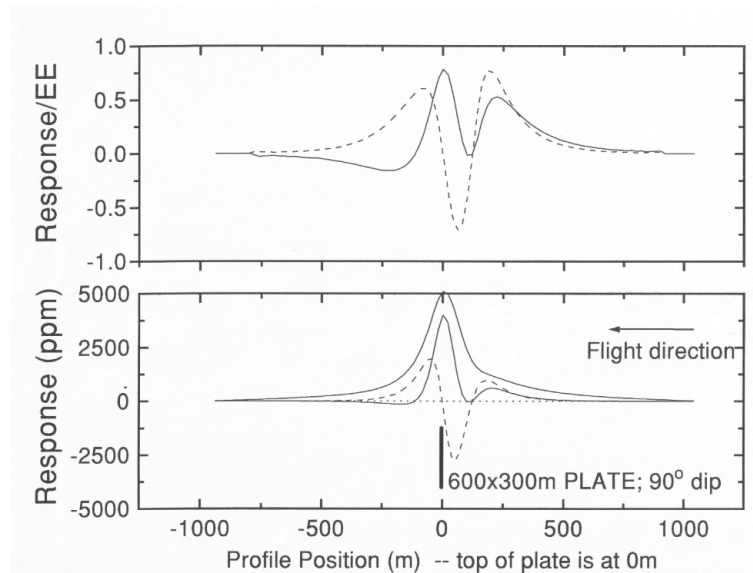


Fig. 4. (Bottom) the response of a 600 by 300 m plate 120 m below an aircraft flying from right to left. The plotting point for the response is below the receiver. The x-component response is the smaller amplitude solid line, the z-component is the dashed line, and the y-component response is the dotted line. The larger amplitude solid line is the “energy envelope” of all three components. (Top) the z and x-components normalized by the energy envelope. These and all subsequent curves are for a delay time of 0.4 ms after the transmitter current is turned off.

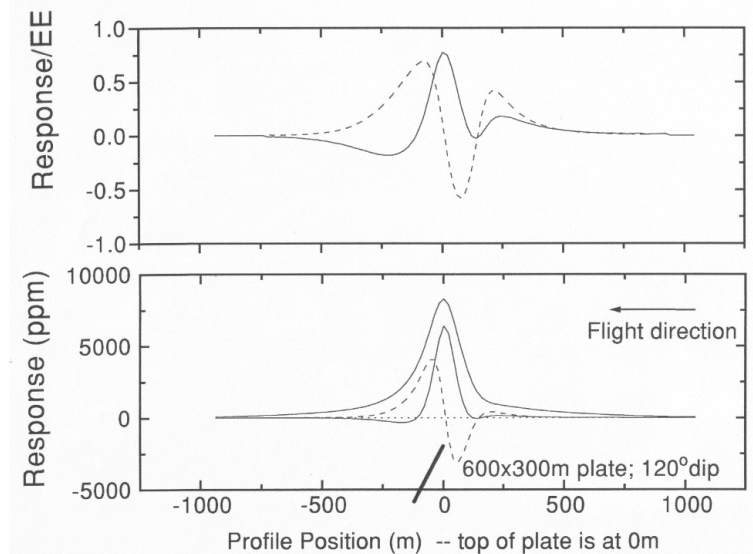


Fig. 5 (Bottom) same as Figure 4, except the plate is now dipping at 120°. On the top graph note the down-dip (left) peak on the normalized z-component response is larger than the right peak (c.f. Figure 4).

What little asymmetry remains in the energy envelope is a good indication of the coupling of the AEM system to the conductor. If the response profile for each component is normalized by the energy envelope, then the effect of system coupling will be removed (at least partially) and the profiles will appear more symmetric. For example, the top part of Figure 4 shows the  $V_x$  and  $V_z$  normalized by the energy envelope at each point. The size of the two x peaks and the two z peaks are now roughly comparable.

## Dip determination

The response of a plate with a dip of  $120^\circ$  is shown on Figure 5. For the  $V_x/EE$  and  $V_z/EE$  profiles, the peak on the down dip side is larger. For shallow dips, it becomes difficult to identify both  $V_x/EE$  peaks, but the two positive  $V_z/EE$  peaks remain discernable. Plotting the ratio of the magnitudes of these two  $V_z/EE$  peaks, as has been done with solid squares on Figure 6, shows that the ratio is very close to the tangent of the dip divided by 2. Hence, calculating the ratio of the peak amplitudes ( $R$ ) will yield the dip angle  $\theta$  using

$$\theta = 2 \tan^{-1}(R).$$

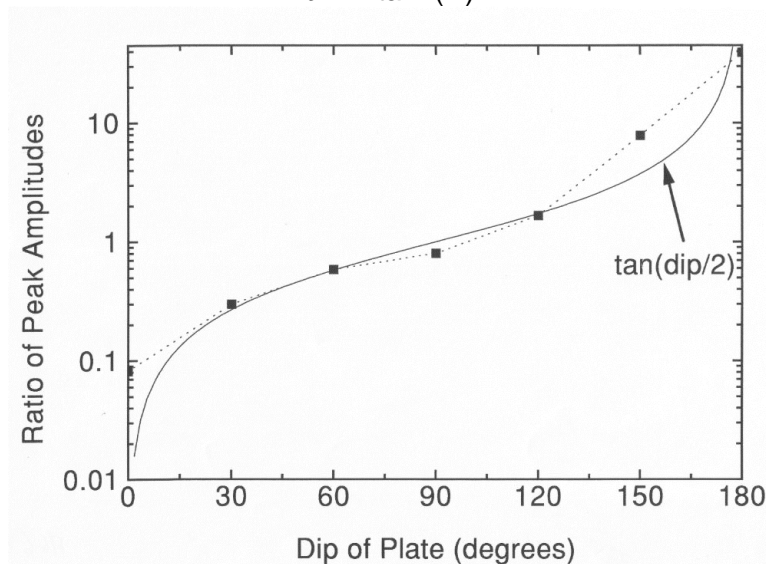


Fig. 6. The ratio of the peak amplitudes of the normalized z-component response (left/right) plotted with solid squares. The ratio plots very close to the tangent of half the dip angle  $\theta$  of the plate.

## Depth Determination

As the depth of the body increases, there is a corresponding increase in the distance between the two positive peaks in the  $V_z/EE$  profile. As an example of this, Figure 7 shows the case of a plate 150 m deeper than the plate of Figure 4. The peaks are now 450 m apart, as compared with 275 m on Figure 4. A plot of the peak-to-peak distances for a range of depths is shown on Figure 8 for plates with  $60^\circ$ ,  $90^\circ$  and  $120^\circ$  dips. Because the points follow a straight line, it can be concluded that for near vertical bodies ( $60^\circ$  to  $120^\circ$  dips), the depth to the top of the body  $d$  can be determined from the measured peak-to-peak distances using the linear relationship depicted in Figure 8. The expected error would be about 25 m. Such an error is tolerable in airborne EM interpretation. More traditional methods for determining  $d$  analyze the rate of decay of the measured response (Palacky and West, 1973). Our method requires only the  $V_z/EE$  response profile at a single delay time. Analyzing this response profile for each delay time allows  $d$  to be determined as a function of delay time, and hence any migration of the current system in the conductor could be tracked.



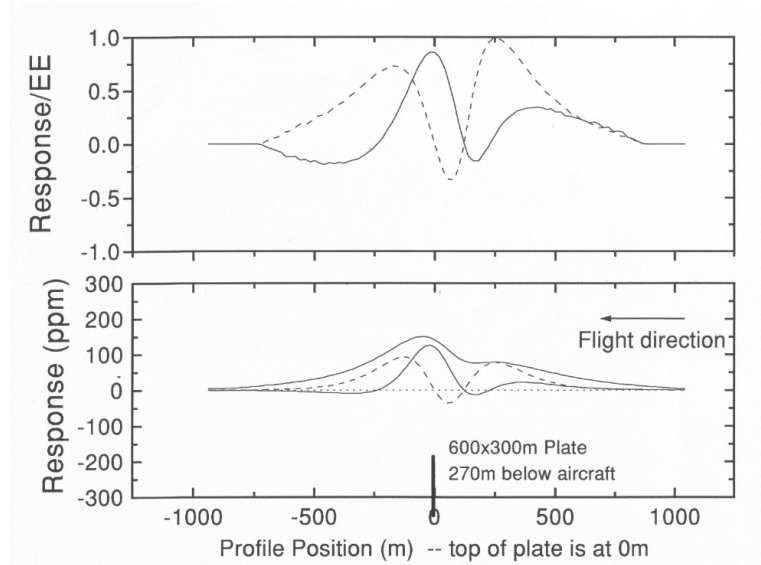


Fig. 7. The same as Figure 4, except the plate is now 270 m below the aircraft. Note that the distance between the z-component peaks is now much greater.

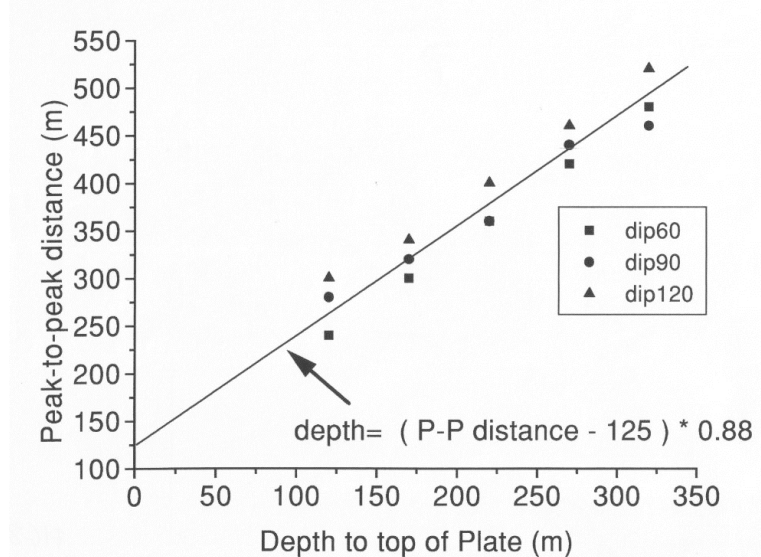


Fig. 8. The peak-to-peak distance as a function of plate depth for three different dip angles  $\theta$ . A variation in dip of  $\pm 30^\circ$  does not result in a large change in the peak to peak distance.

### Strike and offset determination

The response shown in Figure 4 varies in cases when the plate has a strike different from  $90^\circ$  or the flight path is offset from the center of the plate.

Figure 9 shows the response for a plate with zero offset and Figure 10 shows the plate when it is offset by 150 m from the profile line. The calculated voltages  $V_z$  and  $V_x$  are little changed from the no offset case, but the  $V_y$  response, is no longer zero. In fact, the shape of the  $V_y$  curve appears to be the mirror image of the  $V_z$  curve.

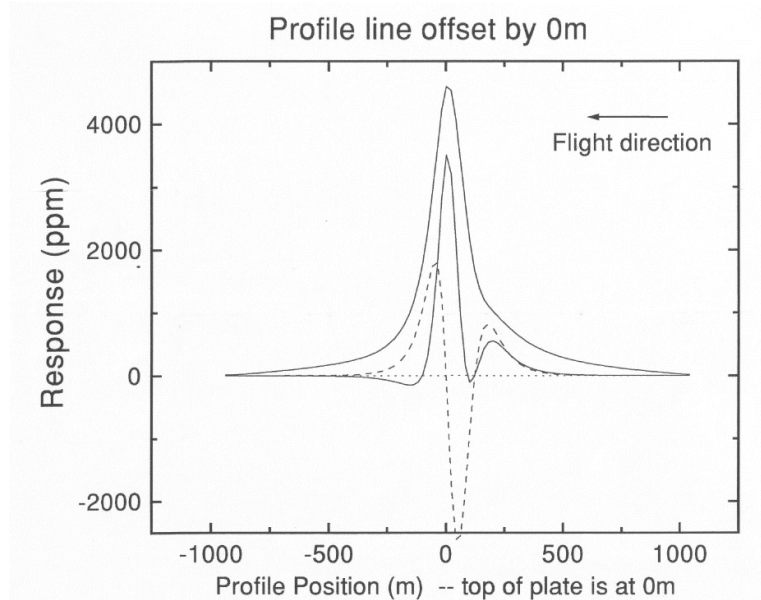


Fig. 9 The response of a 300 by 300 m plate traversed by a profile line crossing the center of the plate in a direction perpendicular to the strike of the plate (the strike angle  $\zeta$  of the plate with respect to the profile line is  $90^\circ$ ).

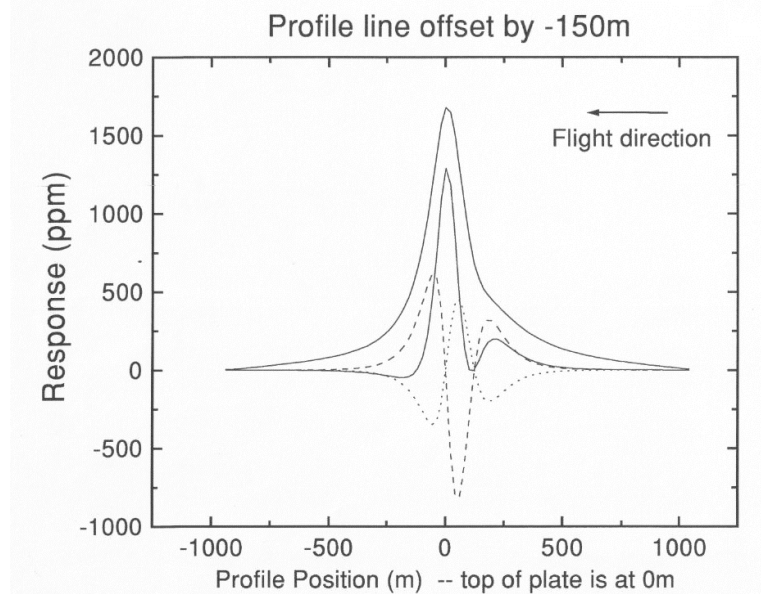


Fig. 10. Same as Figure 9, except the profile line has been offset from the center of the plate by  $-150$  m in the  $y$  direction (equivalent to a  $+150$  m displacement of the plate).

In the case when the plate strikes at  $45^\circ$ , the  $y$  component is similar in shape but opposite in sign to the  $x$  component response (Figure 11).

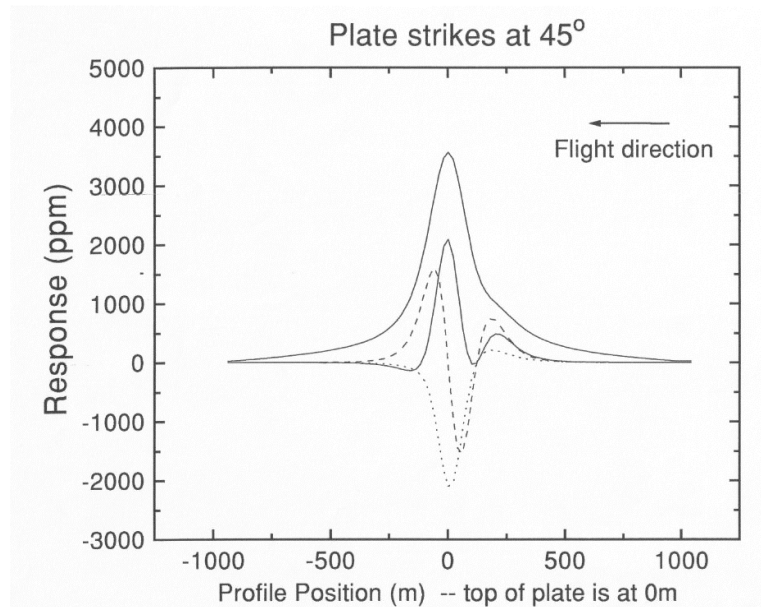


Fig. 11 Same as Figure 9, except the profile line traverses the plate such that the strike angle  $\zeta$  of the plate with respect to the profile line is  $45^\circ$ .

These similarities can be better understood by looking at schematic diagrams of the secondary field from the plate. Figure 12 shows a plate and the field in section. For zero offsets, the field is vertical (z only). As the offset increases, the aircraft and receiver moves to the right and the measured field rotates into the y component.

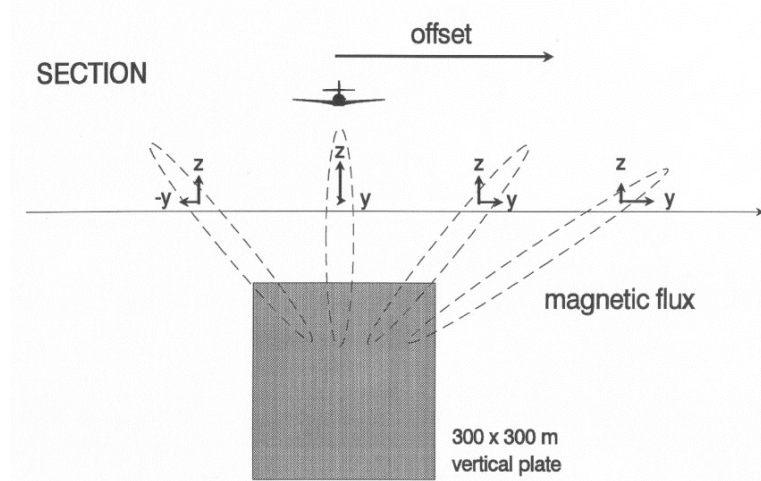


Fig. 12. A schematic diagram of the plate and the magnetic flux of the secondary field (section view). For increasing offset of the aircraft and receiver from the center of the plate, the magnetic field at the receiver rotates from the z to the y component.

The secondary field is depicted in plan view in Figure 13. Variable strike is simulated by leaving the plate stationary and changing the flight direction. When the strike of the plate is different from  $90^\circ$ , the effective rotation of the EM system means that the secondary field, which was previously measured purely in the x direction, is now also measured in the y direction.

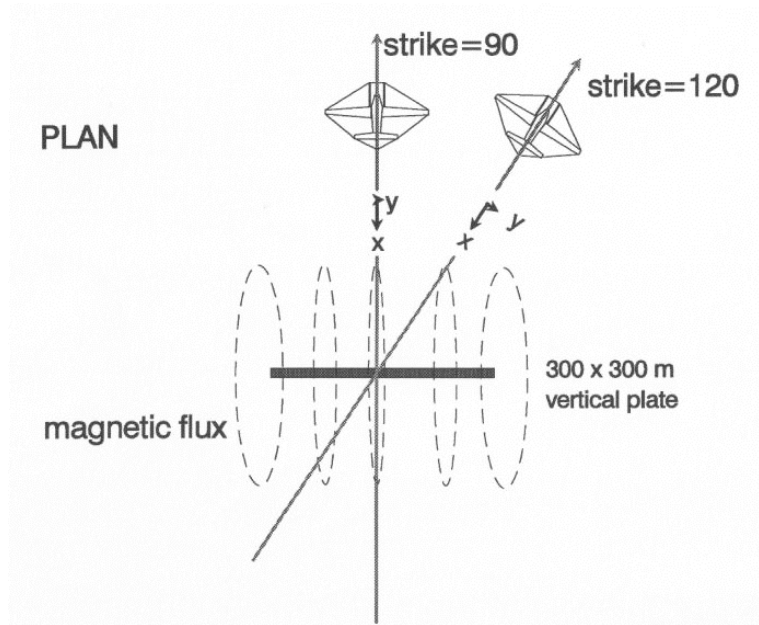


Fig. 13. A schematic diagram of the plate and the magnetic flux of the secondary field (plan view). Here varying strike is depicted by an equivalent variation of the flight direction. As the flight direction rotates from a strike angle of 90°, the receiver rotates so as to measure a greater response in the y direction.

The y component ( $V_y$ ) can thus be considered to be a mixture of  $V_x$  and  $V_z$  components, viz

$$V_y = C_{stk} V_x + C_{off} V_z ,$$

an equation that is only approximate. The response for a variety of strike angles and offset distances has been calculated and in each case the y component response has been decomposed into the x and z components by solving for the constants of proportionality  $C_{stk}$  and  $C_{off}$ .

A plot of  $C_{stk}$  for the case of zero offset and varying strike direction  $\xi$  is seen on Figure 14. The values of  $C_{stk}$  determined from the data are plotted with solid squares and compared with the  $\tan(90^\circ - \xi)$ . Because the agreement is so good, the formula

$$\xi = 90 - \tan^{-1} (C_{stk})$$

can be used to determine the strike. This relation was first obtained by Fraser (1972).



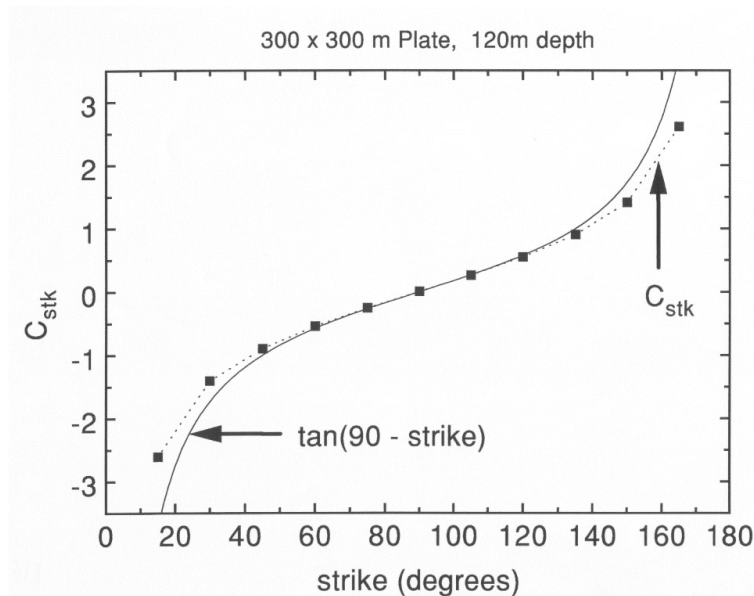


Fig. 14. The ratio  $C_{stk} = V_y/V_x$  plotted as a function of varying strike angle (solid squares). The data agree very closely with the cotangent of the  $\zeta$ .

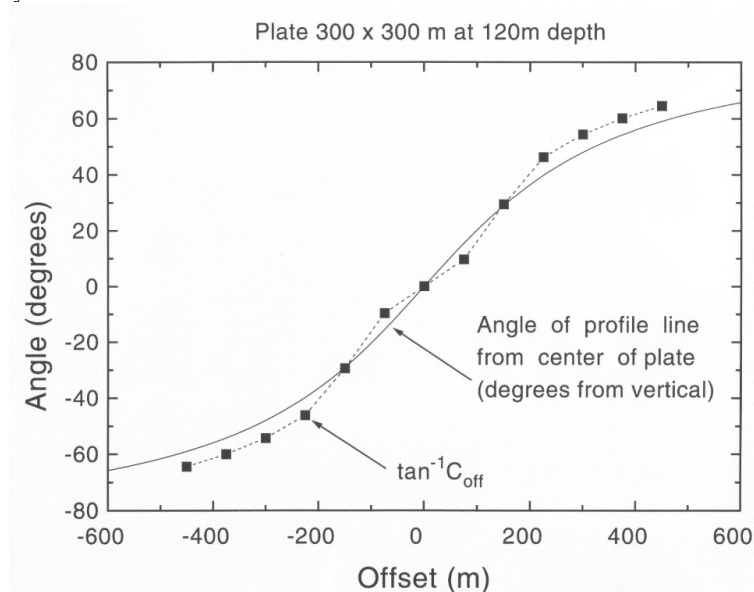


Fig. 15. The arctangent of  $C_{off} = V_y/V_z$ , plotted as a function of varying offset (solid squares). There is good agreement between this quantity and the angle  $\phi$  between a vertical line and the line from the center of the top edge of the plate to the profile line.

When the strike is fixed at  $90^\circ$ , and the offset varies, the corresponding values obtained for  $C_{off}$  have been plotted with solid squares on Figure 15. Again, there is good agreement with the arctangent of  $C_{off}$  and the angle  $\phi$  between a vertical line and the line that joins the center of the top edge of the plate with the position where the aircraft traverse crosses the plane containing the plate. If an estimate of the distance to the top of the conductor  $D$  is already obtained using the method described above, or by the method described in Palacky and West (1973), then

$$D = \sqrt{(O^2 + d^2)},$$

(where  $d$  is the depth below surface). Hence, the offset distance  $O$  can be written as follows

$$\begin{aligned} O &= d \tan(\phi) \\ &= d C_{\text{off}} \\ &= C_{\text{off}} \sqrt{(D^2 - O^2)} \end{aligned}$$

which can be rearranged to give

$$O = C_{\text{off}} D / \sqrt{(1 + C_{\text{off}}^2)}.$$

### Lateral delectability

Figure 12 illustrates that  $V_y$  becomes relatively strong as the lateral displacement from the conductor is increased. Thus, if  $V_y$  is measured, then the total signal will remain above the noise level at larger lateral displacements of the traverse line from the conductor. This has been illustrated by assuming a flat-lying conductor, here approximated by a wire-loop circuit of radius 125 m (Figure 16). The  $x$ ,  $y$  and  $z$  components of the response have been computed using the formula for the large-loop magnetic fields in Wait (1982). The results are plotted on Figure 17 as a function of increasing lateral displacement  $L$  of the transmitter/receiver from the center of the conductor. The transmitter and receiver are separated in a direction perpendicular  $L$  to simulate the case when the system is maximal coupled to the conductor, but the flight line misses the target by an increasing amount. The effect of varying the conductance or measurement time has been removed by normalizing the response to the total response measured when the system is at zero displacement. At displacements greater than 80 m, the  $y$  component is clearly larger than any other component. Assuming the same sensitivity and noise level for each component (which is a realistic assumption if the data are corrected for coil rotation and the spheric activity is low), it is clearly an advantage to measure  $V_y$ , as this will increase the chances of detecting the target when the flight line has not passed directly over the conductor.

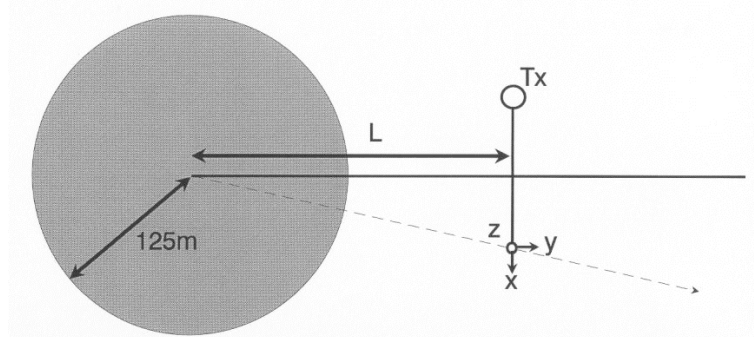


Fig. 16. Plan view of a flat-lying conductor (a circular loop with a radius of 125 m). The AEM system is offset a distance  $L$  from the center of the conductor in a direction perpendicular to the traverse direction. The traverse direction of the system is from the bottom to the top of the figure.

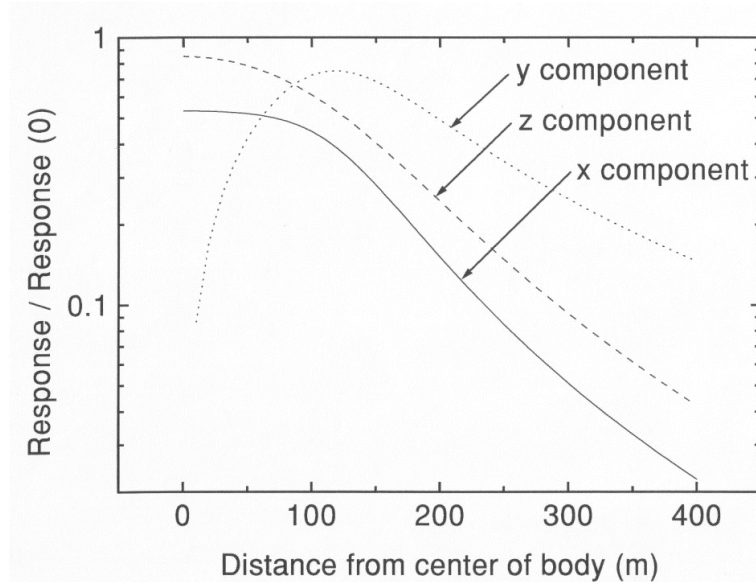


Fig. 17. The normalized response of the EM system plotted as a function of increasing offset distance  $L$ . The  $x$  component falls off most rapidly and the  $y$  component most slowly with increasing offset distance.

## CONCLUSIONS

AEM systems measuring three components of the response can be used to infer more and/or better information than those systems that measure with only one component, i.e.,  $V_x$ .

The  $z$  component data enhances the ability of the AEM system to resolve layered structures as the  $z$  component has a larger signal and a smaller proportion of spheric noise than any other component. If all the components are employed to correct for coil rotation, then the data quality and resolving power is increased further, as individual components are not contaminated by another component. Having better signal-to-noise and greater fidelity in the data will allow deeper layers to be interpreted with confidence.

A non-zero  $y$  component is helpful in identifying when the conductivity structure has a lateral inhomogeneity that is not symmetric about the flight line.

All components can be used to calculate the energy envelope, which is a valuable quantity to image. The energy envelope has a single peak over a vertical conductor and two peaks over a dipping conductor (one at either end). The asymmetry in the response profile of each individual component can be reduced by normalizing each profile by the energy envelope.

All three components are of great use in determining the characteristics of discrete conductors. For example, the distance between the two positive peaks in the  $V_z/EE$  profile can be employed to determine the depth. Also, the ratio of the magnitude of the two  $V_z/EE$  peaks helps to ascertain the dip of the conductor. The  $x$  component has been used in the past for these purposes, but is not as versatile, as it requires the data at all delay times, or an ability to identify a very small peak.

The  $y$ -component can be utilized to extract information about the conductor that cannot be

obtained from single component AEM data. The degree of mixing between the y and z components can give the lateral offset of the conductor (provided the depth is known), while the mixing between the y and x component gives the strike of a vertical conductor.

Finally, because the y component decreases most slowly with increasing lateral offset, this component gives an enhanced ability to detect a conductor positioned at relatively large lateral distances from the profile line, either between lines or beyond the edge of a survey boundary.

### ACKNOWLEDGMENTS

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# **Appendix E**

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## Data Archive Description

## Data Archive Description:

### Survey Details

Survey Area Name:	Renabie
Project number:	13416
Client:	Conquest Resources Limited
Survey Company Name:	Fugro Airborne Surveys
Flown Dates:	March 15, 2013
Archive Creation Date:	March, 2013

### Survey Specifications

Traverse Line Azimuth:	004°-184°
Traverse Line Spacing:	200 m
Tie Line Azimuth:	094°-274°
Tie Line Spacing:	1000 m
Flying Elevation:	120 m Mean Terrain Clearance
Average Aircraft Speed:	65 m/s

### Geodetic Information for map products

Projection:	Universal Transverse Mercator
Datum:	NAD83
Central meridian:	81° West
False Easting:	500000 metres
False Northing:	0 metres
Scale factor:	0.9996
UTM Zone:	17 North
I.G.R.F. Model:	2010
I.G.R.F. Correction Date:	2013/03/15

## Equipment Specifications:

### Navigation

GPS Receiver:	NovAtel OEM4, 12 Channels
Aircraft:	Casa (Twin Turbo Propeller)
Video Camera:	Digital video

### Magnetics

Type:	Scintrex CS-3 Caesium Vapour
Installation:	Towed bird
Sensitivity:	0.01 nT
Sampling:	0.1 s

## Electromagnetics

Type: GEOTEM®, 30 channel multi-coil system  
 Installation: Vertical axis loop (231m<sup>2</sup> area with 6 turns) mounted on aircraft  
 Receiver coils in a towed bird.  
 Coil Orientation: X, Y and Z  
 Frequency: 90 Hz  
 Pulse width: 2127µs  
 Off-time: 3407µs  
 Geometry: Tx-Rx horizontal separation of ~130 m  
 Tx-Rx vertical separation of ~45 m  
 Sampling: 0.25 s

## Data Windows:

Channel	Start (p)	End (p)	Width (p)	Start (ms)	End (ms)	Width (ms)	Mid (ms)
1	3	13	11	0.022	0.141	0.119	0.081
2	14	73	60	0.141	0.792	0.651	0.467
3	74	132	59	0.792	1.432	0.640	1.112
4	133	192	60	1.432	2.083	0.651	1.758
5	193	203	11	2.083	2.203	0.119	2.143
6	204	205	2	2.203	2.224	0.022	2.214
7	206	207	2	2.224	2.246	0.022	2.235
8	208	209	2	2.246	2.268	0.022	2.257
9	210	211	2	2.268	2.289	0.022	2.279
10	212	214	3	2.289	2.322	0.033	2.306
11	215	217	3	2.322	2.355	0.033	2.338
12	218	220	3	2.355	2.387	0.033	2.371
13	221	224	4	2.387	2.431	0.043	2.409
14	225	228	4	2.431	2.474	0.043	2.452
15	229	233	5	2.474	2.528	0.054	2.501
16	234	239	6	2.528	2.593	0.065	2.561
17	240	246	7	2.593	2.669	0.076	2.631
18	247	254	8	2.669	2.756	0.087	2.713
19	255	264	10	2.756	2.865	0.109	2.810
20	265	276	12	2.865	2.995	0.130	2.930
21	277	290	14	2.995	3.147	0.152	3.071
22	291	306	16	3.147	3.320	0.174	3.234
23	307	324	18	3.320	3.516	0.195	3.418
24	325	344	20	3.516	3.733	0.217	3.624
25	345	366	22	3.733	3.971	0.239	3.852
26	367	390	24	3.971	4.232	0.260	4.102
27	391	416	26	4.232	4.514	0.282	4.373
28	417	444	28	4.514	4.818	0.304	4.666
29	445	474	30	4.818	5.143	0.326	4.980
30	475	512	38	5.143	5.556	0.412	5.349

### ASCII and Geosoft Line Archive File Layout (AREA\_ascii.xyz & AREA.gdb):

Field	Variable	Description	Units
1	line	Line Number	line+part
2	fiducial	Universal Time (Seconds Since Midnight)	s
3	flight	Flight Number	-
4	date	Date of Survey Flight	yyyymmdd
5	lat_nad83	Latitude in NAD83	degrees
6	long_nad83	Longitude in NAD83	degrees
7	x_nad83	Easting (X) in NAD83 UTM Zone 17N	m
8	y_nad83	Northing (Y) in NAD83 UTM Zone 17N	m
9	gpsz	GPS Elevation (Referenced to Mean Sea Level)	m
10	alt_radar	Radar Altimeter	m
11	dtm	Terrain (Referenced to Mean Sea Level)	m
12	diurnal	Ground Magnetic Intensity (Diurnal)	nT
13	mag_raw	Raw Total Magnetic Intensity (TMI)	nT
14	igrf	International Geomagnetic Reference Field	nT
15	mag_rmi	Residual Magnetic Intensity (Levelled)	nT
16	mag_tmi	Total Magnetic Intensity (Levelled)	nT
17	vd1	First Vertical Derivative of Residual Magnetic Intensity	nT/m
18	primary_field	Electromagnetic X-Coil Primary Field	μV
19	hz_monitor	Powerline Monitor (60 Hz)	μV
20-49	emx_db	Final dB/dt X-Coil Channels 01 to 30	nT/s
50-79	emy_db	Final dB/dt Y-Coil Channels 01 to 30	nT/s
80-109	emz_db	Final dB/dt Z-Coil Channels 01 to 30	nT/s
110-139	emx_bf	Final B-Field X-Coil Channels 01 to 30	pT
140-169	emy_bf	Final B-Field Y-Coil Channels 01 to 30	pT
170-199	emz_bf	Final B-Field Z-Coil Channels 01 to 30	pT
200-229	emx_db_raw	Raw dB/dt X-Coil Channels 01 to 30	nT/s
230-259	emy_db_raw	Raw dB/dt Y-Coil Channels 01 to 30	nT/s
260-289	emz_db_raw	Raw dB/dt Z-Coil Channels 01 to 30	nT/s
290-319	emx_bf_raw	Raw B-Field X-Coil Channels 01 to 30	pT
320-349	emy_bf_raw	Raw B-Field Y-Coil Channels 01 to 30	pT
350-379	emz_bf_raw	Raw B-Field Z-Coil Channels 01 to 30	pT
380	cond_z01	Apparent Conductance Derived from dB/dt Z-Coil Channel 01	mS
381	tauBx12_30	Decay Constant (Tau) Derived from B-Field X-Coil Channels 12 to 30	μs
382	tee08	Total Energy Envelope Derived from dB/dt Channel 08	nT/s

Note – The null values in the ASCII archive are displayed as \*

### Grid Archive File Description:

The grids are in Geosoft format. A grid cell size of 50 m was used for all area grids.

File	Description	Units
Renabie_rmi.grd	Residual Magnetic Intensity	nT
Renabie_vd1.grd	First Vertical Derivative of Residual Magnetic Intensity	nT/m
Renabie_cond_z01(_deh).grd	Apparent Conductance Derived from dB/dt Z-Coil Channel 01	mS
Renabie_tauBx12_30(_deh).grd	Decay Constant (Tau) Derived from B-Field X-Coil Channels 12 to 30	$\mu$ s
Renabie_tee08(_deh).grd	Total Energy Envelope Derived from dB/dt Channel 08	nT/s

The \*\_deh files are the grid files corrected for asymmetry (“de-herringboned”).

## **Appendix F**

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### Map Product Grids

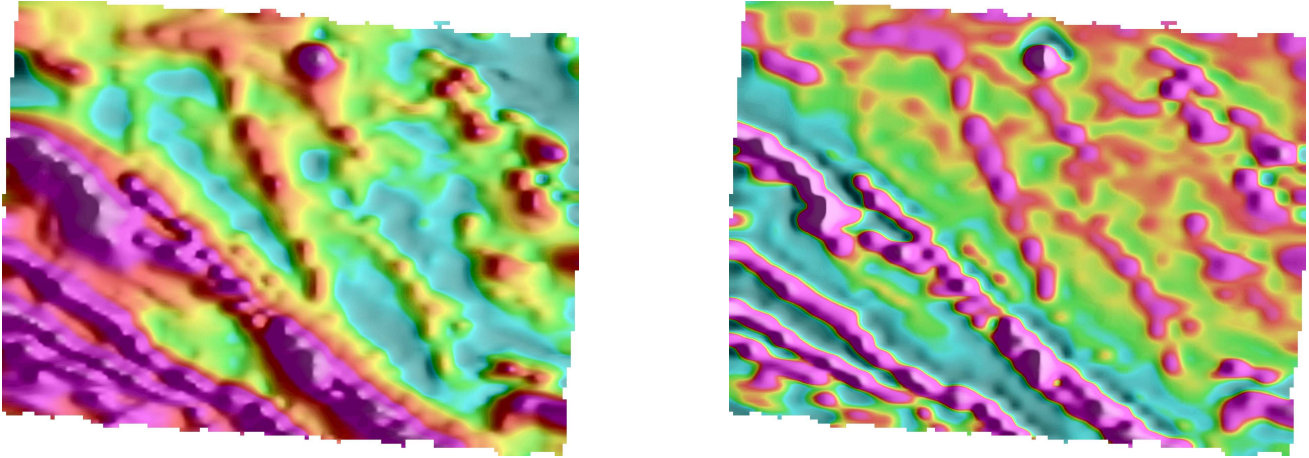


Figure 1. Residual Magnetic Intensity (left), First Vertical Derivative of RMI (right)

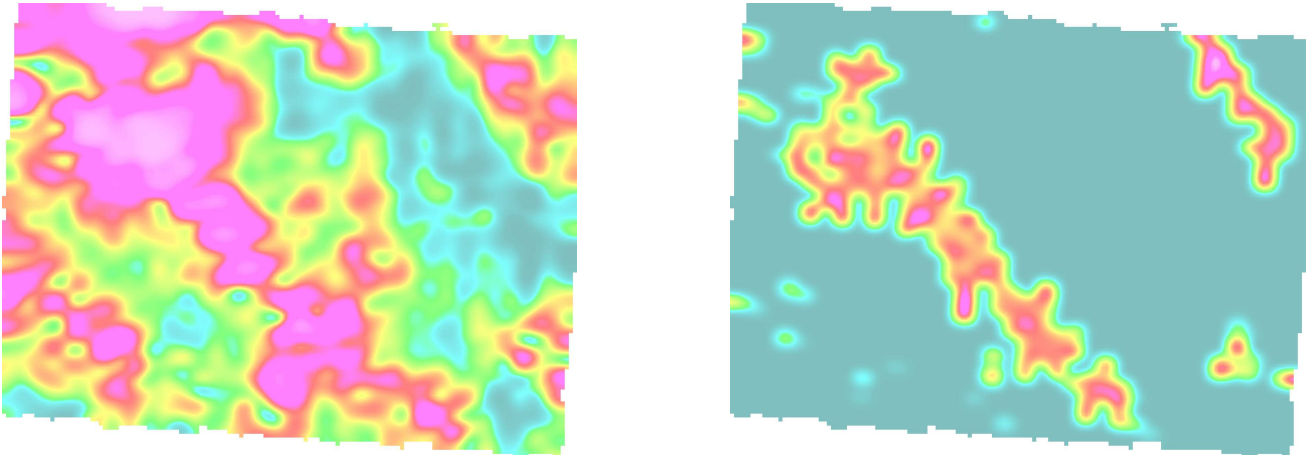


Figure 2. Apparent Conductance (left) and Decay Constant (right)

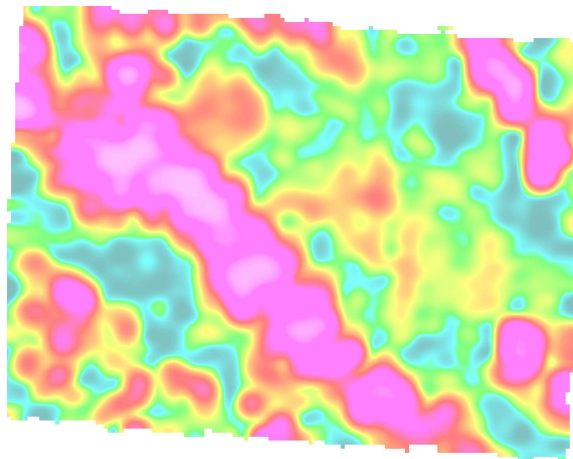


Figure 3. Total Energy Envelope



## **Appendix G**

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### Reference Waveform

## Reference Waveform Descriptor:

The information shown is only an example. The actual reference waveforms are provided on CD-ROM or DVD and will have been renamed to ptaFLTpre.out / ptaFLTpost.out, "FLT" represents the flight number.

The reference waveform can be divided into four main sections, which are described below.

### Section 1

This section contains the name of the raw reference waveform file (i.e. **D0050704.002**). The approximate horizontal and vertical offsets (i.e. **125** m and **50** m) of the EM bird position in meters are listed. These are followed by the base frequency (i.e. **90**Hz) in Hertz and the sample interval (i.e. **8.14**  $\mu$ s) in microseconds.

```
GEOTEM Calibration Data - Version 31 July 1998
'D0050704.002' = Name of original saved parameter table file
125.000000000000000 = Horizontal TX-RX separation in meters
50.000000000000000 = Vertical TX-RX separation in meters
90.000000000000000 = Base Frequency in Hertz
8.1380208 = Sample Interval in micro-seconds
```

### Section 2

This section displays the gate configuration for channels 1 to 30.

30 Time Gates: First and Last Sample number, RMS chart position:


Start & end samples of each channel

1	4	18	1
2	19	176	2
3	177	333	3
4	334	491	4
5	492	507	5
6	508	512	6
7	513	518	7
8	519	526	8
9	527	536	9
10	537	548	10
11	549	562	11
12	563	578	12
13	579	596	13
14	597	616	14
15	617	641	15
16	642	671	16
17	672	706	17
18	707	746	18
19	747	791	19
20	792	841	20
21	842	901	21
22	902	971	22
23	972	1051	23
24	1052	1141	24
25	1142	1241	25
26	1242	1361	26
27	1362	1501	27
28	1502	1661	28
29	1662	1841	29
30	1842	2048	30

Channels 1 to 30

### Section 3

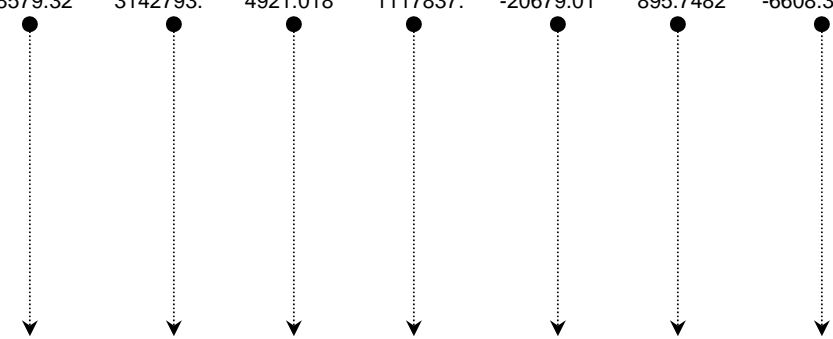
This section contains the different types of conversion factors for each of the components. If the data are provided in ppm the standard procedure is to normalize the data based on the individual components. Three different conversion factors are provided. The first factor converts the data to ppm based on the peak voltages of each individual component. The second factor converts the data to ppm based on the “total” peak voltage, which is actually the RMS value of the 3 components. The third factor converts each component to standard SI units, which are Teslas per second for the dB/dt data and Teslas for the B-field data.

Component:						
IndivPPM_per_DataUnit:	0.1112428E-01	1.106797	0.2890714E-01	0.1519028E-01	1.670836	0.3945841E-01
TotalPPM_per_DataUnit:	0.1038160E-01	0.1038160E-01	0.1038160E-01	0.1417559E-01	0.1417559E-01	0.1417559E-01
SI_Units_per_DataUnit:	0.1000000E-11	0.1000000E-11	0.1000000E-11	0.1000000E-14	0.1000000E-14	0.1000000E-14

### Section 4

The last section contains the reference waveform. Each column represents a component (i.e. dBx/dt). The data units (i.e. pT/s) for each component are displayed in the second row. The first column is the sample number. The transmitter channel (TX) values have been converted to transmitter moment value (transmitter current x loop area x number of turns)

For this example there are 2048 samples.

Component:	TX	dBx/dt	dBy/dt	dBz/dt	Bx	By	Bz
DataUnits:	Am <sup>2</sup>	pT/s	pT/s	pT/s	fT	fT	fT
2048 Samples:							
1	1161.572	38526.84	924.7899	14929.63	-164386.1	447.8830	-57789.91
2	1182.870	37266.81	431.6681	14992.79	-162768.6	466.6186	-57139.19
3	2471.644	130950.8	4966.128	46394.08	-157085.0	682.1624	-55125.55
4	18579.32	3142793.	4921.018	1117837.	-20679.01	895.7482	-6608.309
							
2045	-1089.137	-41458.52	-304.3134	-15670.09	166886.3	-413.2204	58736.67
2046	-1000.563	-40324.45	-44.37050	-14808.83	165136.1	-415.1462	58093.93
2047	-1006.226	-39601.37	61.07057	-14660.24	163417.3	-412.4956	57457.63
2048	-1134.029	-39423.19	-546.7703	-15091.65	161706.2	-436.2269	56802.61

## Appendix H

---

### TDEM Anomaly Selection

## Current approach to TDEM anomaly selection

The current routine for the selection and fitting of EM anomalies is still based on the University of Toronto plate program which fits the response (at the anomaly peak) from the X-coil channels to a vertical plate nomogram. Given that the current GEOTEM<sup>®</sup>, TEMPEST<sup>®</sup>, MEGATEM<sup>®</sup>, and HeliTEM<sup>®</sup> systems have evolved to offer the response from coils of 3 different orientations (X, Y and Z) and from dB/dt and B-Field, this approach to the classification of the anomalies is limited and no longer fully reflects all the information being measured by the system. The resulting shortcomings are:

- All anomaly peaks, from the X-coil response, are fitted to a vertical plate model of fixed dimensions, regardless of the nature of the conductive source. The derived CTP and depth-to-source values are then only valid if the conductor can be properly represented by a vertical plate. CTP and depth values derived from “non vertical plate” type conductors will be erroneous. In some conductive terrains, marked by prominent conductive overburden or surface alteration of other sorts, “non-vertical” type conductors may represent 90 % or more of the conductive response.
- The response from the conductor must deflect a minimum of 6 channels above the background to be fitted to the nomogram. CTP or depth-to-source values will not be calculated for a valid but weaker response (a weak or deep source).
- Only the response from the X-coil channels are used in the selection and fitting. A response appearing only on the Z-coil will not be identified. This is sometimes the case for a very deep source. As the depth to the conductor increases, although it may have considerable depth extent, the system becomes less sensitive to the vertical extent of the body and conductors will appear to be more flat-lying than vertical. As a result, as depth of burial increases, the coupling of the response may disappear on the X-coil response but will persist on the Z-coil response (see figure 1, anomalies A and B).
- The fitting of the response is only done from the amplitudes at the peak position of the anomaly and does not take-in the full shape of the anomaly or relate the difference in response between the X-coil and Z-coil. This does not allow for the distinction between vertical, flat-lying or dipping plates.

Fugro is presently working on the development of a new anomaly selection and classification routine which will use a window of data centered about the anomaly peak (to properly define the entire anomaly shape), using both the response from the X and Z-coils and fitting to a suite of models from flat-lying to dipping to vertical plates and spheres. This will hopefully address all the above shortcomings of the present method.

Unfortunately the current anomaly fitting program must continue to be in use until this new routine is made available. Until such time, our approach is to present the full information being measured by the system within the confines of the present program’s limitations. Some responses may not be visible on the regular channels of the dB/dt X-coil data but will be identified on either the B-Field response or the Z-coil response. The initial selection of the anomalies is still being derived from the X-coil channels of the dB/dt data but at the “review” stage (via a graphic screen editor), the response from all components (X and Z, dB/dt and B-Field) are examined. All significant responses from any of the components are inserted in the anomaly field. Since all anomaly edits are still being updated by the same routine, again only fitting the X-coil response from the dB/dt data, many of these “other” responses which have no measurable signature on the dB/dt X-coil data will only be flagged as an

anomaly location with no measurable response suitable for fitting to the reference nomogram. Although improperly represented, these “other” anomalies, at the very least, are identified by their location in the EM anomaly database (listing) and on the anomaly map.

Figures 1 and 2 provide examples of a typical display of the channel data used when reviewing the EM anomaly selection. Given the limited space available on a computer screen, a good display can include every even numbered channel, 8 to 30 (in more resistive areas, often all off-time channels can be displayed) for X and Z for both dB/dt and B-Field, along with the Hz monitor and the radar altimeter ( the EM primary field can also be very useful).

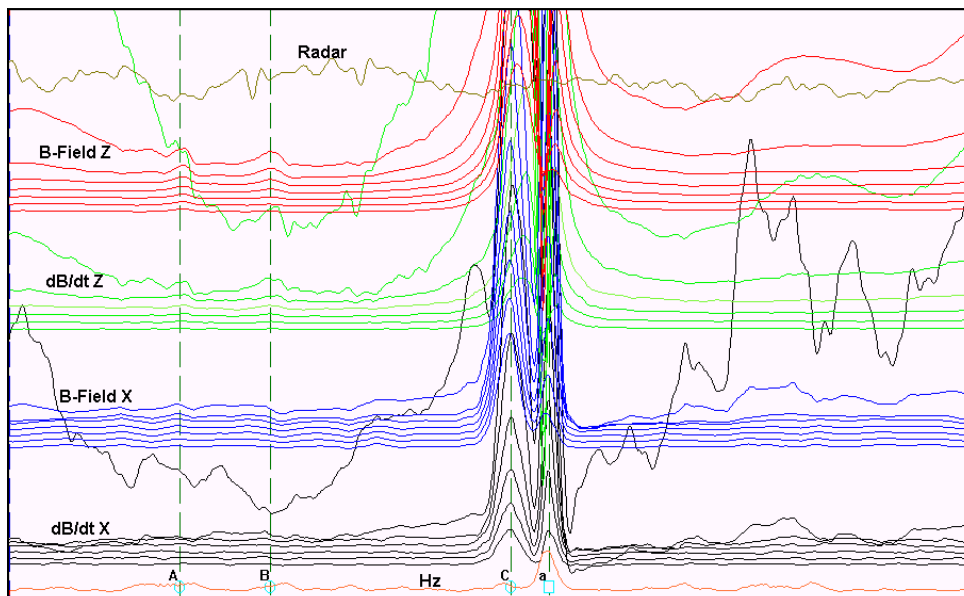


Figure 1

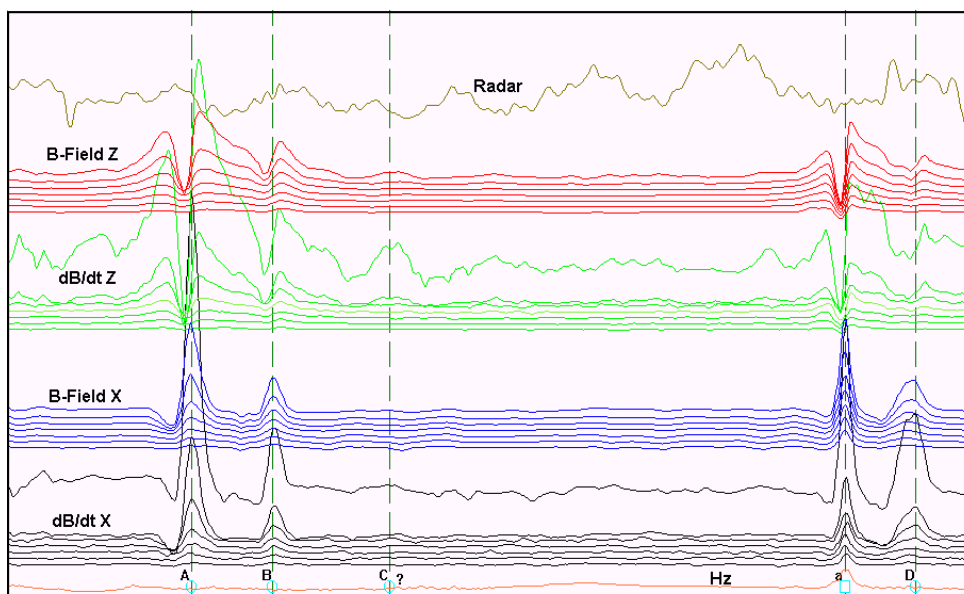


Figure 2

In **Figure 1**, anomalies identified as A and B could be indicative of a very deep source, where the coupling of the response is lost on the X-coil but persists on the Z-coil. The dB/dt X-coil response is devoid of any response while a weak response is visible of the Z-coil of the dB/dt response. Greater support for this selection is provided by the response on the B-Field. Although weak and very questionable on the X-coil (close to the noise level) the response is clearly marked on the Z-coil, as a low amplitude response of slow decay. This is a good indication of a high conductance body having a long time constant and therefore enhanced by the B-Field. These two anomalies may be prime targets for mineralization but will be indistinguishable from weak surface responses, as represented on the anomaly map or in the anomaly listing. Hence, the importance of always reviewing the EM anomaly selection against the data profiles.

In **Figure 2**, anomaly C? is similar to the responses discussed in Figure 1 above, in that it presents no measurable signature on the X-coil response (for both dB/dt and B-Field) but a weak response on the Z-coil response. The difference here however, is that the B-Field response does not show an enhancement of the response on the Z-coil, but an attenuation. This is indicative of a conductive response with a short time constant and hence more likely a weak surficial source.

The anomalies discussed in Figures 1 and 2 have very similar characteristics on dB/dt X and Z and B-Field X and yet may reflect very different conductive sources, one being of potentially economic interest. The only distinguishing signature is offered by the B-Field Z-coil response. Be aware that these differences are not properly accounted for in the current anomaly selection and presentation process.

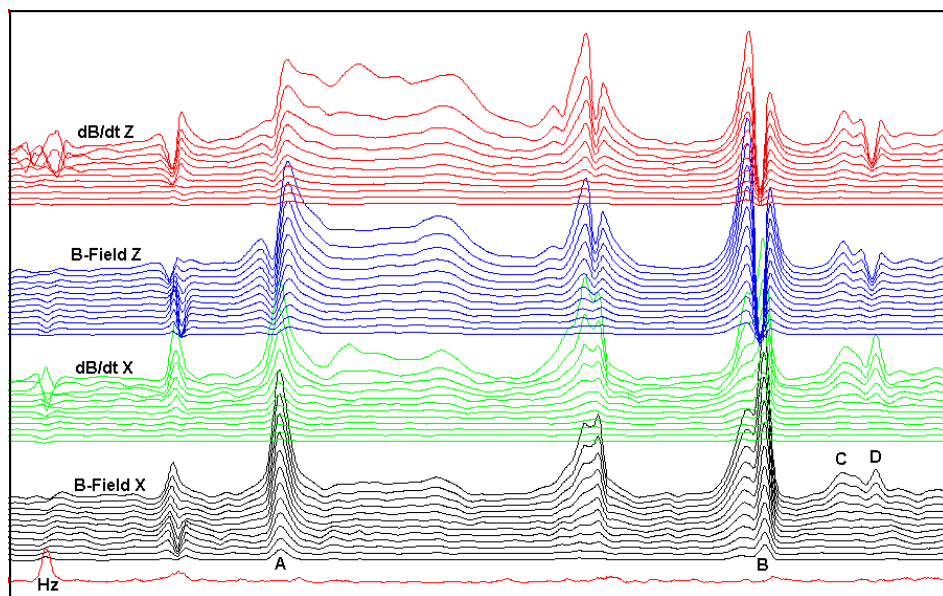


Figure 3

**Figure 3** shows the importance of looking at the response with all components when evaluating the

possible source of a conductive response. Anomalies identified as A and B look quite similar on the X-coil response and could both be interpreted as narrow, vertical conductors. However, looking at the response on the Z-coil clearly shows that anomaly A is the leading edge of a broad tabular body, better displayed on Z because of the enhanced coupling with the Z-coil, whereas anomaly B does reflect a narrow, near-vertical conductor. Anomalies C and D again look very similar on the response from the X-coil and could be interpreted as related to the same source. However, the response on the Z-coil indicates that the response at C is from a tabular or flat-lying source whereas the response at D is from a vertical source.



## **Appendix I**

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### **General Review of the Geophysical Response**

## Renabie Area

On March 15, 2013, Fugro Airborne Surveys conducted a GEOTEM<sup>®</sup> airborne electromagnetic and magnetic survey of Renabie on behalf of Conquest Resources Limited. Using Wawa, Ontario as the base of operations, a total of 175 line kilometres of data were collected over the survey area.

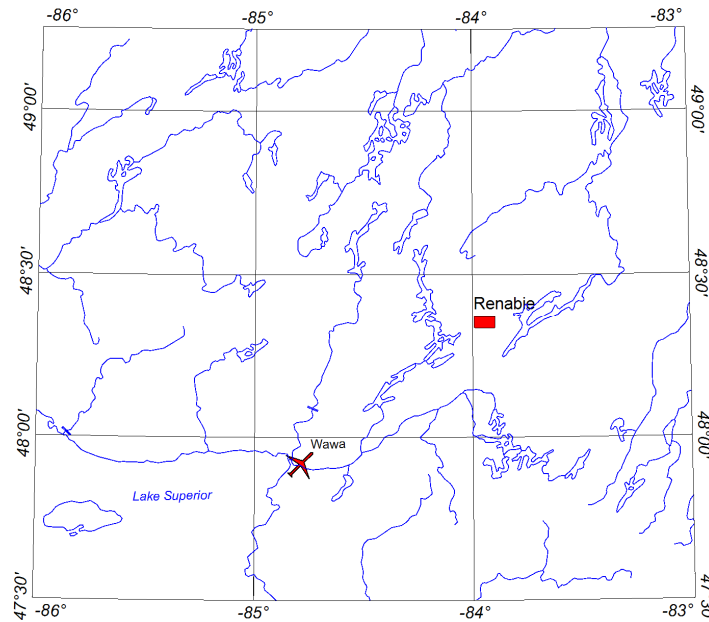
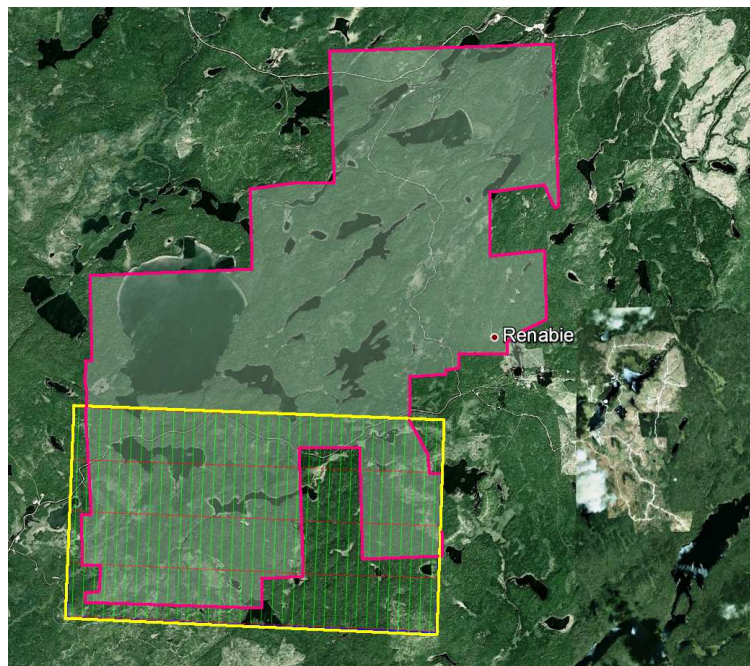


Fig. 1 Location of the survey area



Survey area & Claim blocks

Fig. 2

The image above (Fig. 2) shows the survey boundary (in yellow) and the claim boundaries (in magenta). The claims cover most of the survey area except a section in the SE corner of the block.

The next image (Fig. 3) shows the known mineral occurrences in the survey area (information obtained from the OGS database). These are mainly gold with minor copper and silver showings.

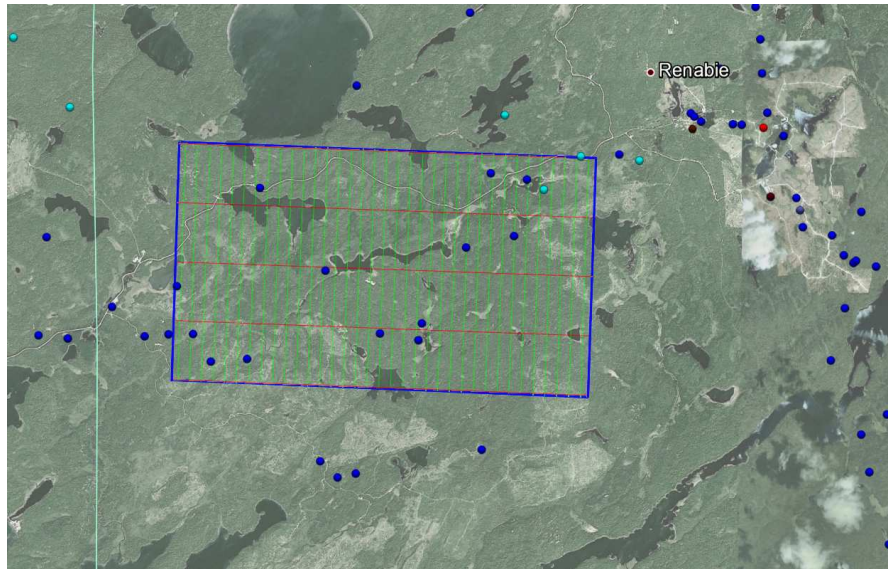


Fig. 3

The next image (Fig. 4) shows the general response obtained over the survey area from the airborne program.

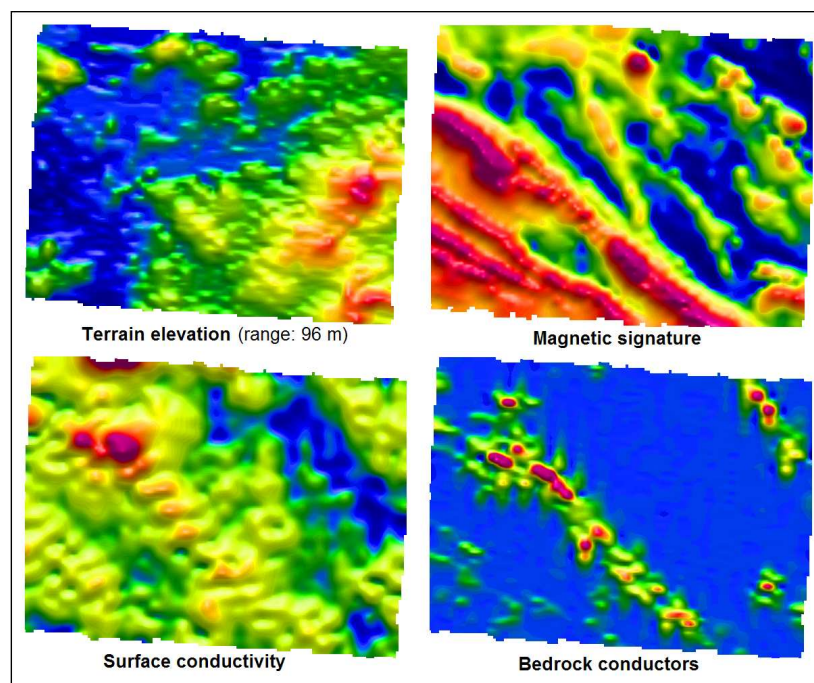


Fig 4.

## General review of the geophysical response over the survey area

The general geology, extracted from the OGS 1/250,000 bedrock geology map, shows a mix of generally NW – SE trending belts of mafic and felsic metavolcanics. This is generally supported by the magnetic signature although there appears to be a distinct division between more mafic material in the SW quadrant of the block with less mafic material to the NE. This division more or less coincides with a dominant NW – SE fault shown on the geology map (refer to Fig. 5).

The magnetic signature also suggests the presence of a small mafic intrusion along the north central edge of the block, which is not shown on the published geology map (Fig.6).

Numerous gold occurrences are shown to correlate with the location of either contacts between the mafic and more felsic metavolcanics or along NE – SW faults or shear zones, as inferred by the magnetic signature.

The surface conductivity shows localized increases in the NW section of the block. These correspond to Baltimore Lake and Rennie Lake and likely reflect weakly conductive lake bottom sediments. The remainder of the area is considered to be highly resistive (Fig 7.)

A number of bedrock conductors are noted which follow the trend of the metavolcanics and, more specifically lie along major faults between the mafic and felsic volcanics (Fig 8.). These do not particularly relate to the gold mineralization but could reflect either sulphide mineralization or graphitic material developed along the contacts or faults. The gold mineralization itself relates more to the faults and contacts in the volcanics and these features are well defined by the magnetic signature.

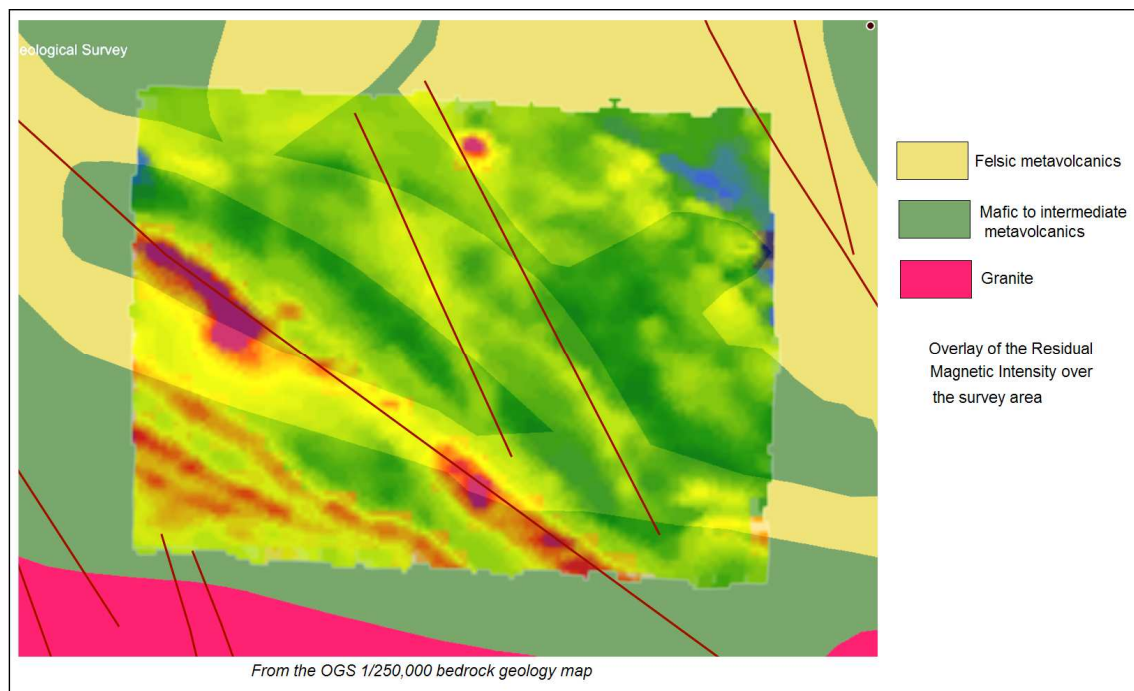


Fig. 5



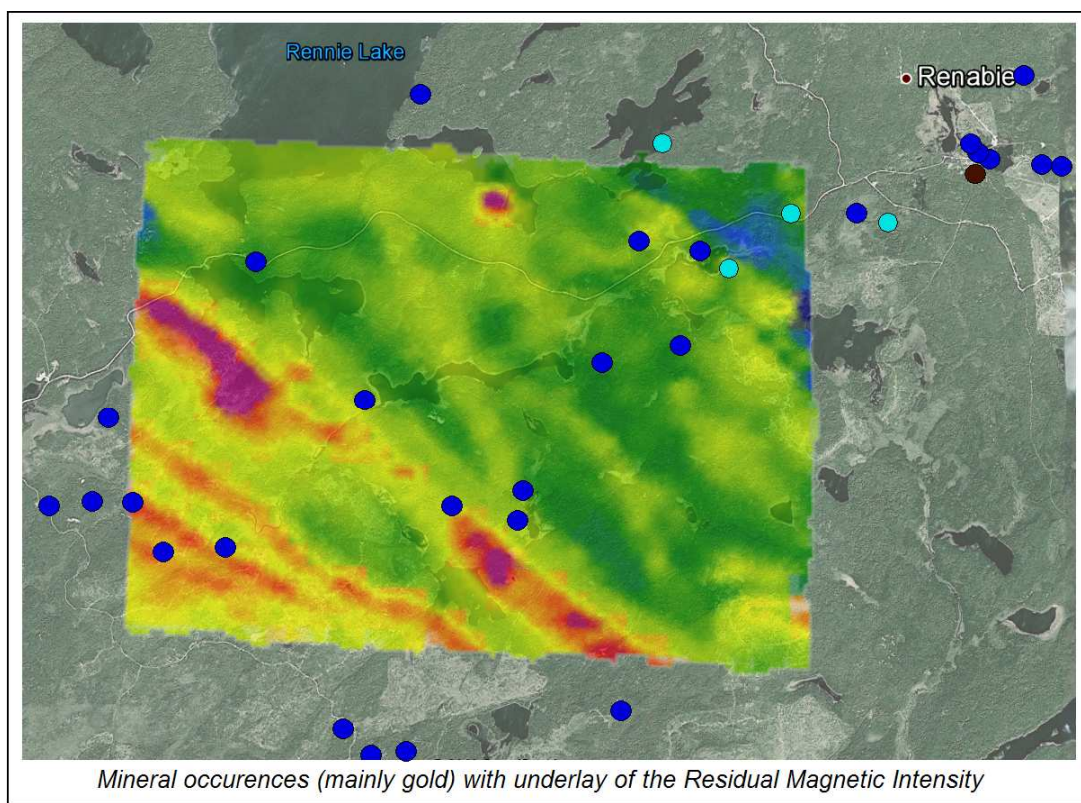


Fig. 6

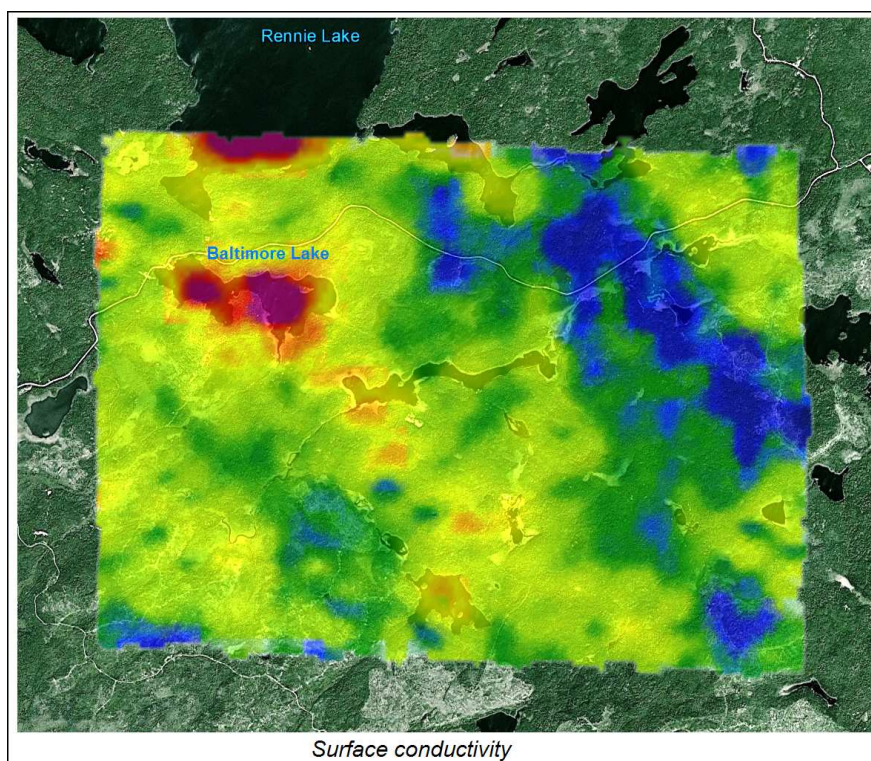


Fig. 7

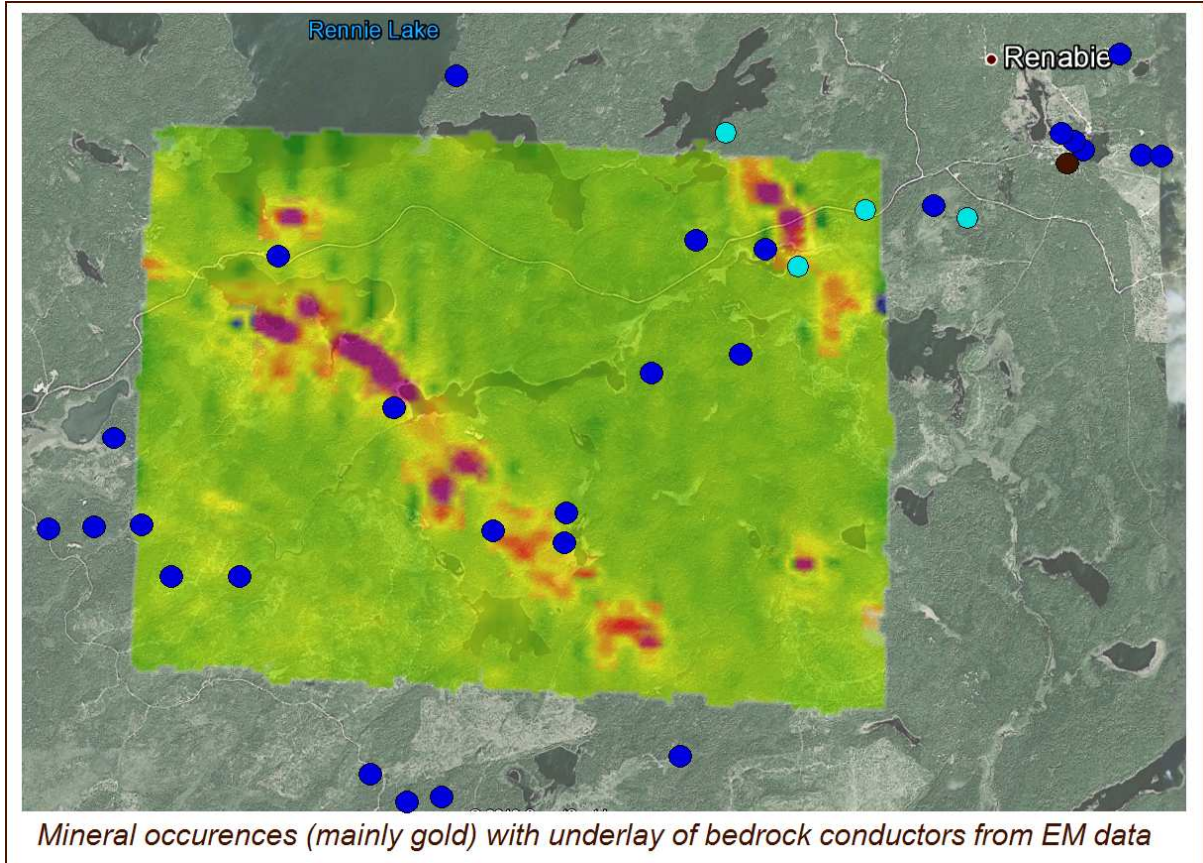


Fig. 8

The last figure (9) shows the GEOTEM electromagnetic profile response over a section of the main bedrock conductive trend which follows the contact between the mafic and felsic metavolcanics and which may also be associated with a major fault zone. This conductor appears to be of good conductance as reflected by the general character of the anomaly (moderate to low amplitude and slow decay). The correlation between the X-coil and Z-coil response (single peak on X and double peaks on Z) suggests a narrow sheet-like body of vertical extent. These characteristics are often associated with mineralization in the bedrock.



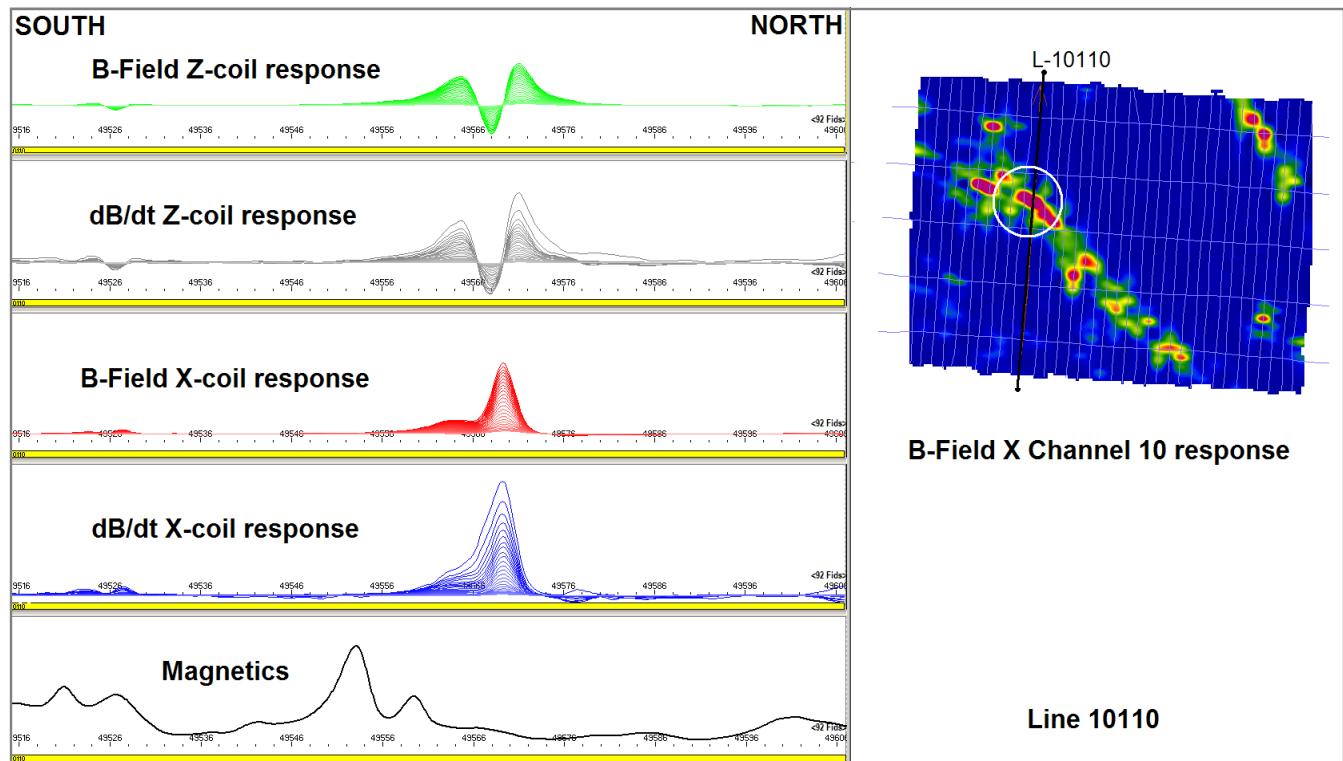


Fig. 9

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**Fugro Airborne Surveys**  
*Interpretation & Consulting Services*  
 Ottawa, CANADA  
 March 27<sup>th</sup>, 2013